

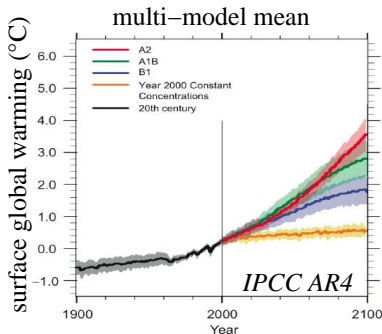
The added value of tropospheric water vapor isotopic measurements for evaluation cloud and precipitation processes in climate models

Camille Risi

LMD/IPSL/CNRS

Seminar at ITP-CAS, April 2013

Inter-model spread in climate projections



CO₂ increase

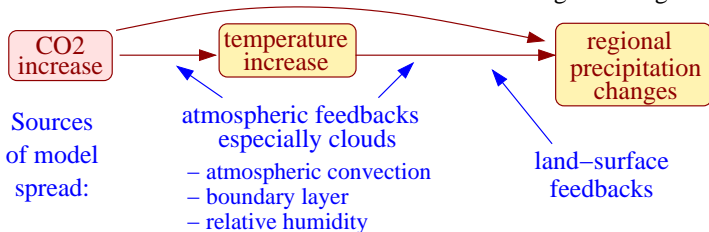
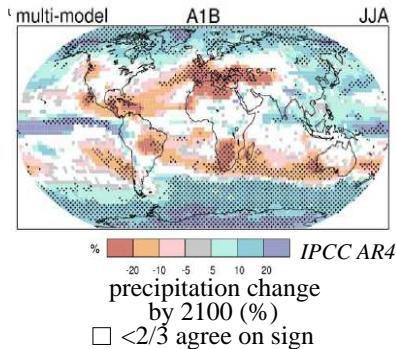
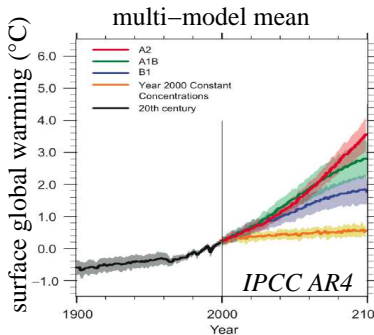
temperature increase

Sources of model spread:

atmospheric feedbacks especially clouds

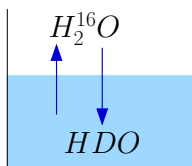
- atmospheric convection
- boundary layer
- relative humidity

Inter-model spread in climate projections



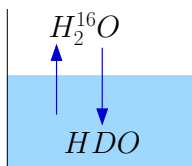
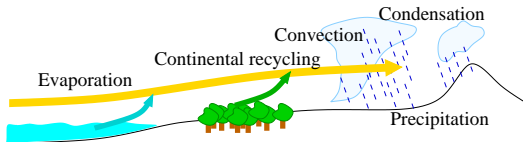
Water isotopic composition

- ▶ $H_2^{16}O$, HDO , $H_2^{18}O$, $H_2^{17}O$, fractionation



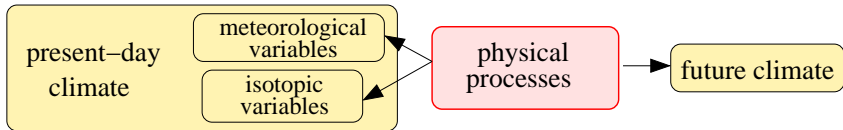
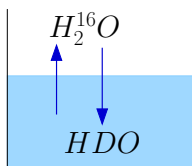
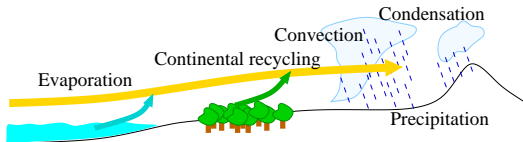
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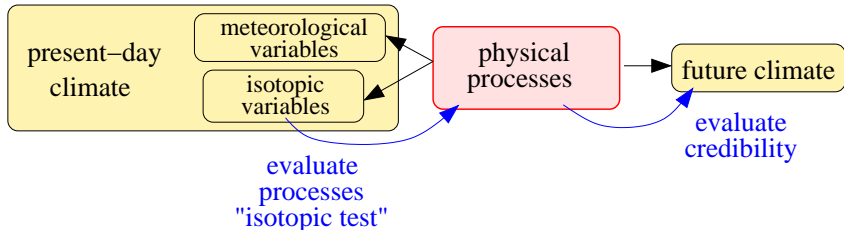
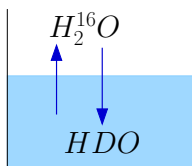
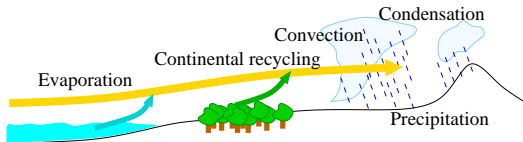
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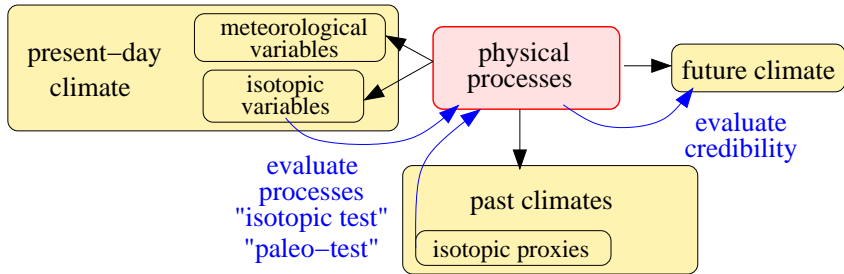
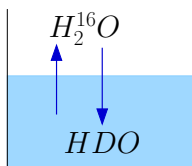
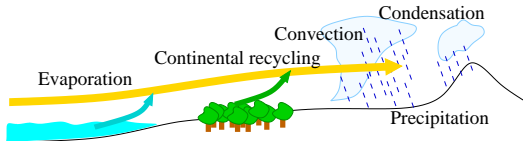
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Overview of my activities

1. evaluation of atmospheric processes
 - ▶ processes controlling humidity
 - ▶ atmospheric deep convection

Overview of my activities

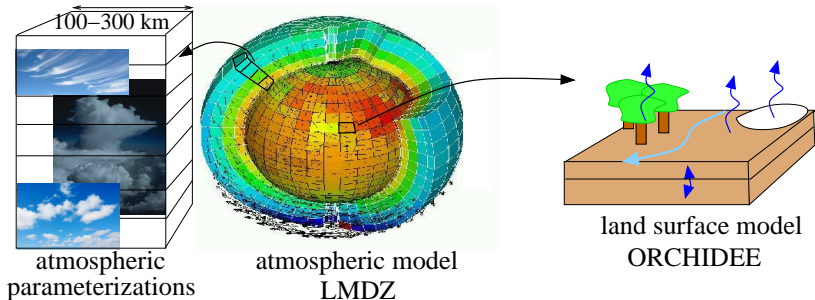
1. evaluation of atmospheric processes
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 - ▶ partitioning of water fluxes at land surface
 - ▶ land-atmosphere feedbacks, continental recycling

Overview of my activities

1. evaluation of atmospheric processes
 - ▶ processes controlling humidity
 - ▶ atmospheric deep convection
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 - ▶ partitioning of water fluxes at land surface
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3. evaluation of tropical precipitation changes
 - ▶ what do tropical water isotopic proxies record
 - ▶ link between past and future behavior (CMIP5)

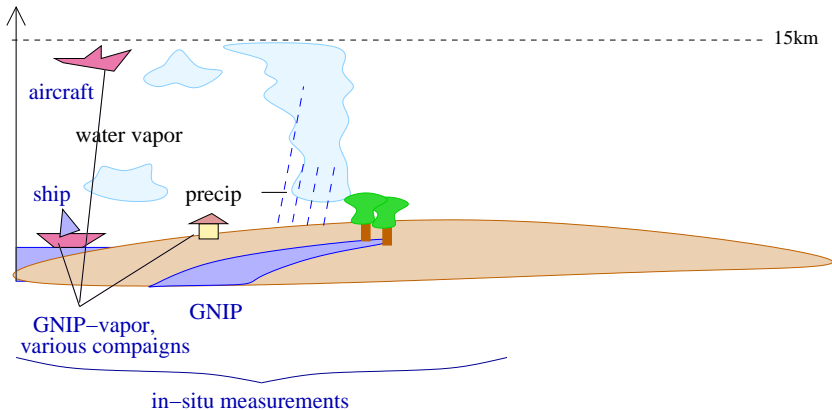
LMDZ and ORCHIDEE models

- ▶ components of IPSL climate model

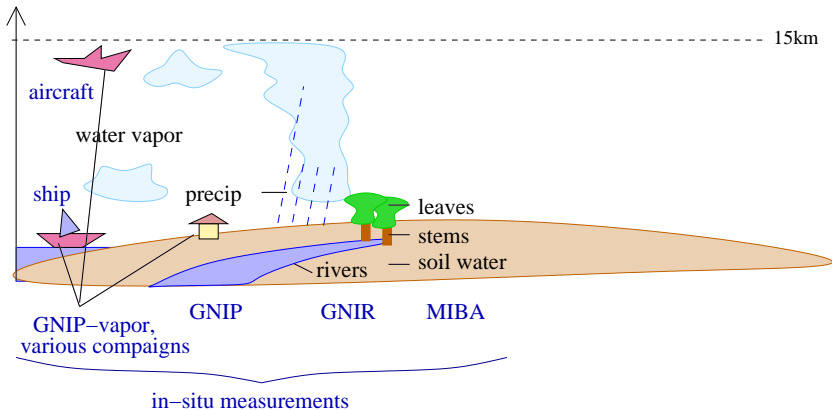


- ▶ isotope-enabled (*Risi et al 2010a*) + water tagging
- ▶ nudging capability \Rightarrow realistic dynamical context
- ▶ zoom capability down to 30km
- ▶ isotope-enabled + water tagging

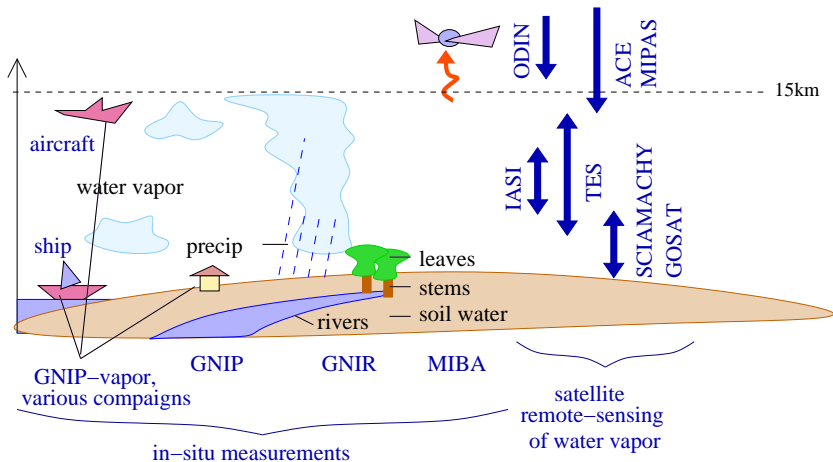
Available measurements



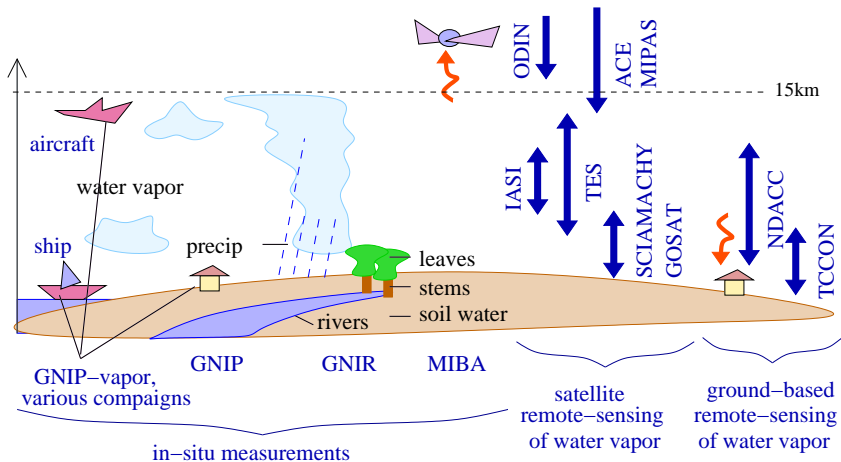
Available measurements



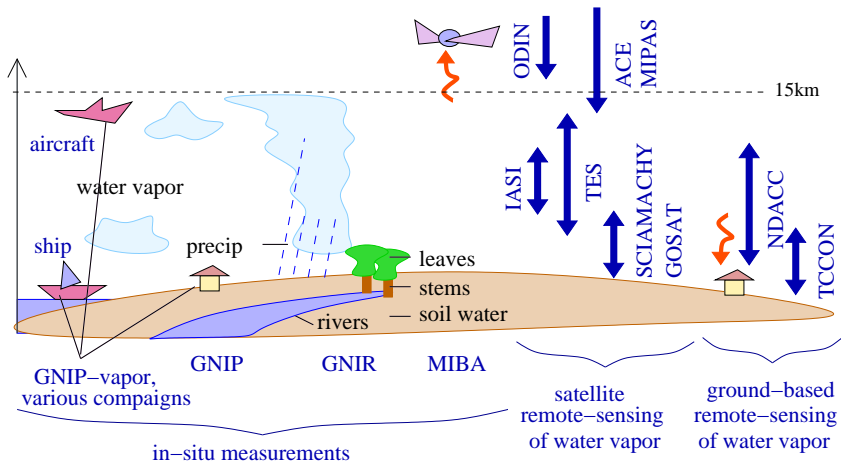
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Available measurements



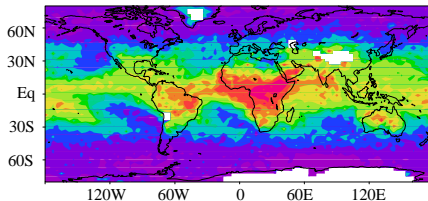
Available measurements



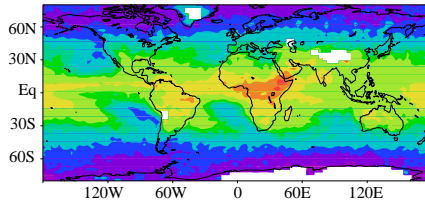
- ▶ for remote-sensing : focus on spatio-temporal variations
- ▶ account for sampling and instrument sensitivity

Evaluation of LMDZ water vapor and precip

TES data

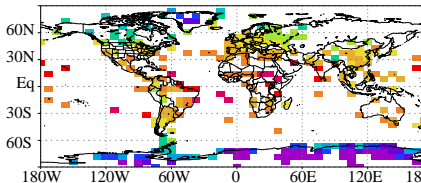


LMDZ

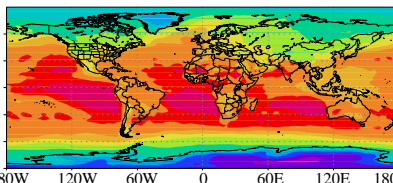


δD_{vapor} (‰) 800hPa (anomaly relatively to the tropical average)

GNIP data

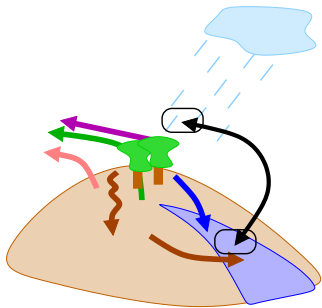


LMDZ

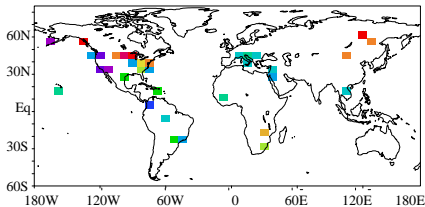


$\delta^{18}O_{precip}$ (‰)

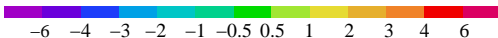
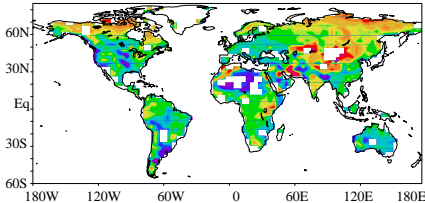
Evaluation of LMDZ-ORCHIDEE precipitation and rivers



GNIR and GNIP data



LMDZ-ORCHIDEE-iso



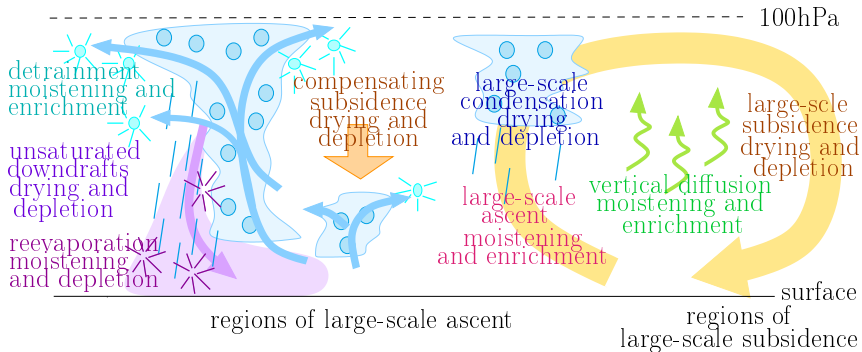
$$\delta^{18}O_{river} - \delta^{18}O_{precip} (\text{‰})$$

1) Using water vapor measurements to evaluate atmospheric processes

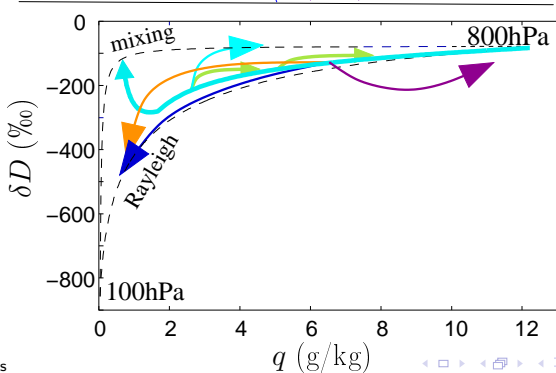
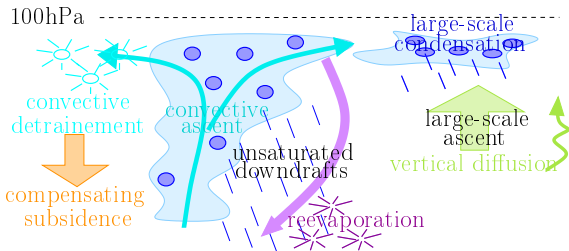
- ▶ what controls the water vapor composition
- ▶ 2 examples

Processes controlling isotopic composition

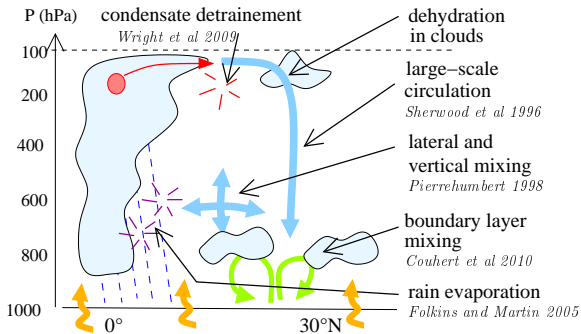
- ▶ observational studies (Risi et al 2008b), in particular at intra-event time scales (Risi et al 2010c, Tremoy et al 2012)
- ▶ modeling studies (Risi et al 2008, 2010b, 2012b)



q - δD complementarity



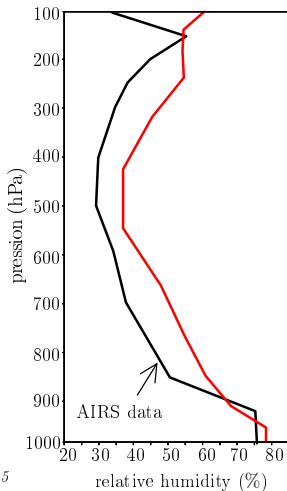
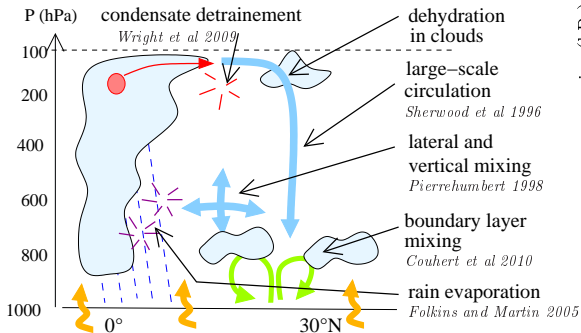
1) Processes controlling humidity



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LMDZ-iso (Risi et al 2010a):

— control: AR4 version (19 levels)

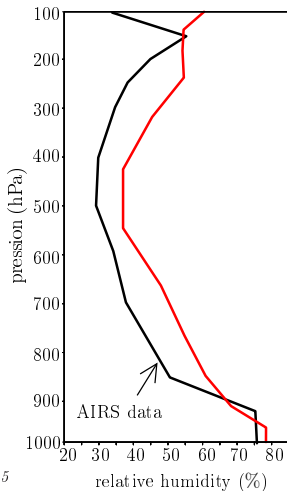
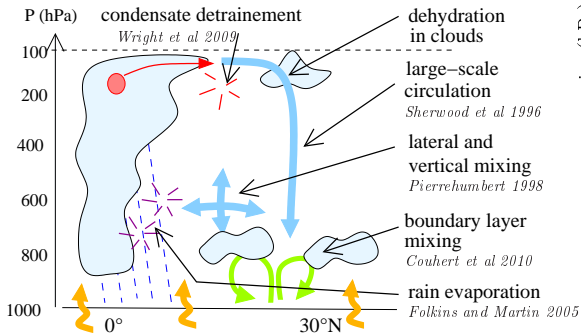


1) Processes controlling humidity

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3 reasons
for a
moist bias



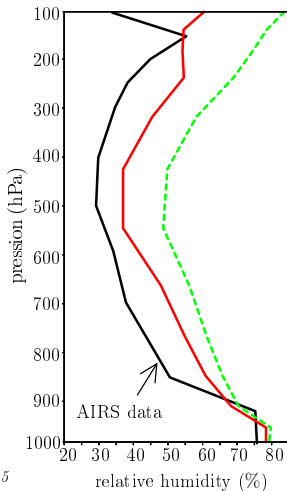
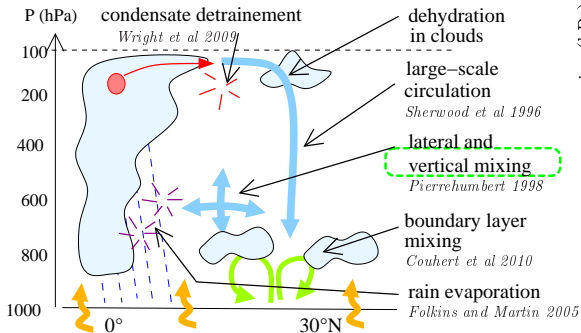
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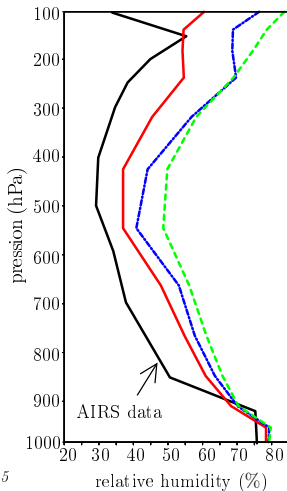
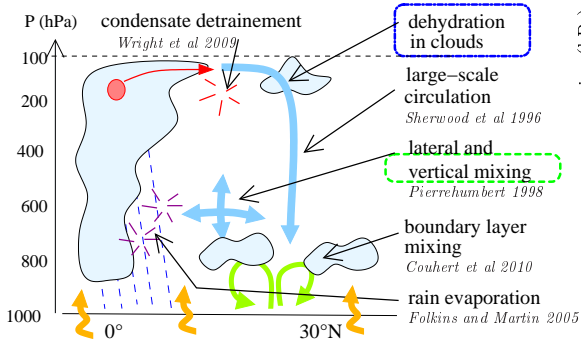


1) Processes controlling humidity

LMDZ-iso (Risi et al 2010a):

- control: AR4 version (19 levels)
- - - diffusive vertical advection
- $\sigma_q/10$

3 reasons
for a
moist bias

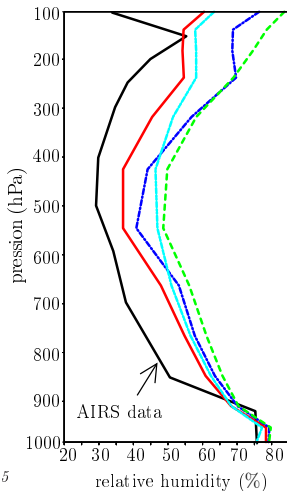
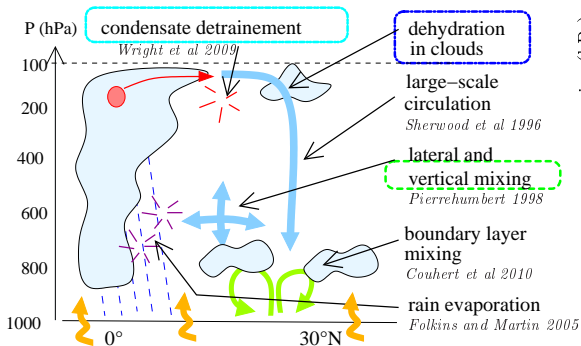


1) Processes controlling humidity

LMDZ-iso (Risi et al 2010):

- control: AR4 version (19 levels)
- - - diffusive vertical advection
- $\sigma_q/10$
- $\epsilon_p/2$

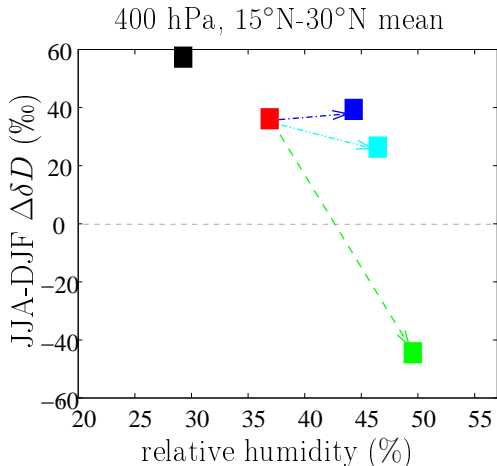
3 reasons
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What causes the moist biases in GCMs?

Sensitivity tests:
with LMDZ:

- Control
- Excessively diffusive vertical advection
- Excessive condensate detrainment
- Insufficient in-situ condensation
- AIRS/ACE data



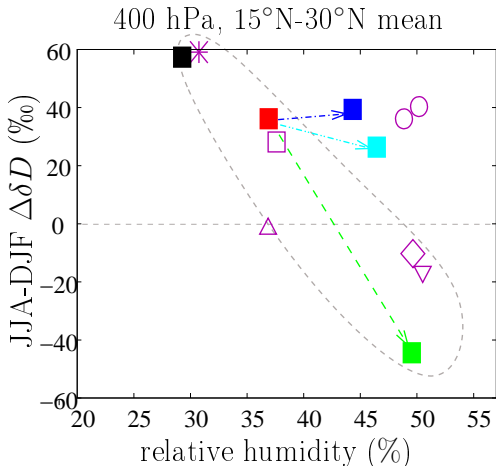
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SWING2 models:

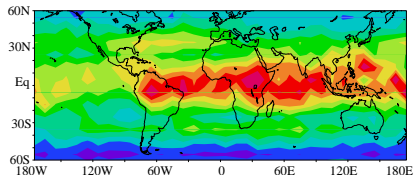
- | | |
|---------|--------|
| □ ECHAM | ◇ CAM2 |
| △ MIROC | ○ GISS |
| * HadAM | ▽ GSM |



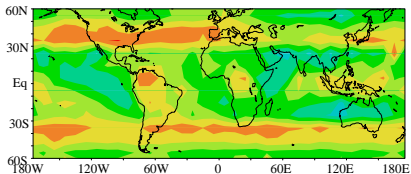
► frequent reason for moist bias=excessively diffusive advection

2) Upper tropospheric convective moistening

MIPAS data at 200hPa, annual



LMDZ control

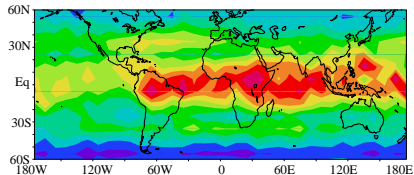


-700 -640 -600 -560 -520 -480 -440 -400 -360 -320

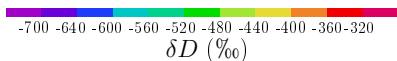
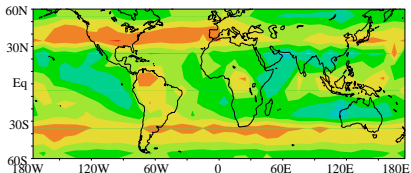
δD (‰)

2) Upper tropospheric convective moistening

MIPAS data at 200hPa, annual

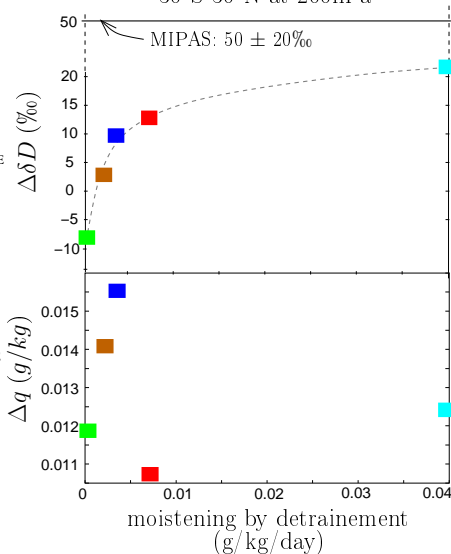


LMDZ control



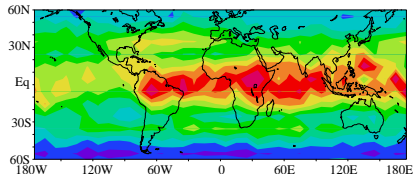
- control
- vertical advection more diffusive
- stronger condensate detrainment
- less in-situ condensation
- less in-situ precipitation

Difference 15°S-15°N minus
30°S-30°N at 200hPa

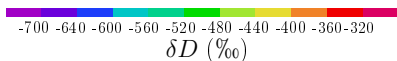
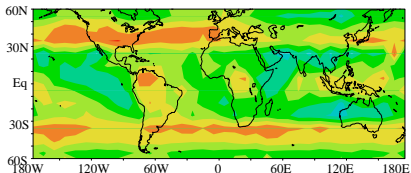


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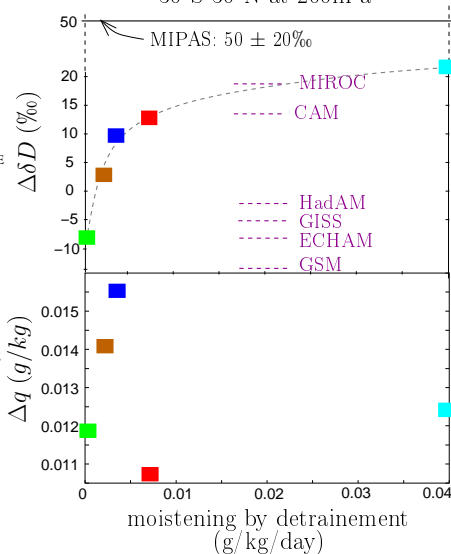


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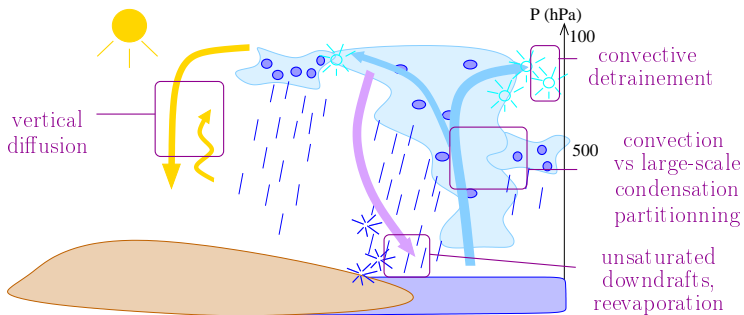
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Difference 15°S-15°N minus 30°S-30°N at 200hPa



Conclusion on atmospheric processes

- ▶ Potential of isotopic measurements to evaluate a broad range of processes in atmospheric models



Perspectives on atmospheric processes

- ▶ Combine q , δD + cloud \Rightarrow better constrain large-scale precip

Perspectives on atmospheric processes

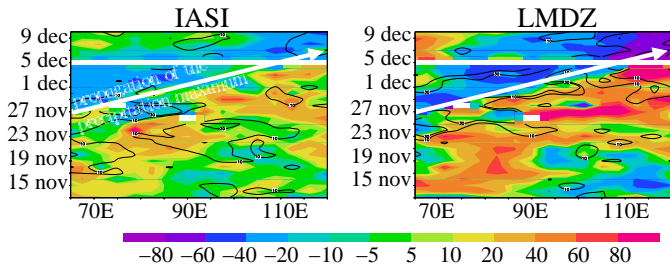
- ▶ Combine q , δD + cloud \Rightarrow better constrain large-scale precip
- ▶ Combine q , δD + chemical tracers : CO , O_3 , ^{10}Be \Rightarrow better characterize fluxes

Perspectives on atmospheric processes

- ▶ Combine q , δD + cloud \Rightarrow better constrain large-scale precip
- ▶ Combine q , δD + chemical tracers : CO, O₃, ¹⁰Be \Rightarrow better characterize fluxes
- ▶ MJO project : cause of models' difficulties? \Rightarrow Relate MJO biases to specific problems in parameterizations, isotopes as additional diagnostic.

Perspectives on atmospheric processes

- ▶ Combine q , δD + cloud \Rightarrow better constrain large-scale precip
- ▶ Combine q , δD + chemical tracers : CO, O₃, ¹⁰Be \Rightarrow better characterize fluxes
- ▶ MJO project : cause of models' difficulties? \Rightarrow Relate MJO biases to specific problems in parameterizations, isotopes as additional diagnostic.
- ▶ IASI data : daily global coverage \Rightarrow convective organization, life cycle

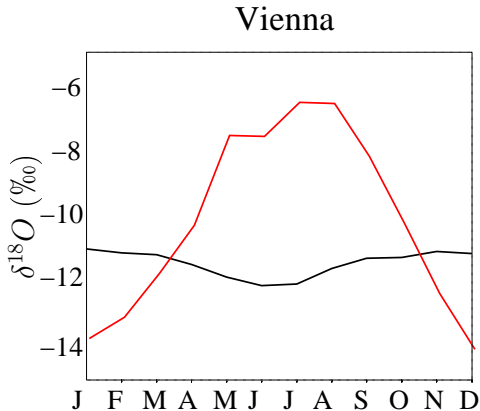
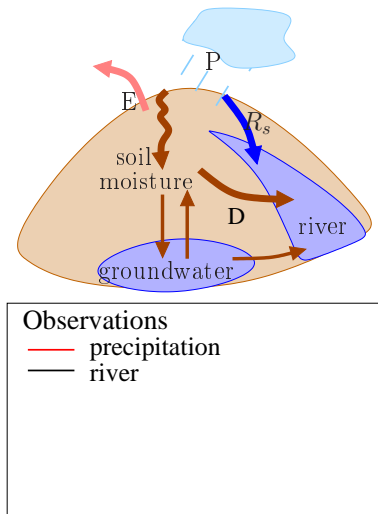


colors: δD anomaly (‰)
contours: q anomalies (‰)

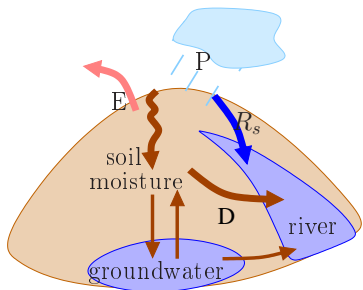
II) Using river water and water vapor measurements to evaluate land surface processes

- ▶ 2 examples

1) Pathways from precipitation to rivers



1) Pathways from precipitation to rivers



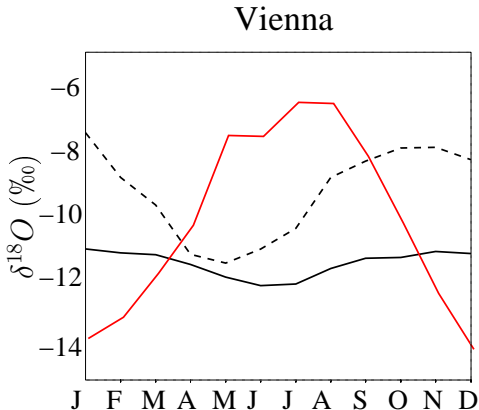
Observations

— precipitation

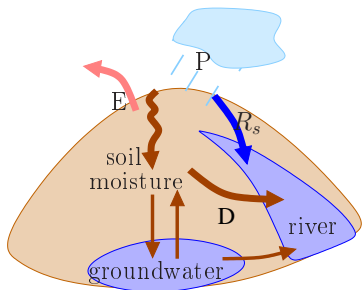
— river

LMDZ-ORCHIDEE-iso

---- control



1) Pathways from precipitation to rivers

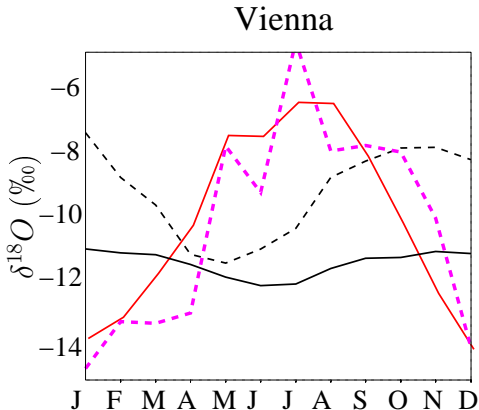


Observations

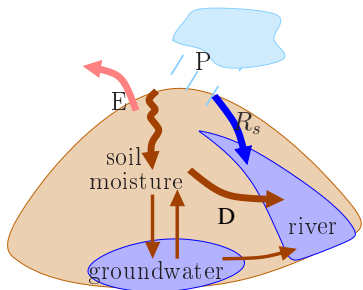
- precipitation
- river

LMDZ-ORCHIDEE-iso

- - - control
- · - · all surface runoff



1) Pathways from precipitation to rivers

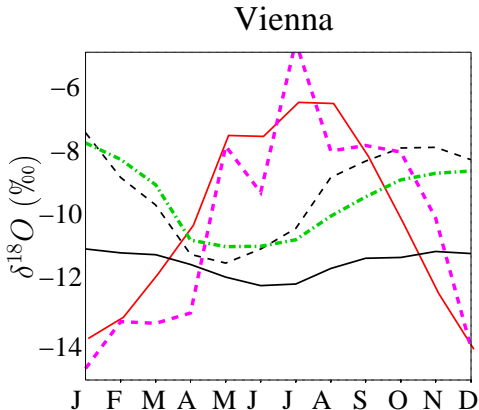


Observations

- precipitation
- river

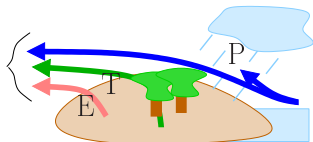
LMDZ-ORCHIDEE-iso

- - - control
- · - · all surface runoff
- · · · slower underground reservoirs



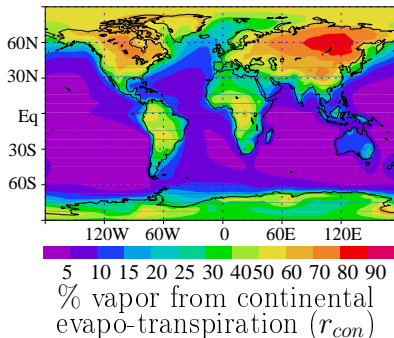
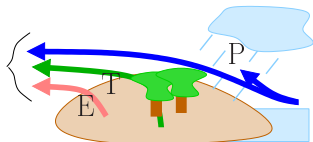
2) Continental recycling

Water tagging:



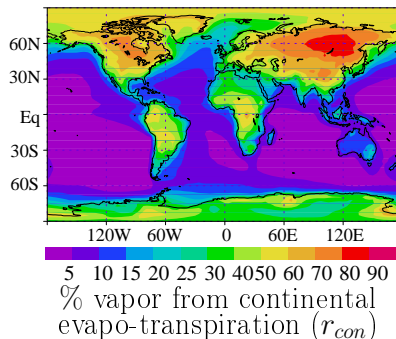
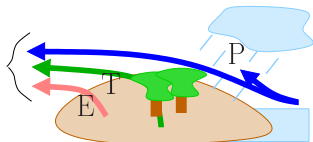
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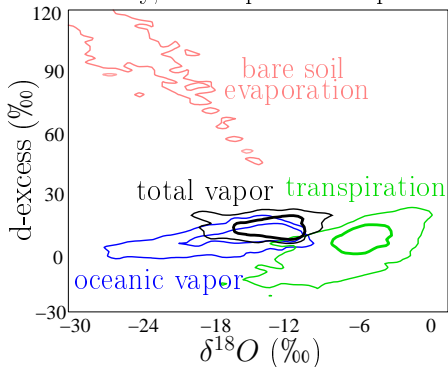


2) Continental recycling

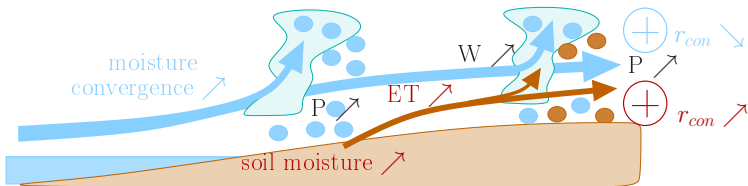
Water tagging:



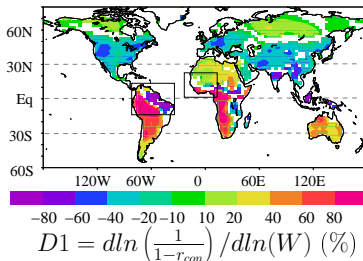
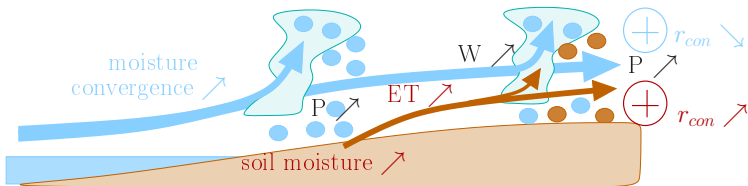
PDF of vapor composition
monthly, all tropical land points



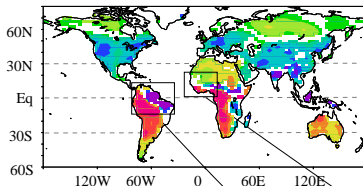
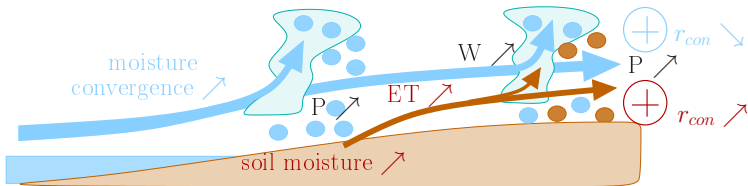
Continental recycling feedbacks



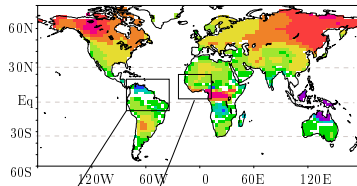
Continental recycling feedbacks



Continental recycling feedbacks



$$D1 = d \ln \left(\frac{1}{1-r_{con}} \right) / d \ln(W) \quad (\%)$$

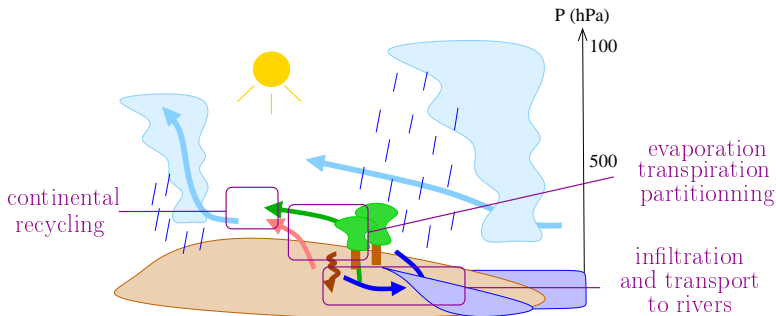


$$D1_{iso} = d \delta D / d \ln(W) \quad (\% / \%)$$

$r=0.88$ $r=0.96$

- ▶ use $D1_{iso}$ to evaluate role of cont recycling

Conclusion on land surface processes

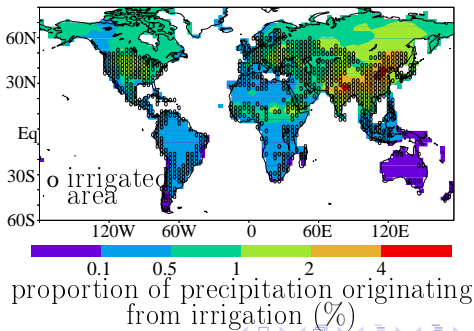
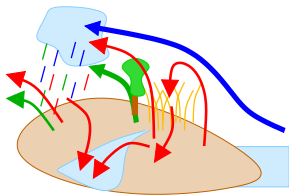


Perspectives on land surface

- ▶ isotopes in 11-layer hydrology of ORCHIDEE \Rightarrow better simulation of soil profiles, more physical runoff-drainage partitioning
- ▶ use d-excess signal in the vapor to constrain evaporation/transpiration partitioning?
- ▶ link between present-day representation of the water cycle and simulated hydrological response to climate changes

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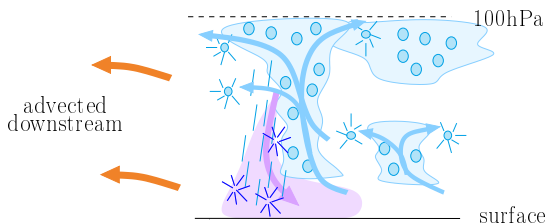


III) What does tropical $\delta^{18}O_p$ record ?

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- ▶ Interpretations in the literature :
 - ▶ Thompson et al 2000 → proxy for temperature
 - ▶ Vuille et al 2005, Pausata et al 2011 → proxy for monsoon intensity, upstream precipitation



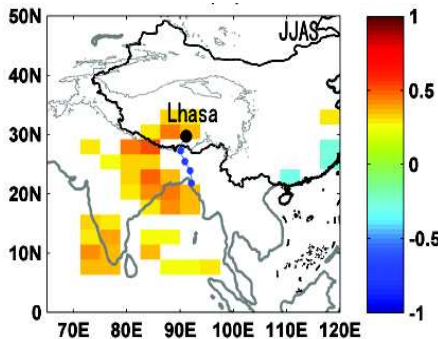
⇒ process studies at the daily time scale (*Gao et al 2011, submitted*) using precip data

But missing link : vapor ?

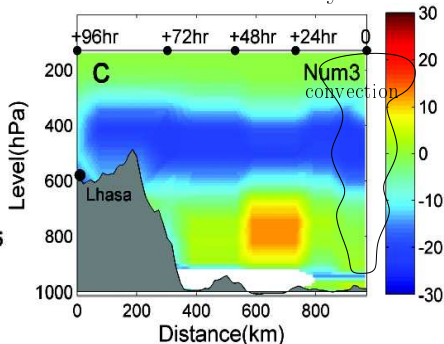
Process study using satellite observations

- ▶ He You's PhD thesis : use of TES data

daily correlation between δD at 400hPa at Lhasa and OLR around in JJAS



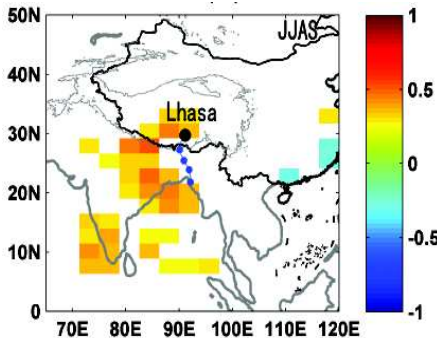
δD composite anomalies (‰) along trajectories to Lhasa when OLR is low 4 days before



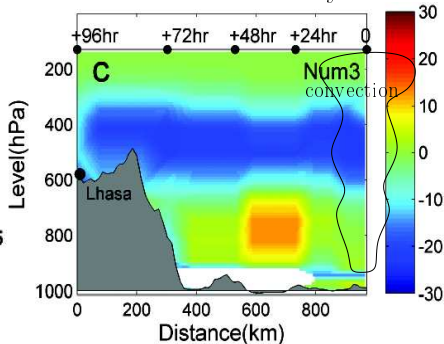
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δD composite anomalies (‰) along trajectories to Lhasa when OLR is low 4 days before



- ▶ convection in India depletes mid-tropospheric vapor
- ▶ depleted anomaly is transported downstream to Lhasa

Does this apply to paleo scales ?

Can we use present-day observations to better understand processes controlling paleo $\delta^{18}O$ and evaluate them in models ?
⇒ modelling study

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 - ▶ paleo relationships between $\delta^{18}O$ and climate ?

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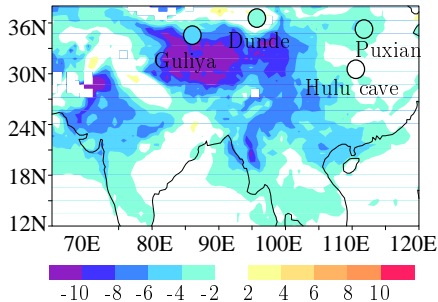
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- ▶ sensitivity tests to model physics and resolution (including 50km zoom)
 - ▶ robustness of simulated relationships ?

Evaluation for LGM and MH

Last Glacial Maximum

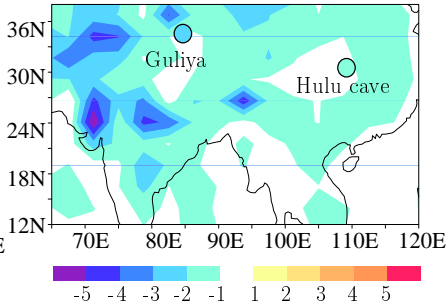


$\Delta\delta^{18}O$ (‰)

(sea water corrected)

LMDZ 50km resolution

Mid Holocene

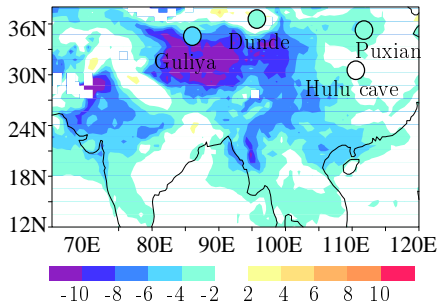


$\Delta\delta^{18}O$ (‰)

LMDZ control

Evaluation for LGM and MH

Last Glacial Maximum

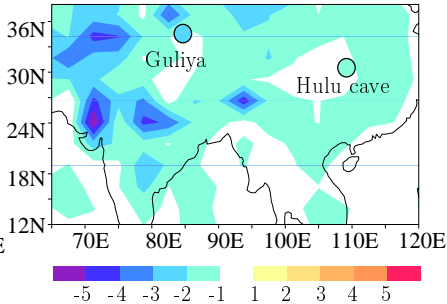


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$\Delta\delta^{18}O$ (‰)

LMDZ control

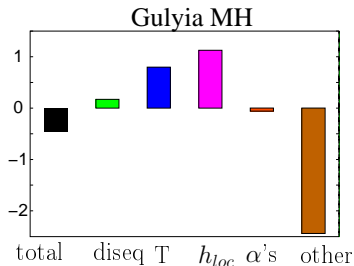
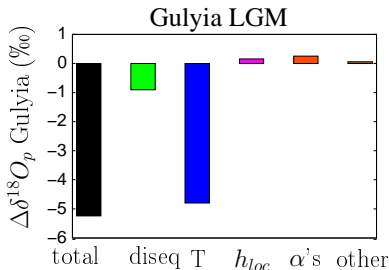
- ▶ LMDZ captures LGM and MH observed depletion

Causes of $\delta^{18}O$ changes?

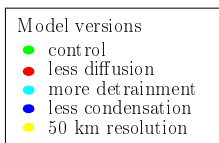
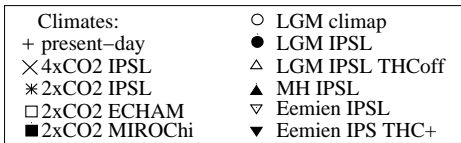
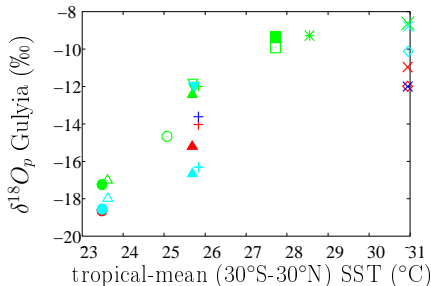
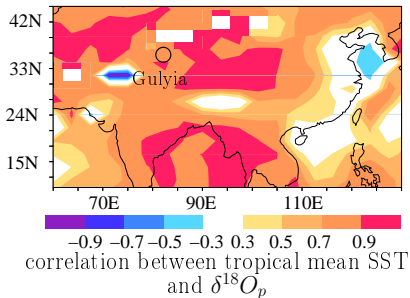
$$R_p = R_v + \underbrace{(R_p - \alpha_{loc} \cdot R_v)}_{\text{vap-cond diseq}} + (\alpha_{loc} \cdot R_v - R_v)$$

$$R_v = \underbrace{\frac{R_{occ} \cdot \alpha_i}{\alpha_K \cdot (1 - h_i) + h_i}}_{\text{initial vapor}} \cdot \underbrace{\left(\frac{h_{loc} \cdot q_s(T_{loc})}{q_s(T_i)} \right)^{\alpha_{loc} - 1}}_{\text{distillation}} + \text{residual}$$

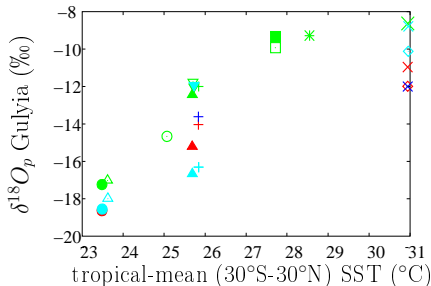
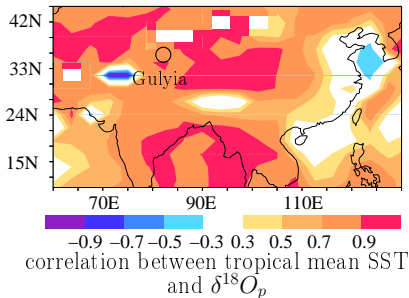
e.g. upstream convection



Is $\delta^{18}O$ a proxy for temperature?



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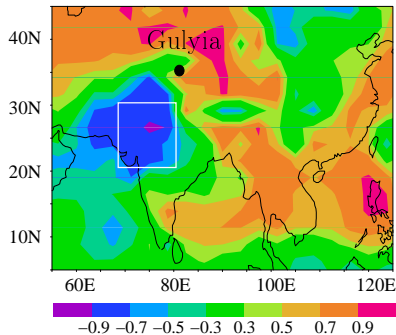


Climates:	
+ present-day	○ LGM climap
× 4xCO ₂ IPSL	● LGM IPSL
* 2xCO ₂ IPSL	△ LGM IPSL THCOff
□ 2xCO ₂ ECHAM	▲ MH IPSL
■ 2xCO ₂ MIROCi	▽ Eemien IPSL
	▼ Eemien IPSL THC+

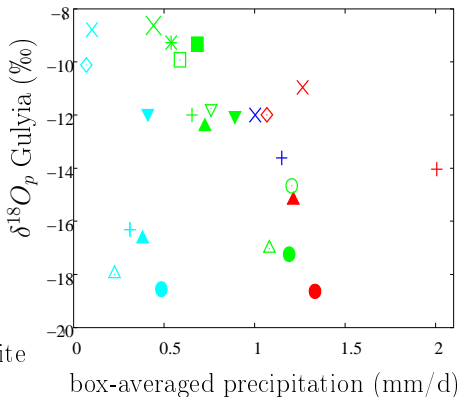
Model versions	
●	control
●	less diffusion
●	more detrainment
●	less condensation
●	50 km resolution

- ▶ temperature = significant control at paleo time scales
- ▶ robust to model physics

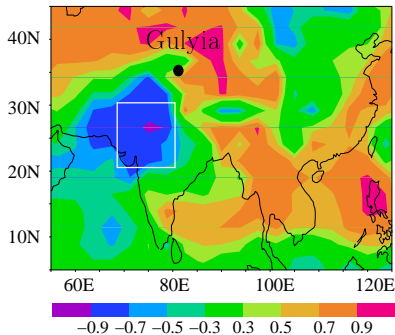
Is $\delta^{18}O$ a proxy for precipitation ?



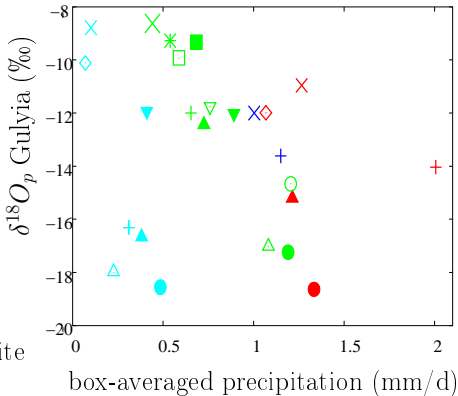
correlation between $\delta^{18}O_p$ at site
and precipitation around



Is $\delta^{18}O$ a proxy for precipitation ?



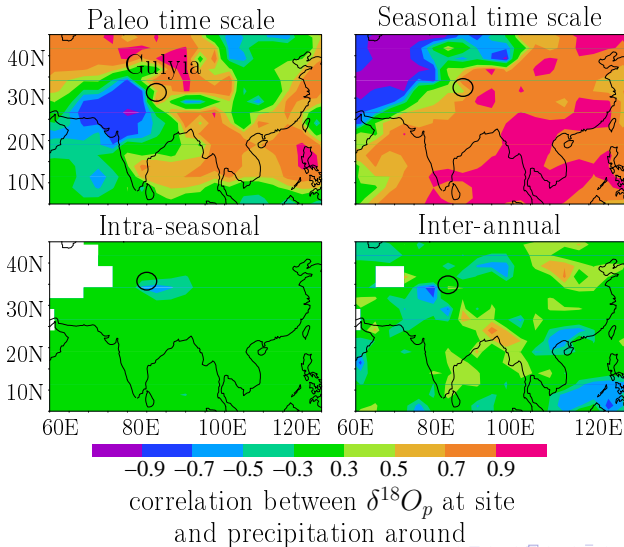
correlation between $\delta^{18}O_p$ at site
and precipitation around



- ▶ Upstream precipitation plays a role at paleo time scales
- ▶ Sensitive to the model physics

$\delta^{18}O$ controls accross time scales

- ▶ Is understanding daily controls enough to understand paleo controls ?



Conclusion on paleo

- ▶ LMDZ can reproduce several aspects of past $\delta^{18}O$ changes
- ▶ At paleo time-scales and especially during LGM, temperature is a major control
- ▶ At paleo time-scales and especially during MH, relationship with upstream precip but sensitive to the model physics
- ▶ Surface and satellite data can help understand processes controlling $\delta^{18}O$ at daily time scale \Rightarrow role of convection
- ▶ But relationship between precip and $\delta^{18}O$ depends on time scale

Perspectives on paleo

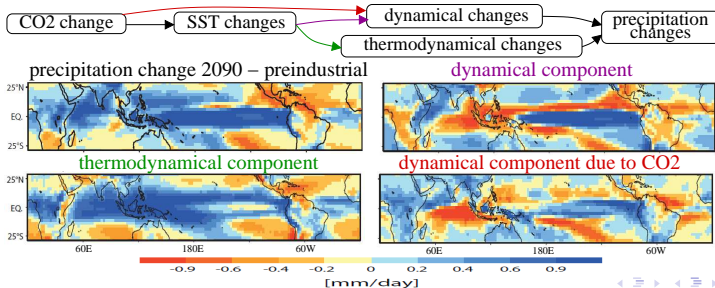
- ▶ Better evaluate climate- $\delta^{18}O$ relationships :
 - ▶ more data synthesis needed for paleo $\delta^{18}O$ to evaluate models
 - ▶ are some sensitivity tests more realistic at daily time scales?
 - ▶ do we expect them to be more realistic for paleo time scales?
 - ▶ compare with other models

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 - ▶ use $\delta^{18}O$ records to better constrain future precip changes?
 - ▶ common behavior in past/ future? common mechanisms
- ⇒ investigation using past and future simulations in CMIP5.

Perspectives on paleo

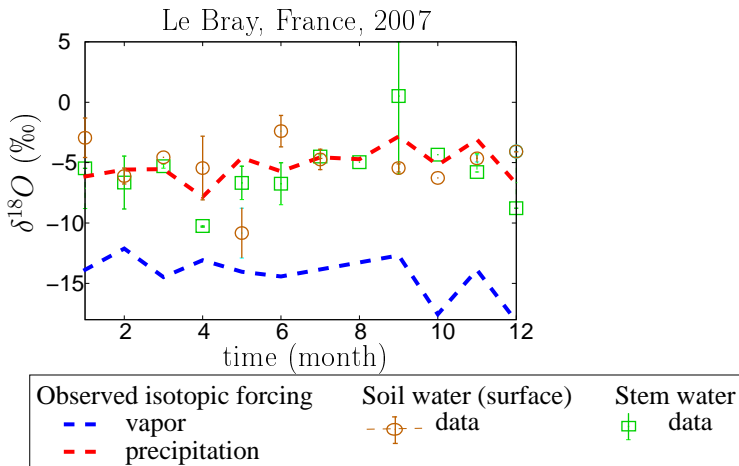
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Bony et al 2013 : decomposition of future precip changes :



Appendix

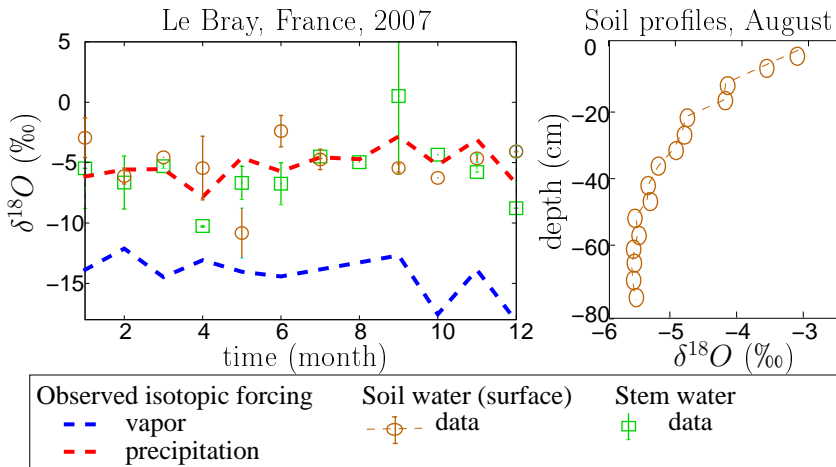
Evaluation of ORCHIDEE land surface isotopes

- ▶ Le Bray (France, *Wingate et al 2009*)



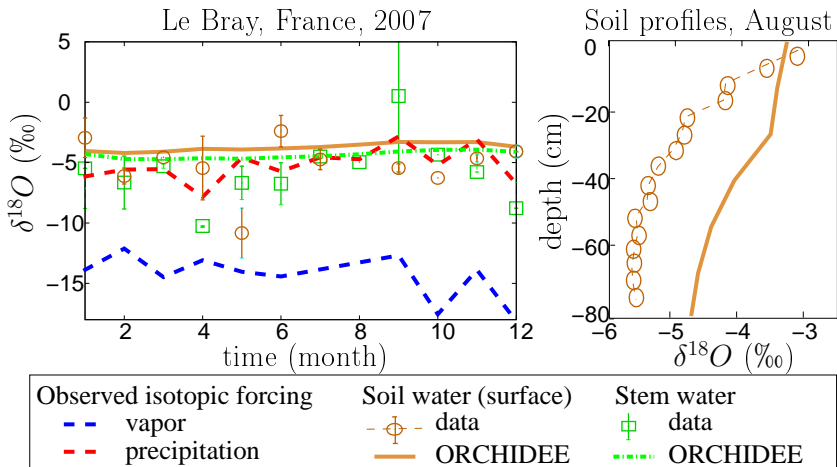
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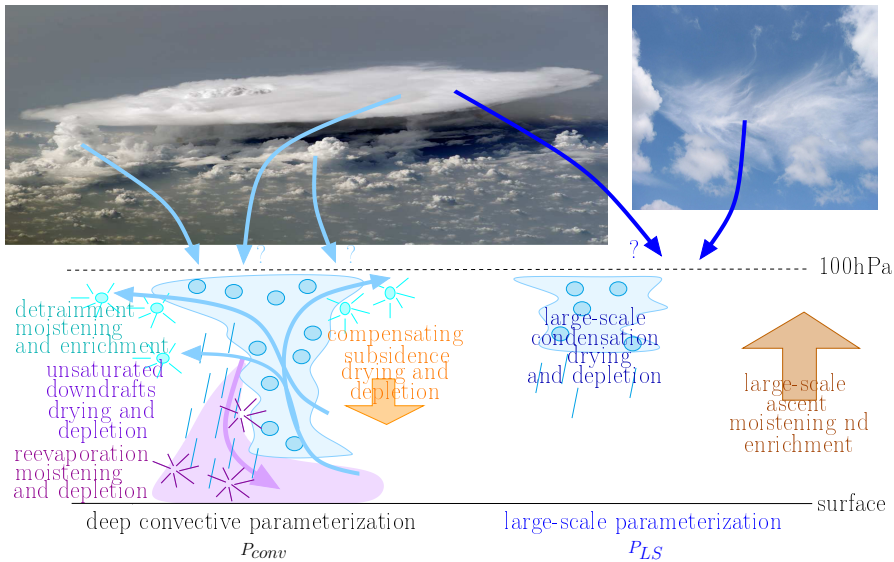


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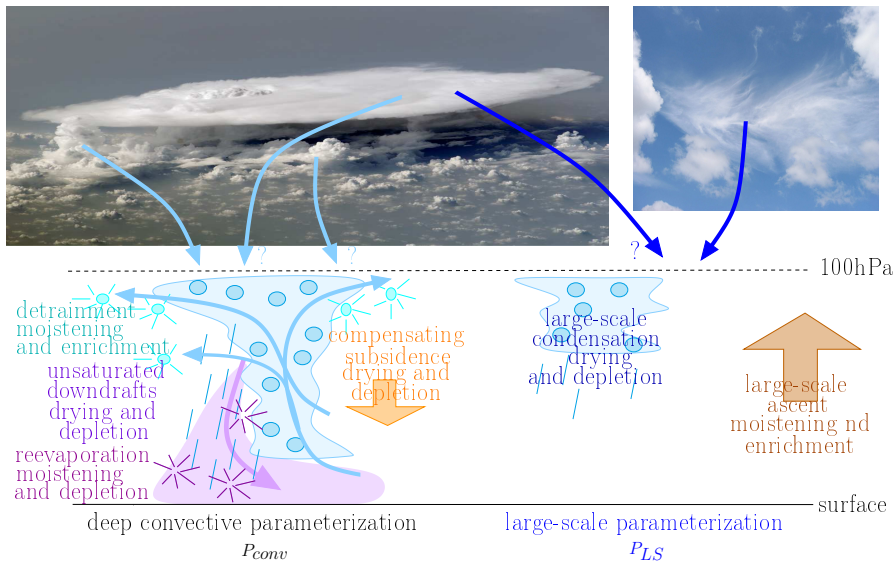
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Interplay convection - large-scale schemes

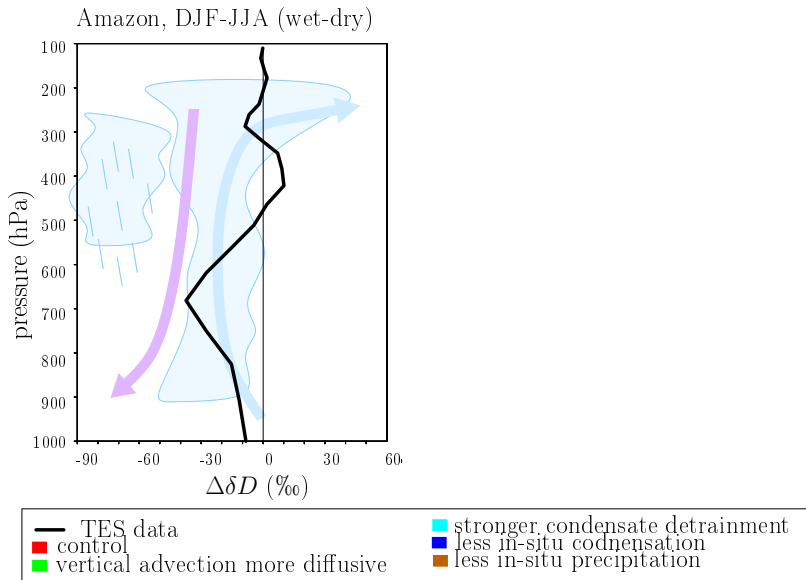


Interplay convection - large-scale schemes

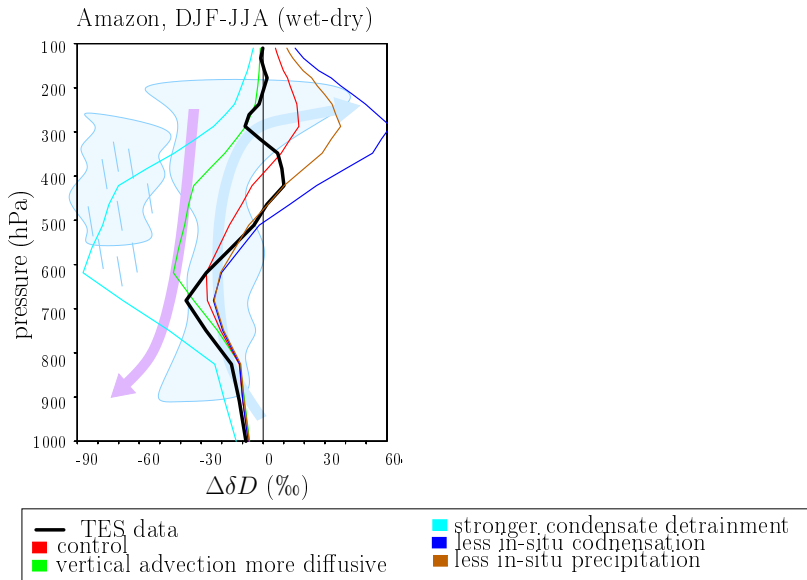


- ▶ P_{LS}/P_{tot} arbitrary, but influences cloudiness, intra-seas. variability, chemical tracer transport

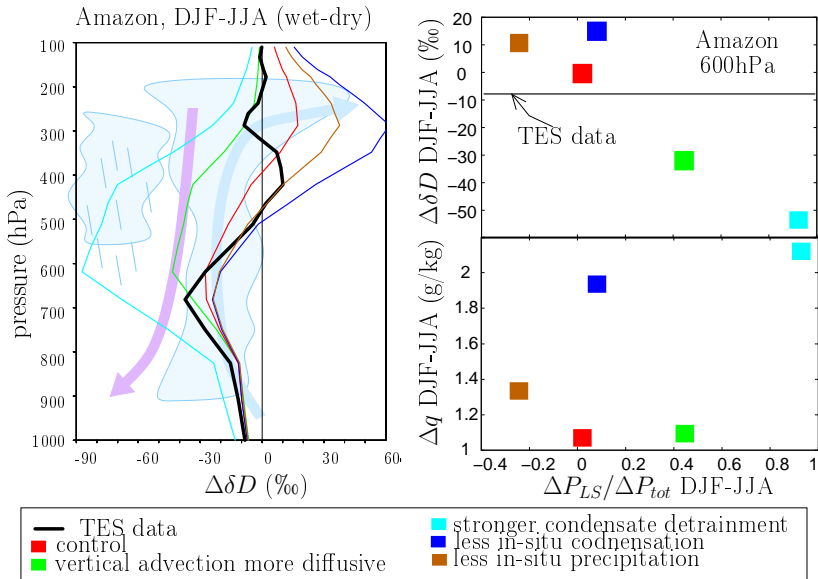
Convection vs large-scale precip



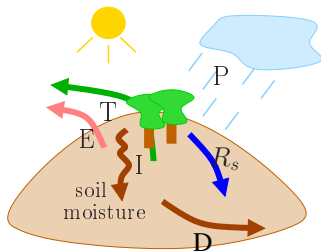
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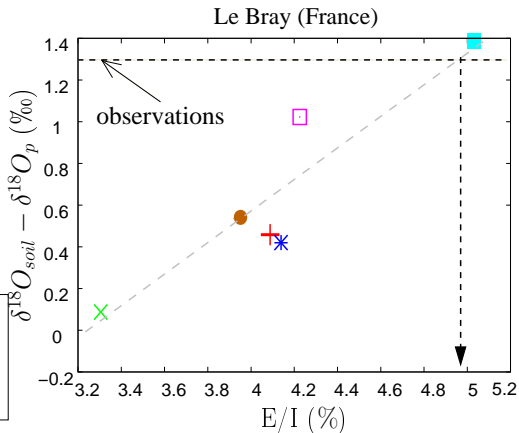
Convection vs large-scale precip



Surface water budget

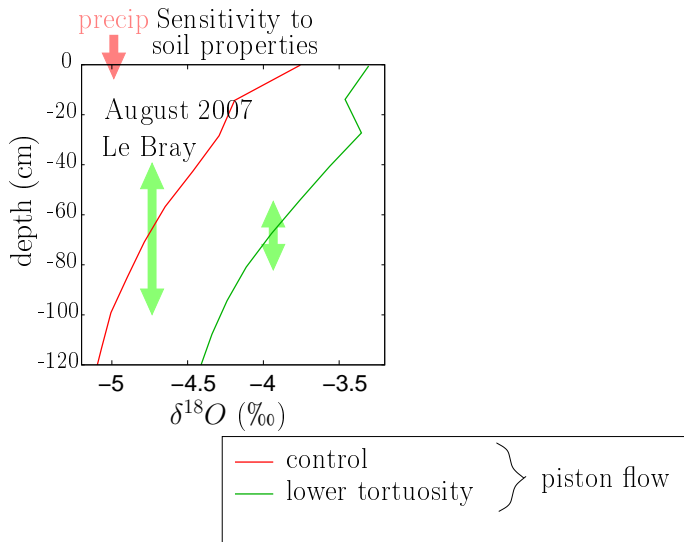


- + control
- × stomatal resistance /5
- no drainage, only surface runoff
- * soil capacity /2
- less vegetation cover
- root extraction depth /4

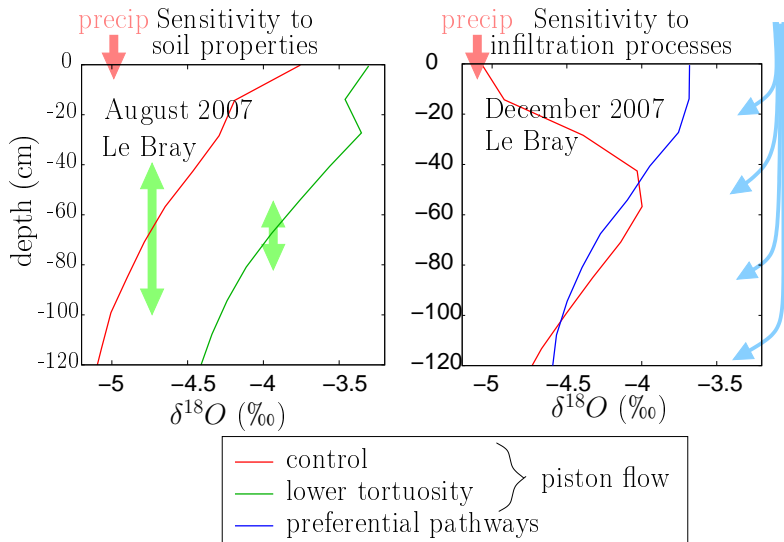


- ▶ soil water isotopic measurements -> bare soil evaporation ratio

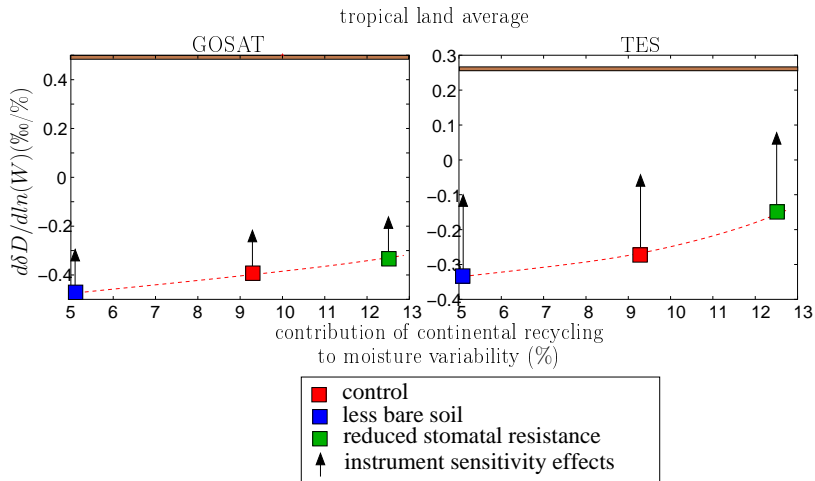
Diffusion/infiltration in soils



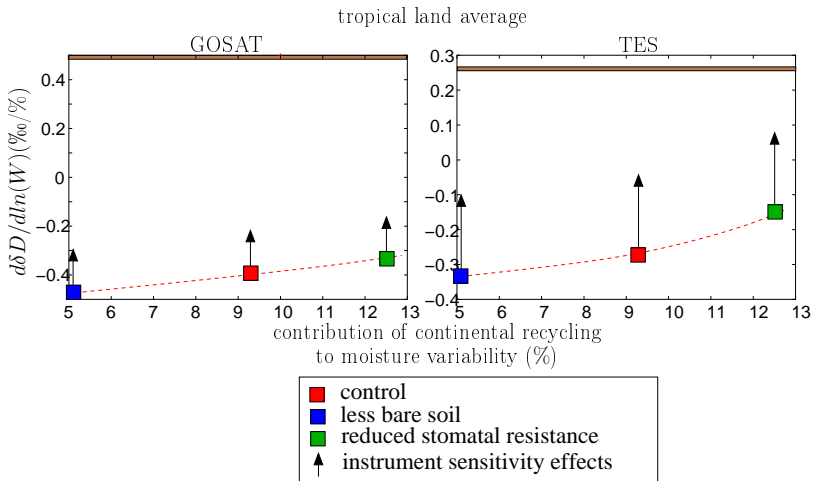
Diffusion/infiltration in soils



Evaluating continental recycling feedbacks



Evaluating continental recycling feedbacks



- ▶ Does LMDZ underestimate the role of continental recycling ?
- ▶ Or atmospheric problems ?