

# Using water stable isotopic measurements to better evaluate the atmospheric and land surface components of climate models

Camille Risi

CIRES, Boulder

with contribution of:

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TES/SCIAMACHY: J Worden, J Lee, D Brown, C Frankenberg

MIPAS/ACE: G Stiller, M Kiefer, B Funke, K Walker, P Bernath,

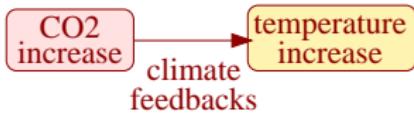
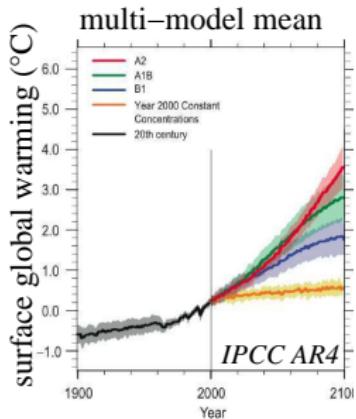
FTIR: M Schneider, D Wunch, P Wennberg, V Sherlock, N Deutscher, D Griffith

in-situ/MIBA: R Uemura, J. Ogée, T. Bariac, L. Wingate, N. Raz-Yaseef

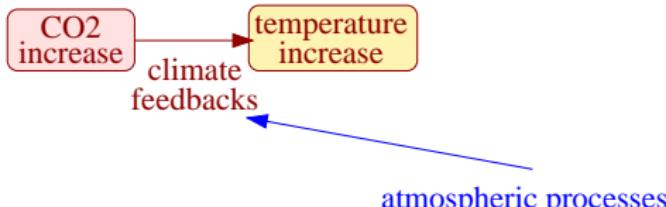
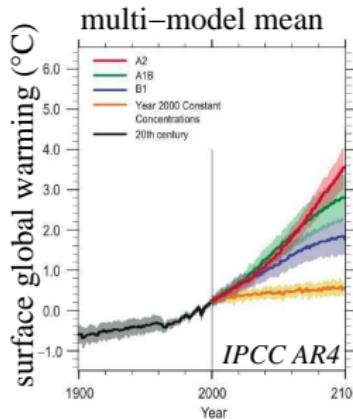
SWING2: C Sturm

Hydrologic Sciences and Water Resources Engineering Seminar  
Series, 27 April 2011

# Uncertainties in climate projections

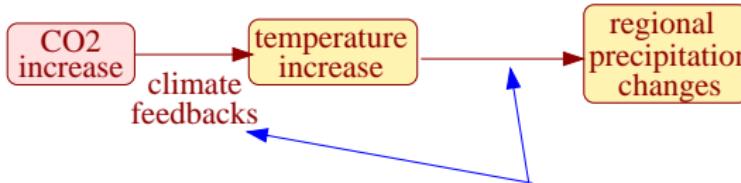
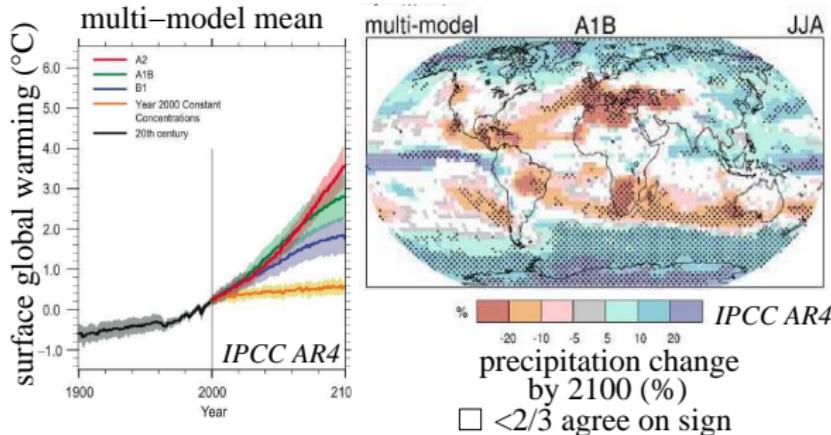


# Uncertainties in climate projections



- clouds
- atmospheric convection
- boundary layer
- relative humidity

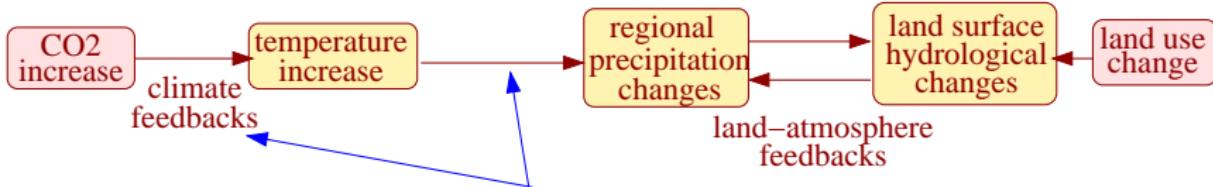
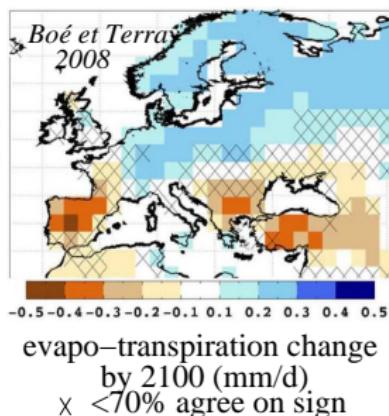
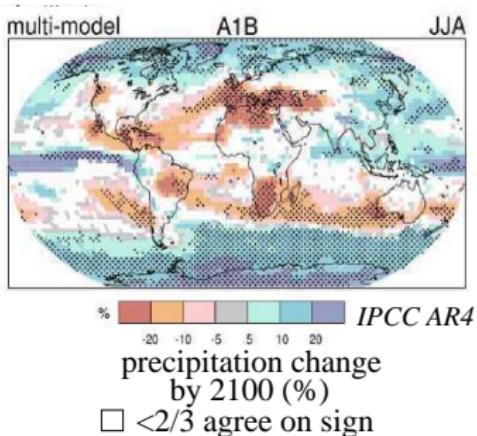
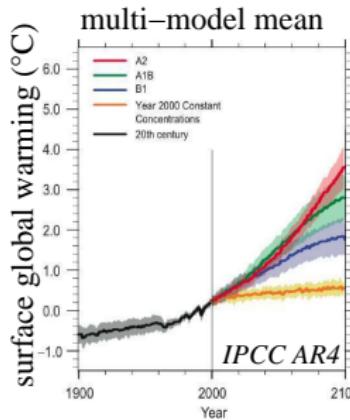
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Key uncertainties  
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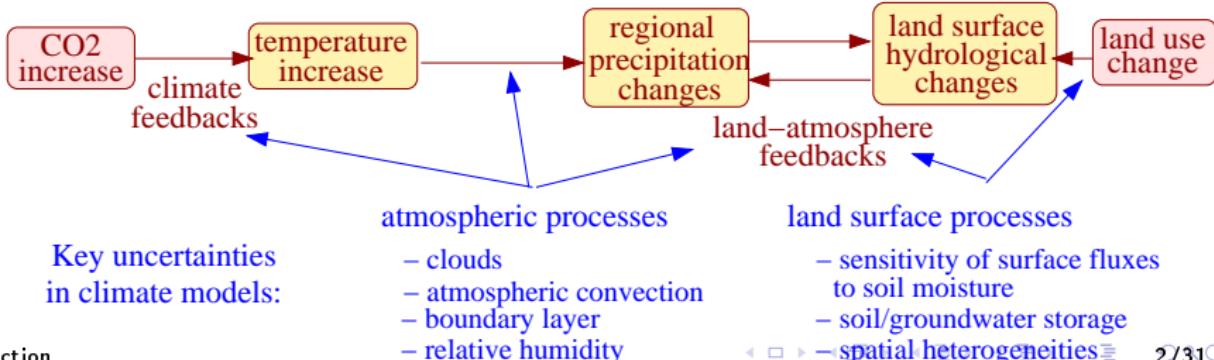
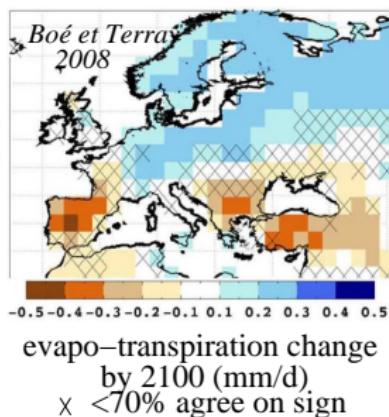
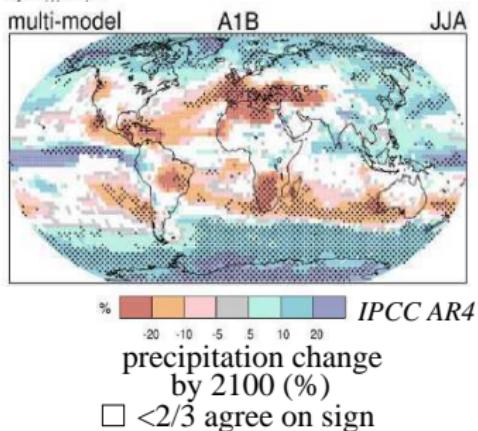
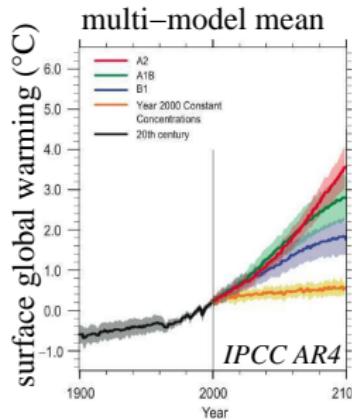
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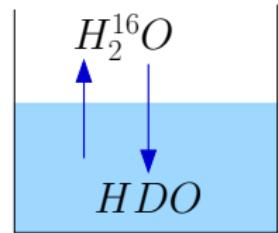
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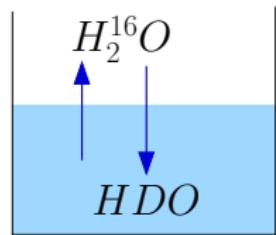
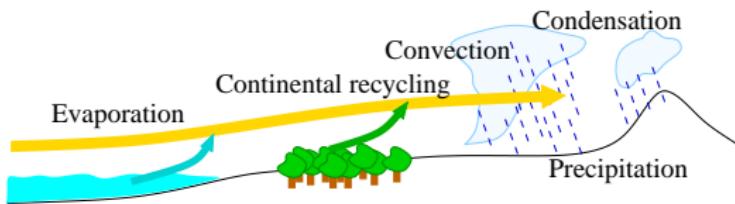
## Water isotopic composition

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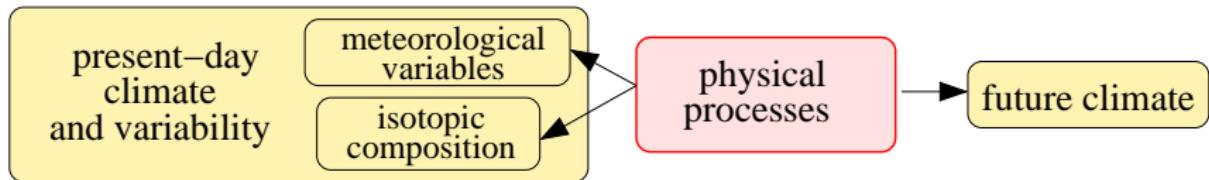
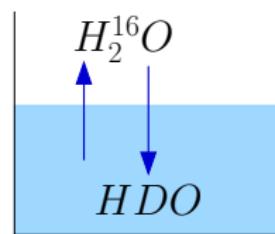
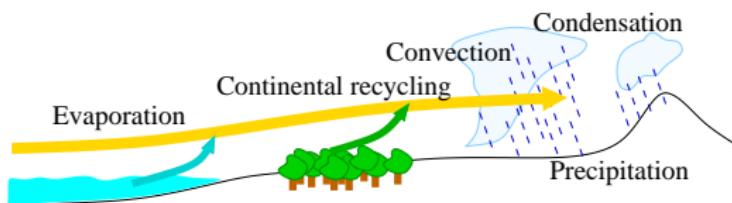
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  - ▶ records phase changes



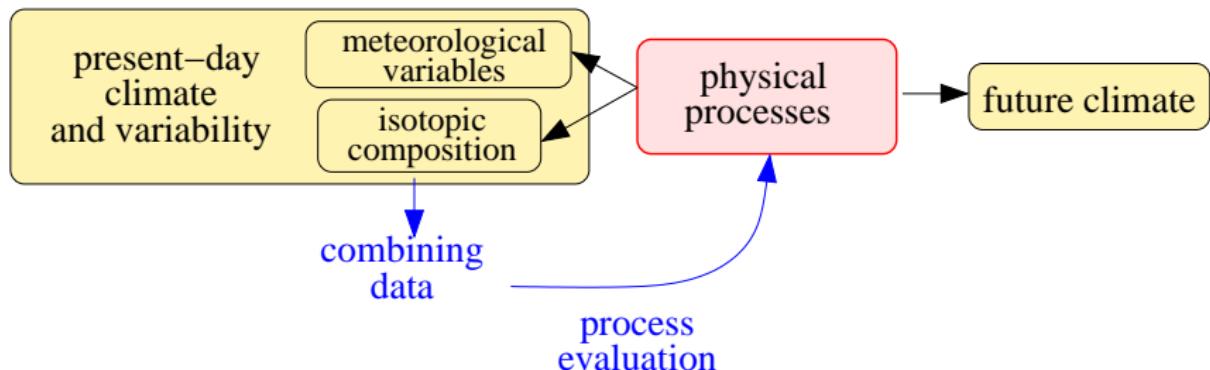
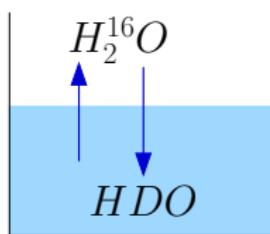
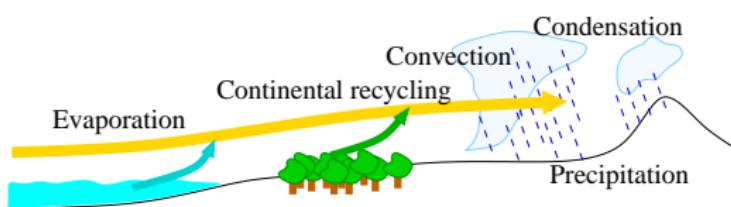
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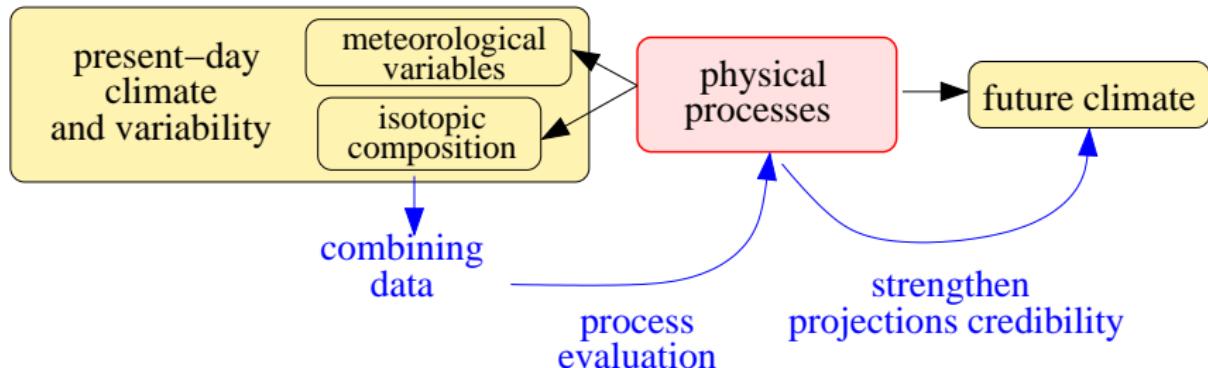
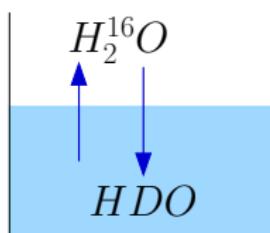
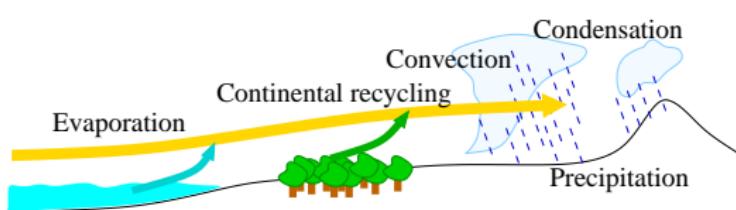
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# General strategy

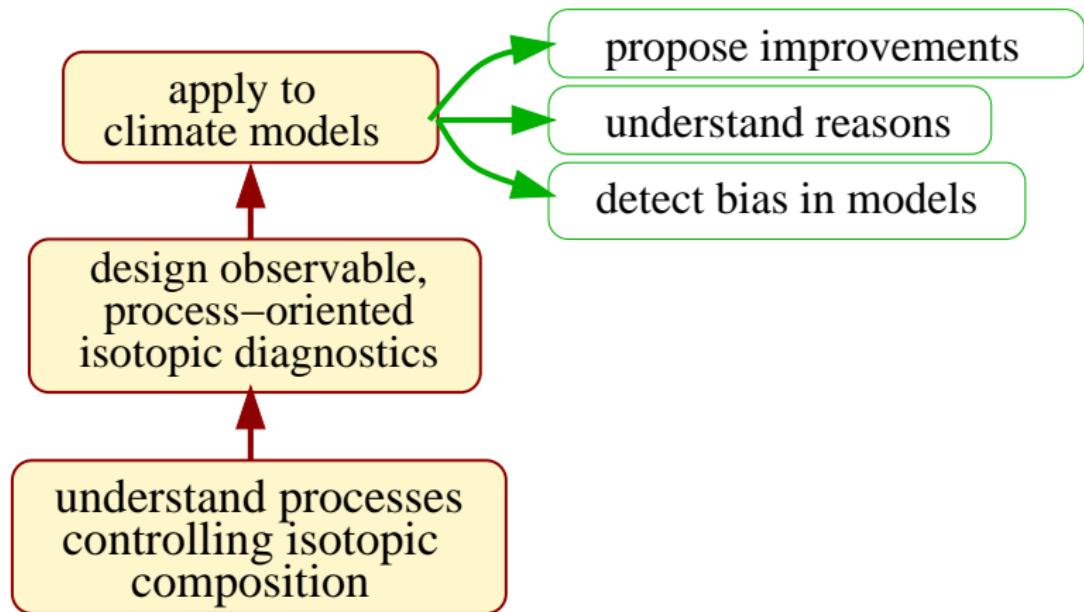
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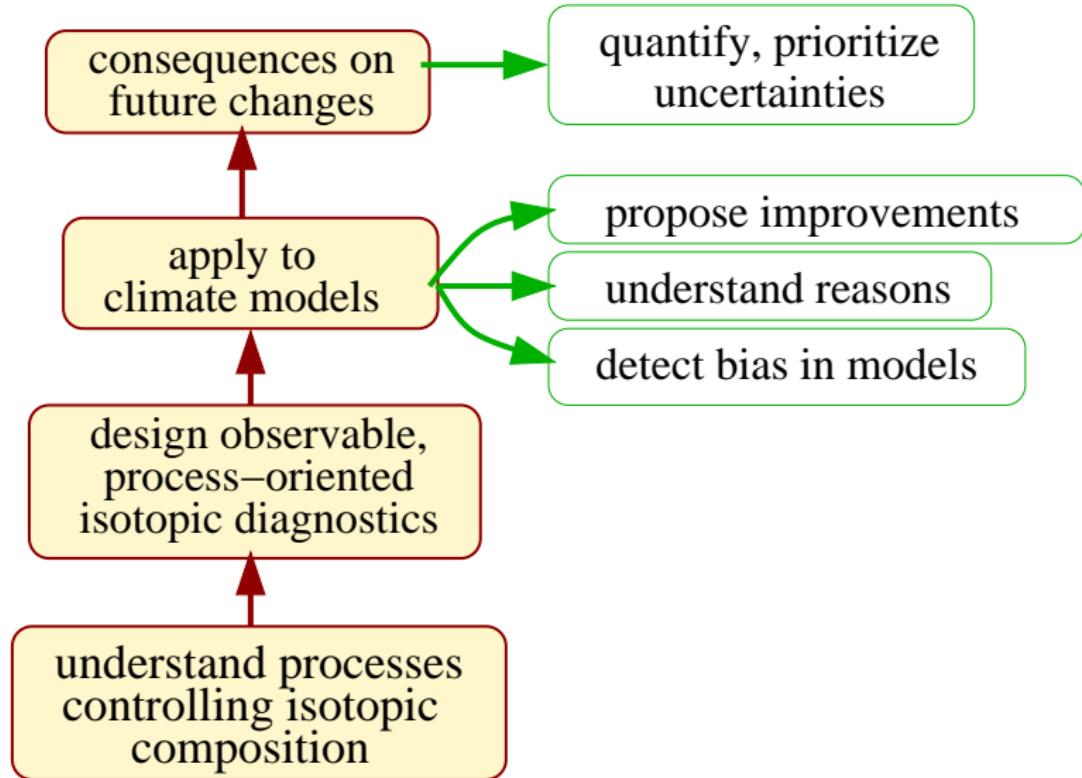
design observable,  
process-oriented  
isotopic diagnostics

understand processes  
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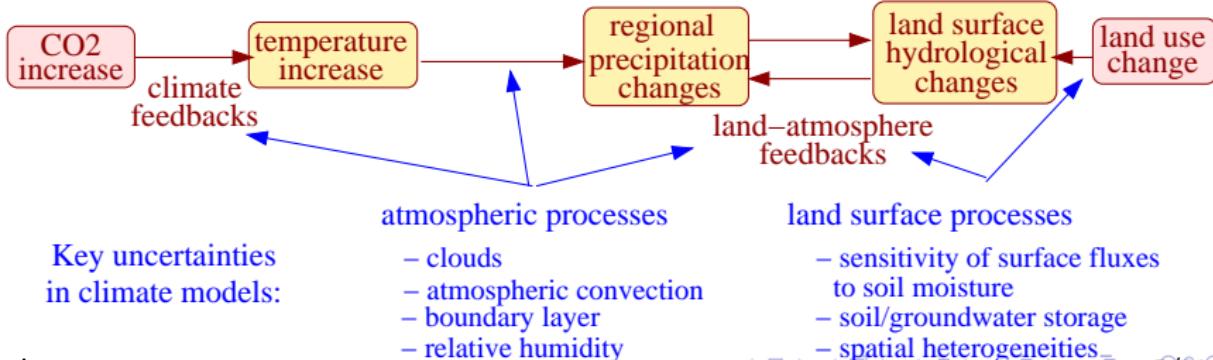
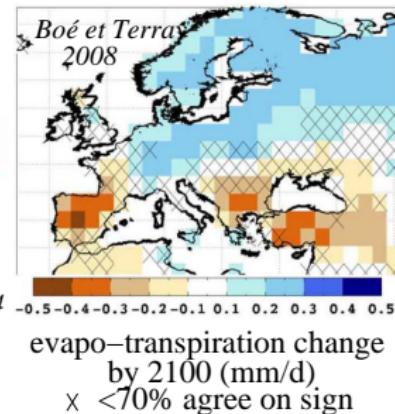
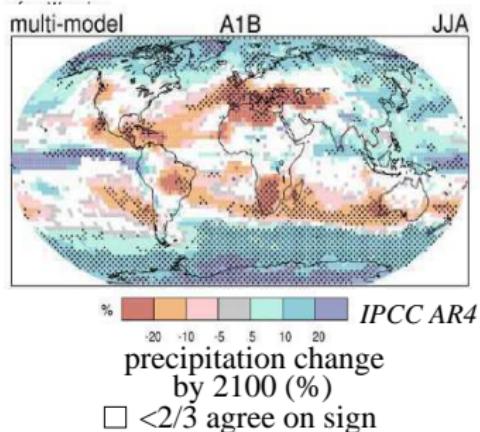
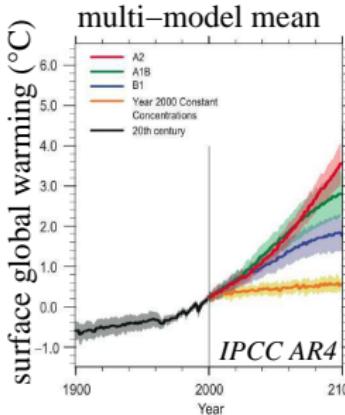
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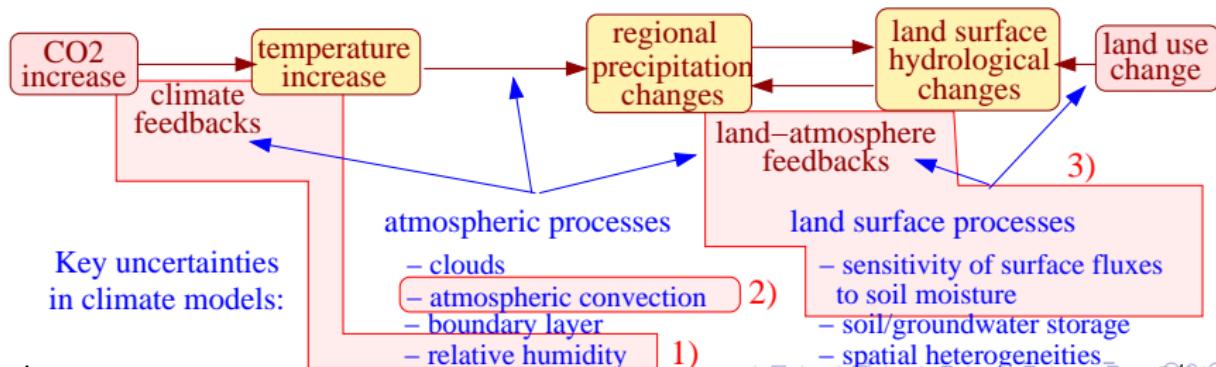
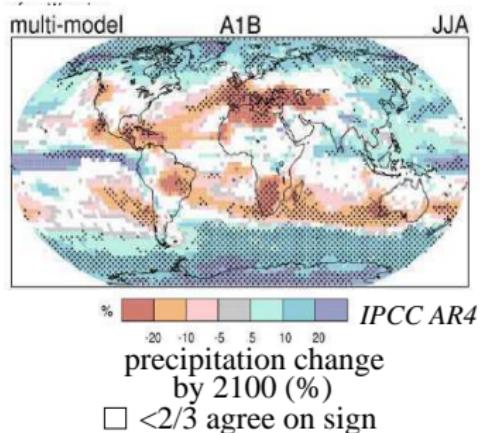
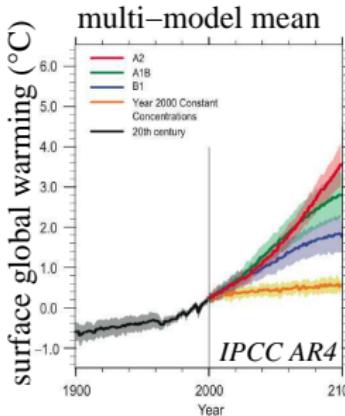
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# Outline



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# 1) Processes controlling relative humidity

- ▶ tropical/subtropical free tropospheric relative humidity (RH) impacts:
  - ▶ water vapor feedback (*Soden et al 2008*)
  - ▶ cloud feedbacks (*Sherwood et al 2010*)

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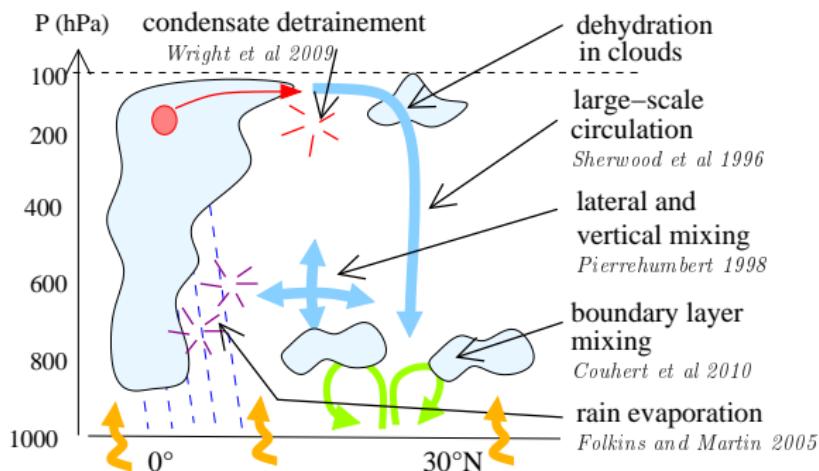
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⇒ need process-based evaluation of RH in climate models

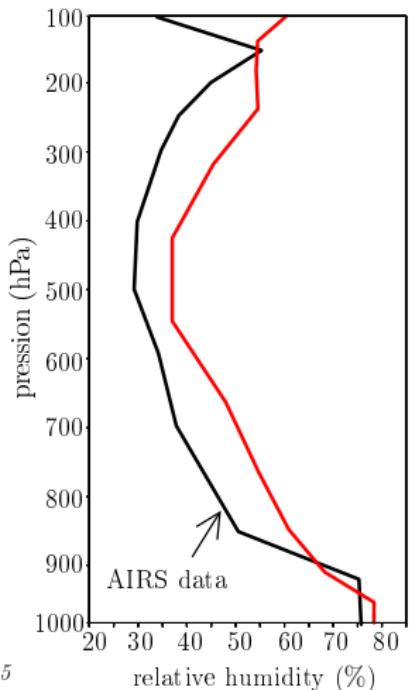
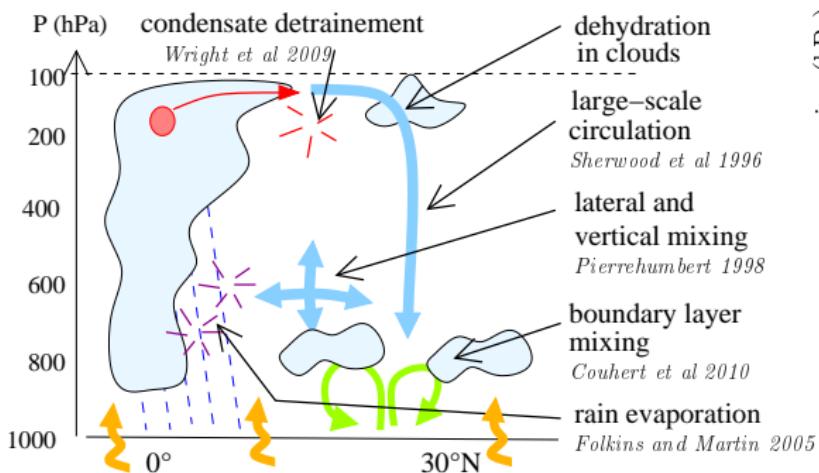
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— control: AR4 version (19 levels)

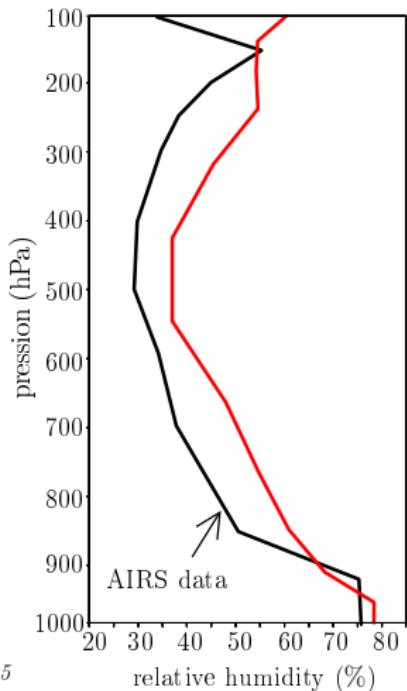
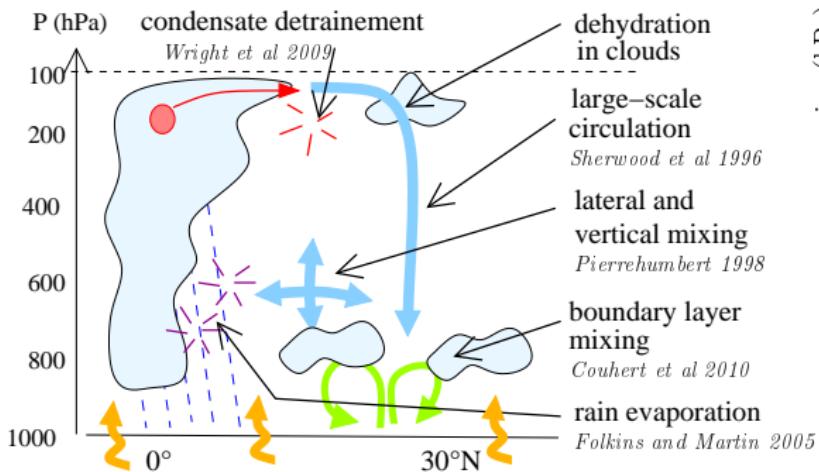


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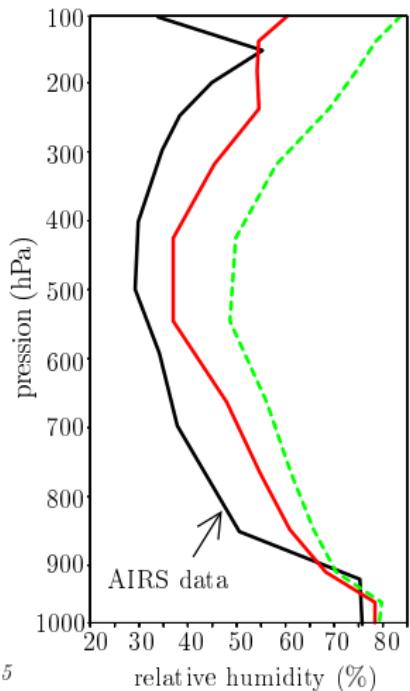
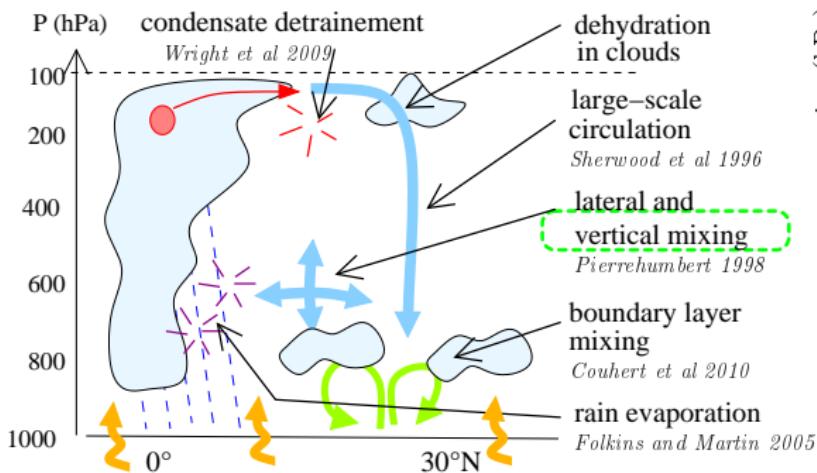


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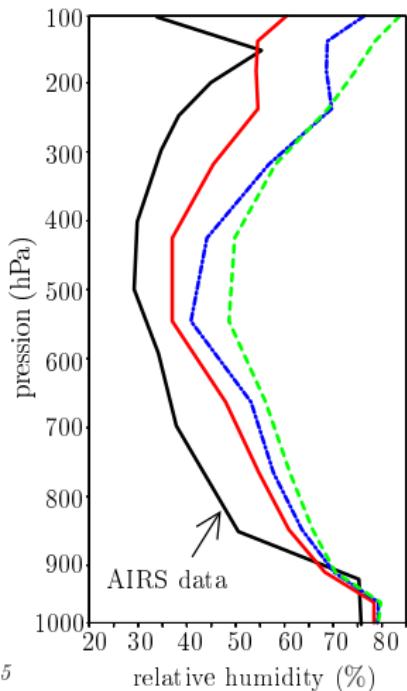
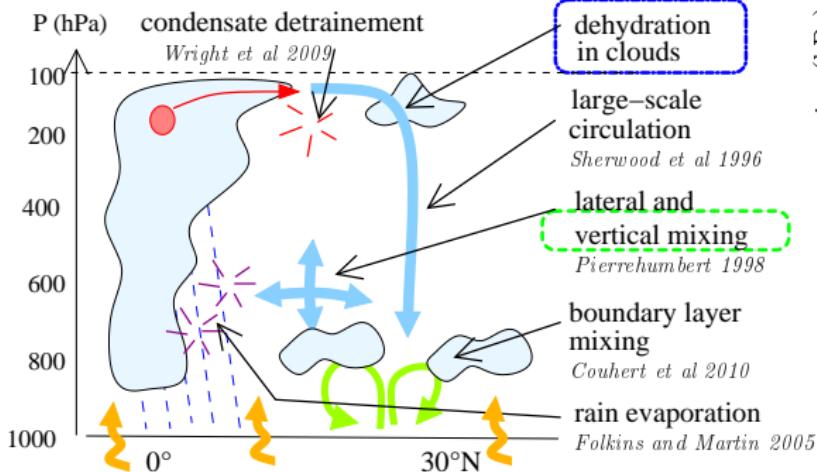


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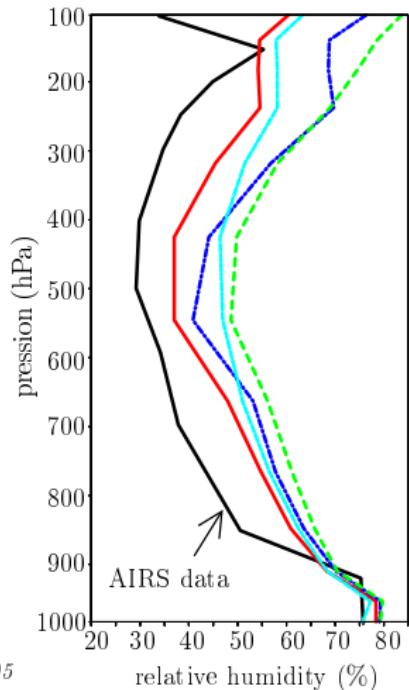
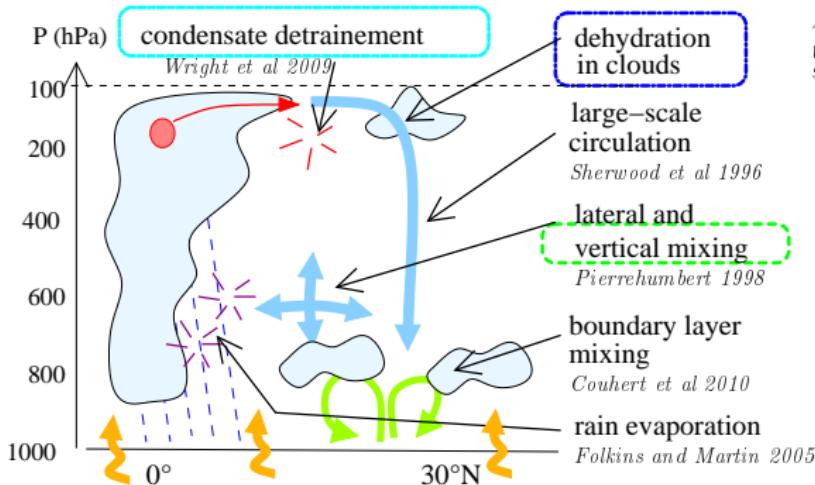


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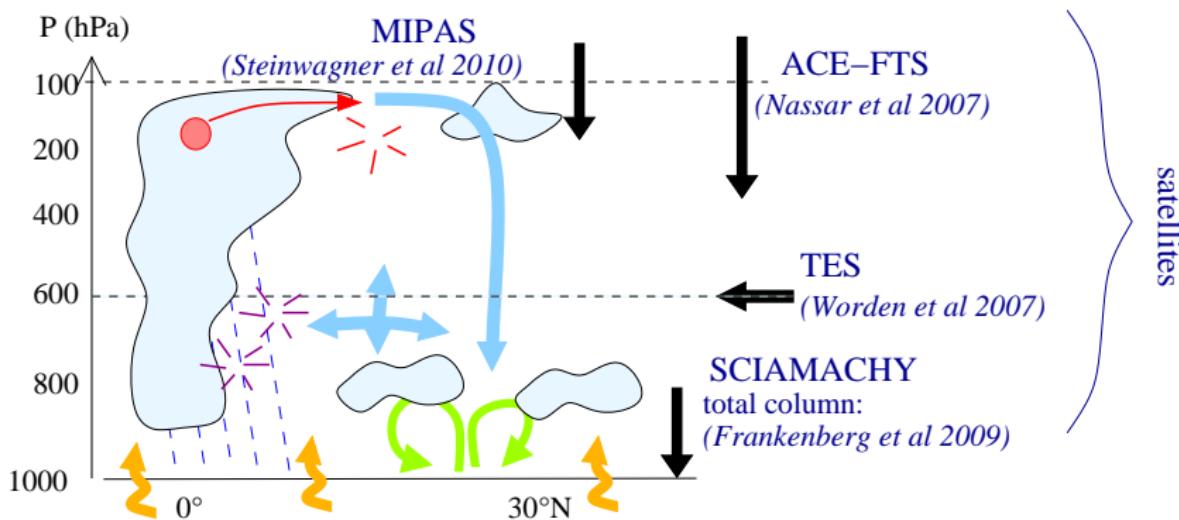
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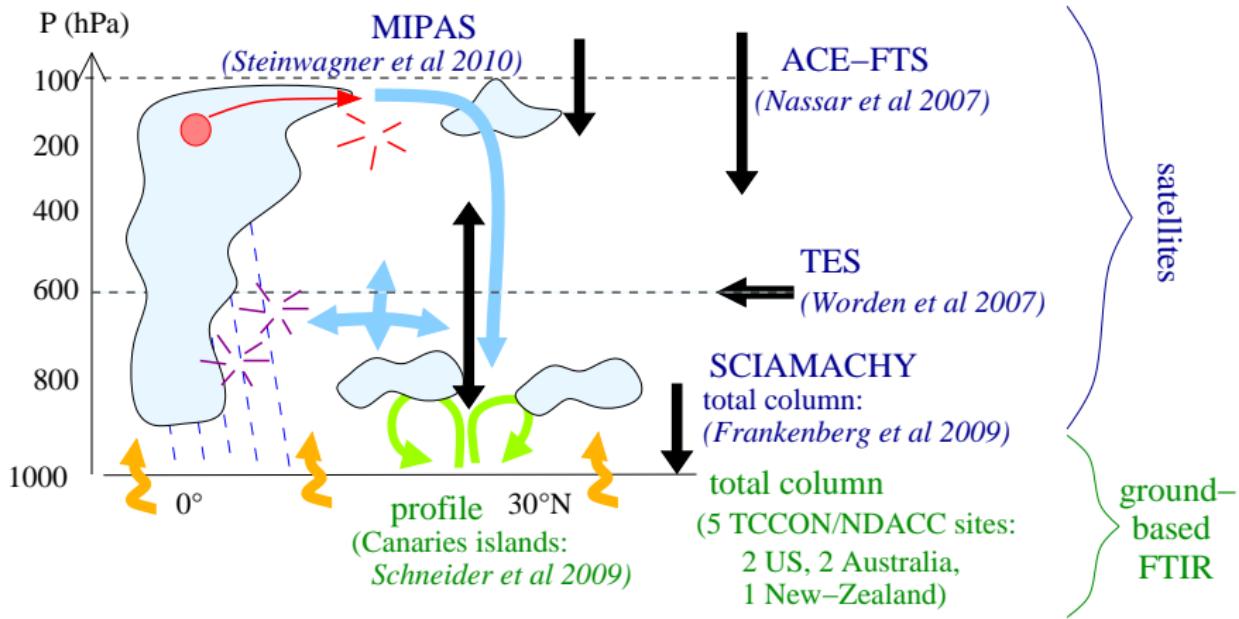
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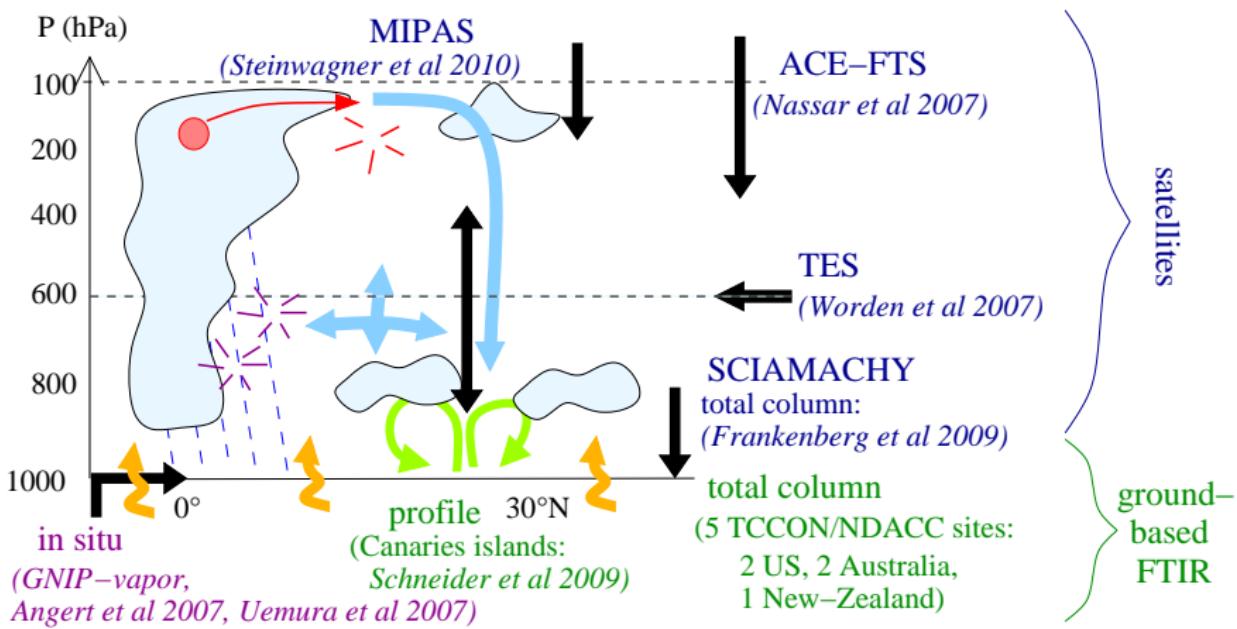
## Isotopic measurements



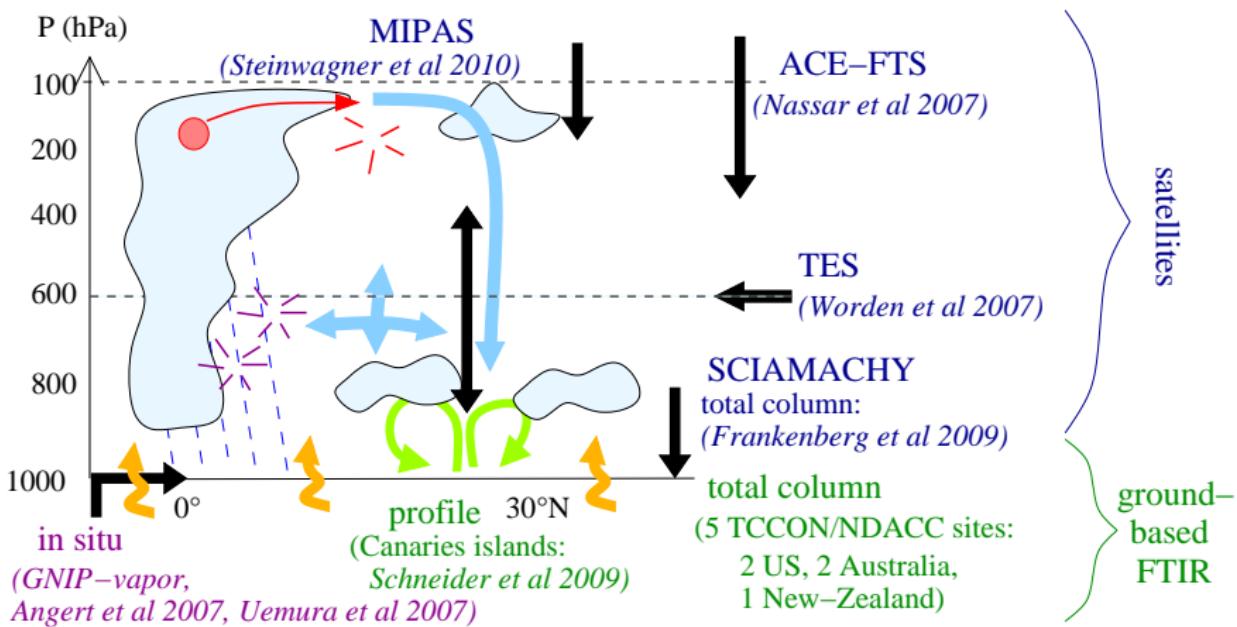
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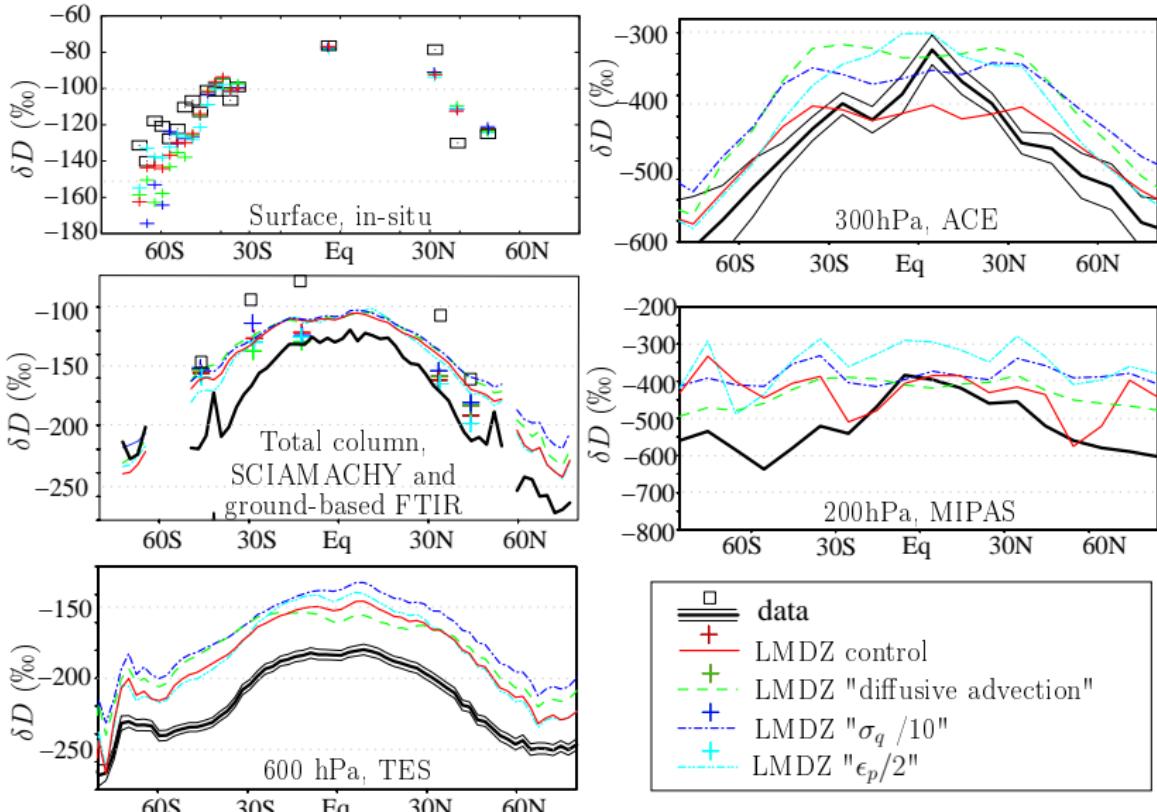


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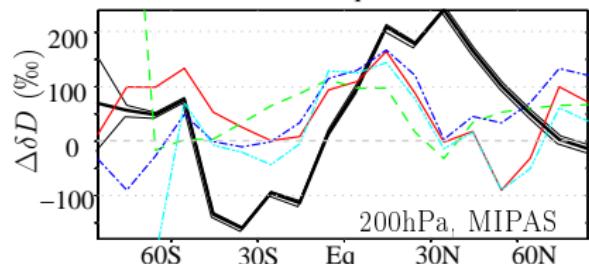
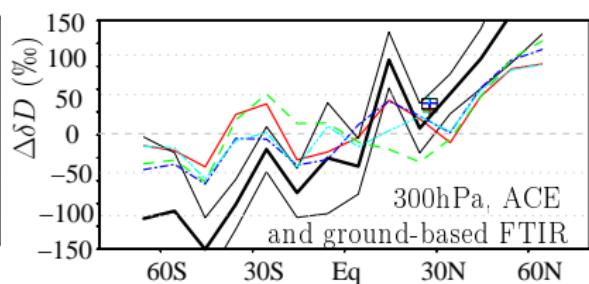
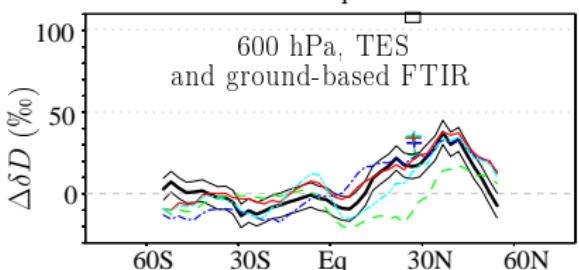
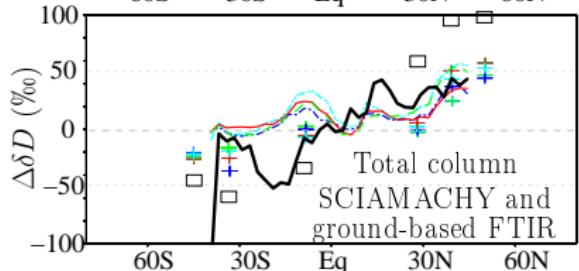
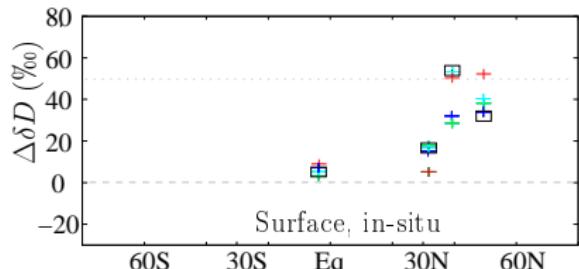


- ▶ model-data comparison: collocation; simulations nudged by ECMWF; averaging kernels; spatial/temporal variations

## Zonal annual mean



## Zonal Seasonal variations (JJA-DJF)



□ data

+ LMDZ control

+ LMDZ "diffusive advection"

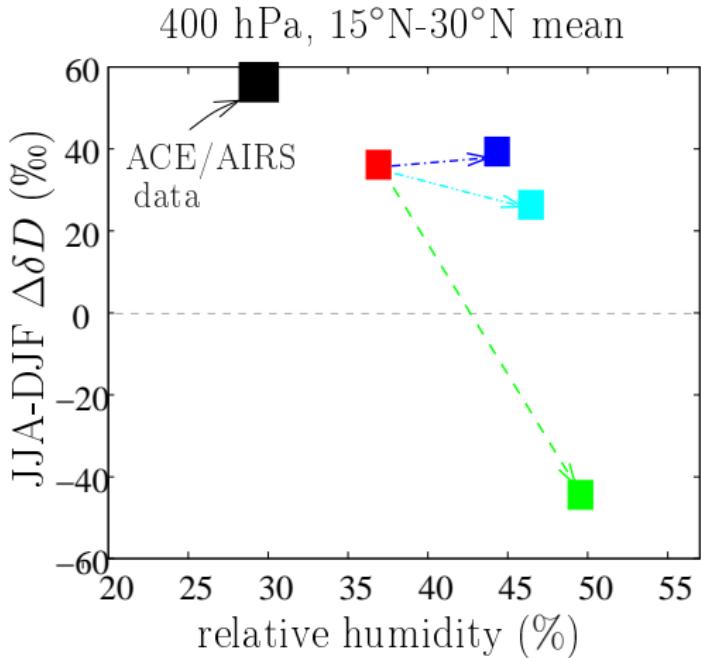
LMDZ "dim usive" /10/

LMDZ +  $\sigma_q / 10$   
+ LMDZ || - / ? ||

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Sensitivity tests:  
with LMDZ:

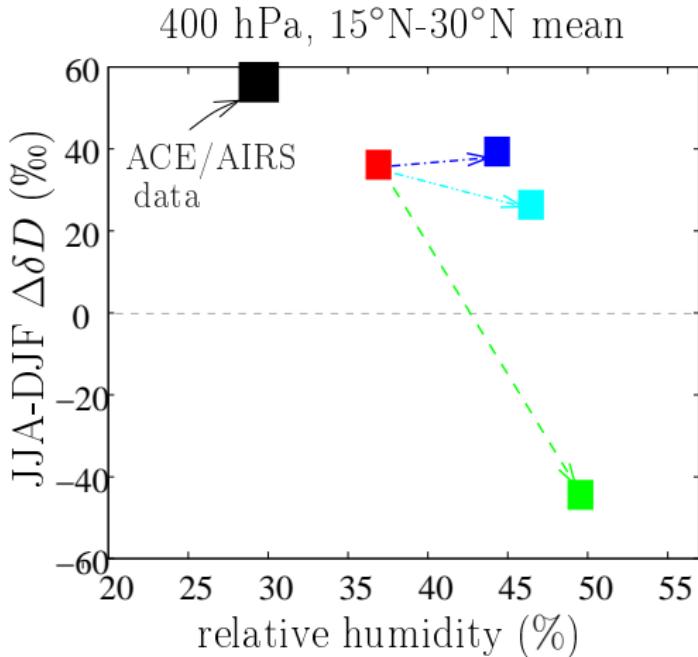
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- ▶ robustness? additional tests, theoretical understanding

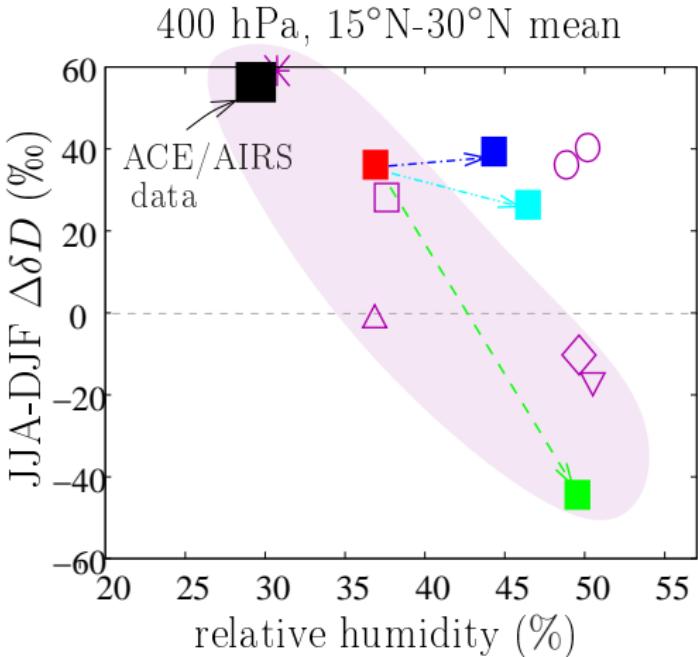
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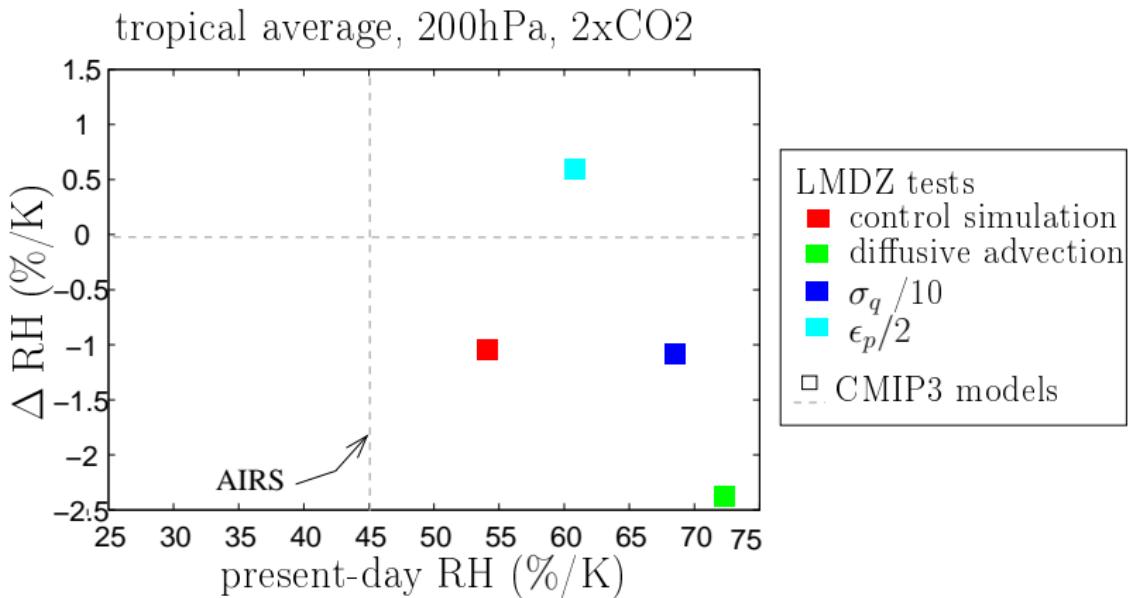
SWING2 models:

- |         |        |
|---------|--------|
| □ ECHAM | ◇ CAM2 |
| △ MIROC | ○ GISS |
| * HadAM | ▽ GSM  |

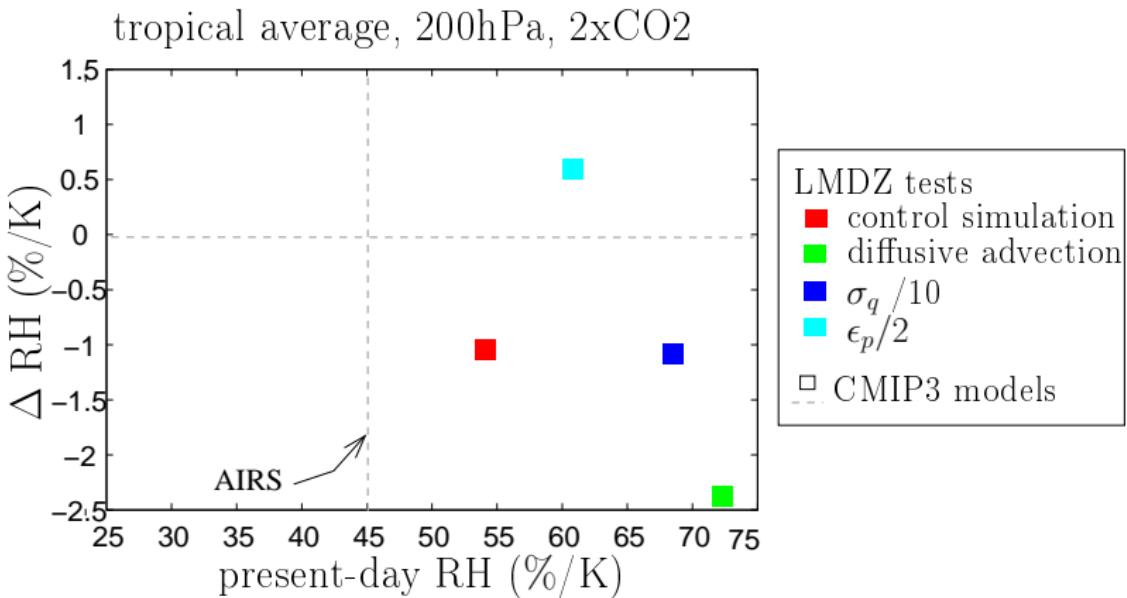


- ▶ robustness? additional tests, theoretical understanding
- ▶ frequent reason for moist bias=excessively diffusive advection

# What impact on humidity projections?

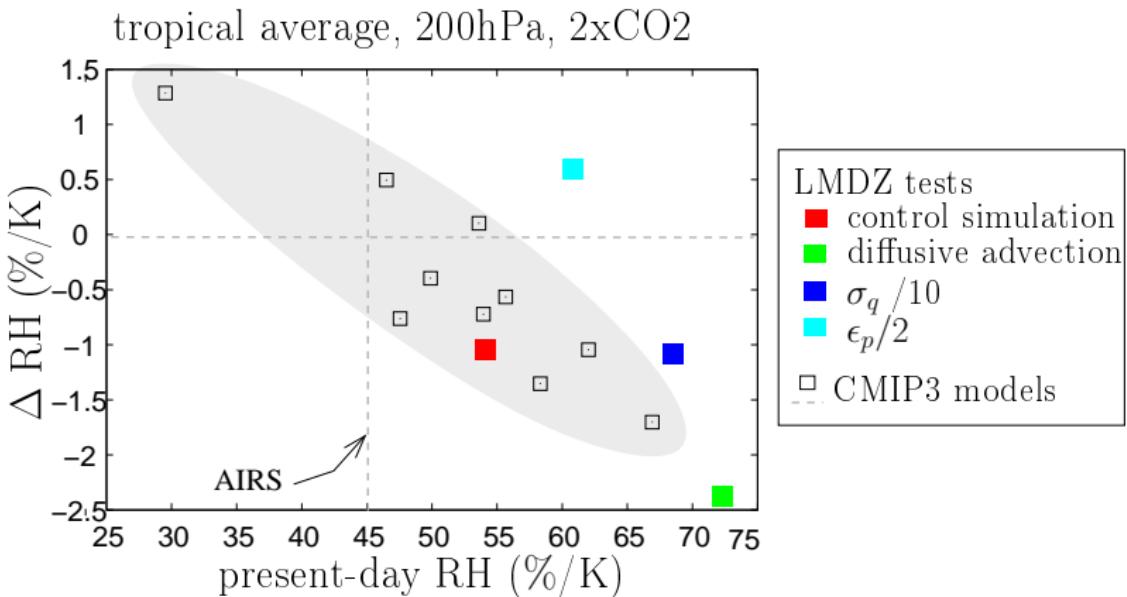


# What impact on humidity projections?



- ▶ How a moist bias affect humidity change projections depends on the reason for the bias

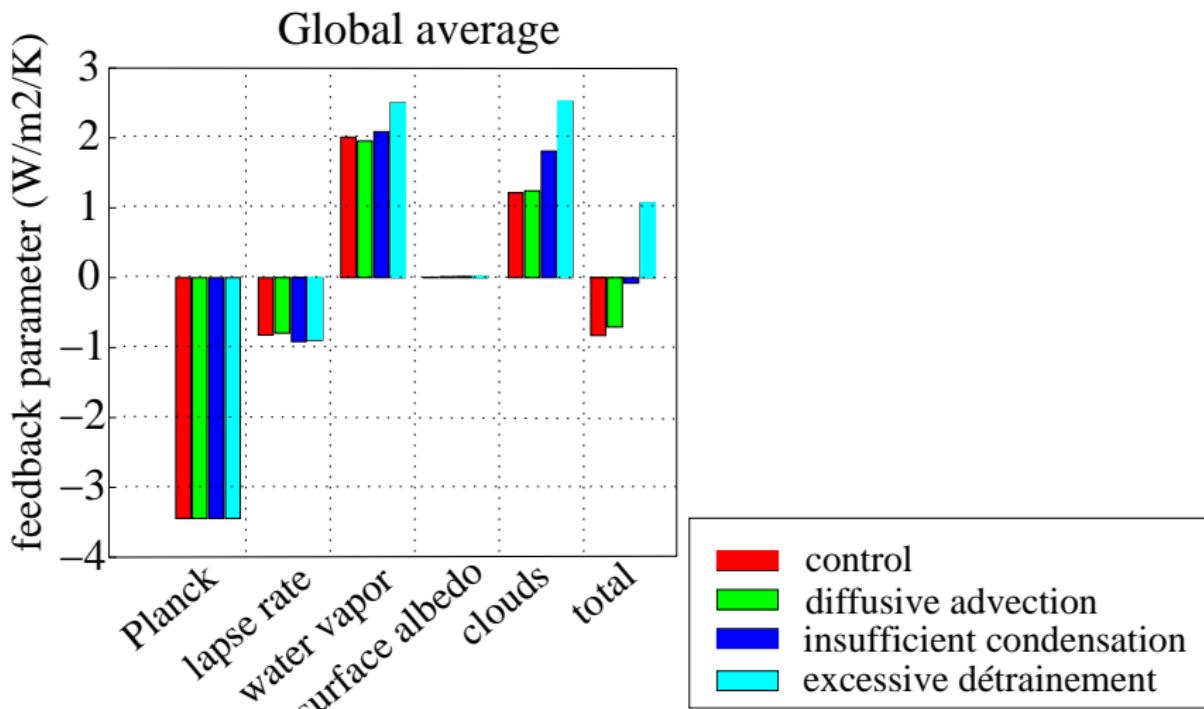
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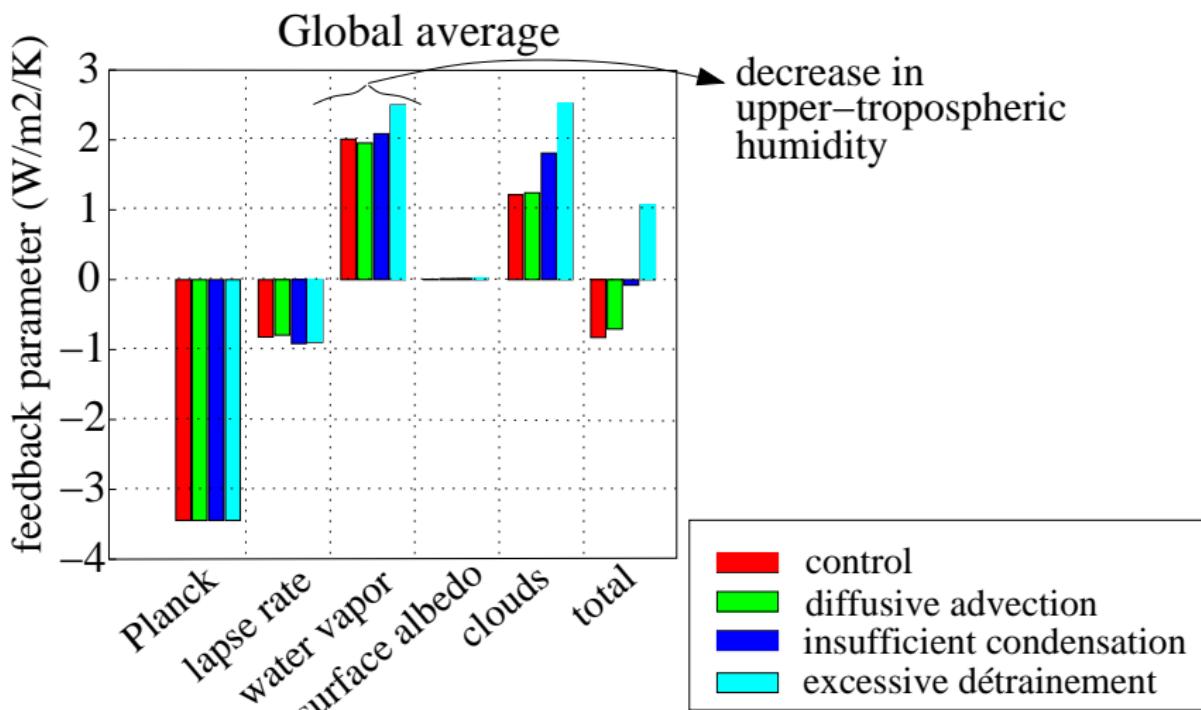
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- radiative kernel decomposition (*Soden et al 2008*)



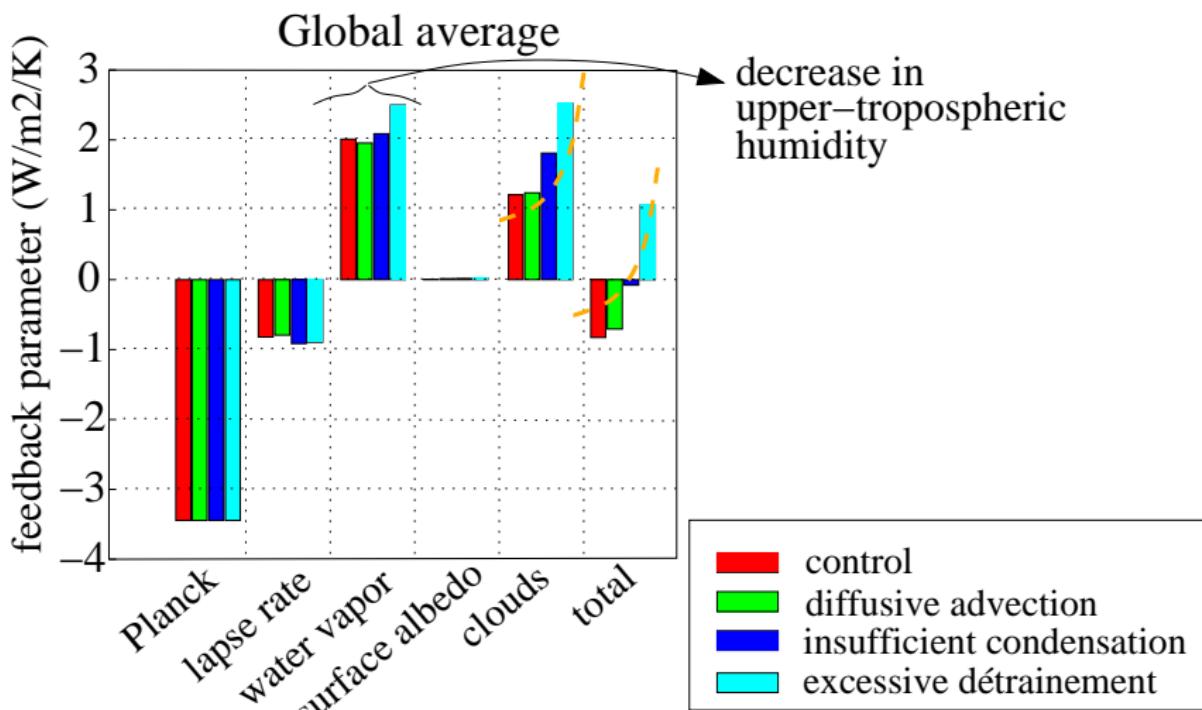
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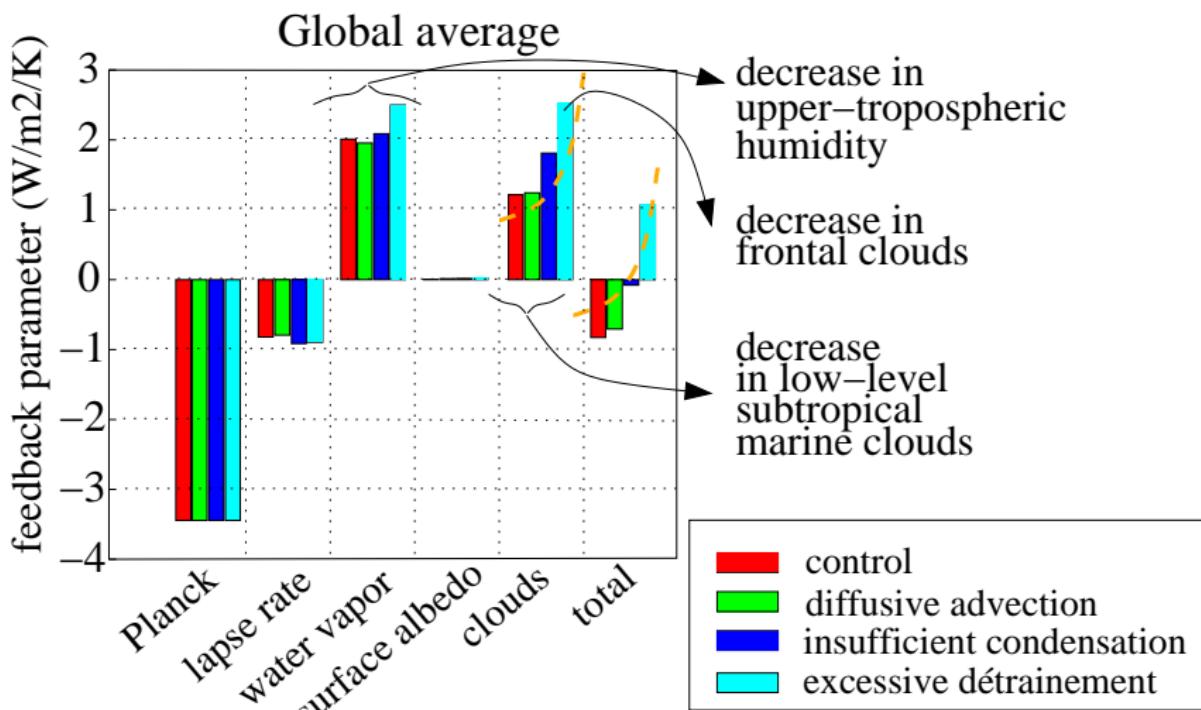
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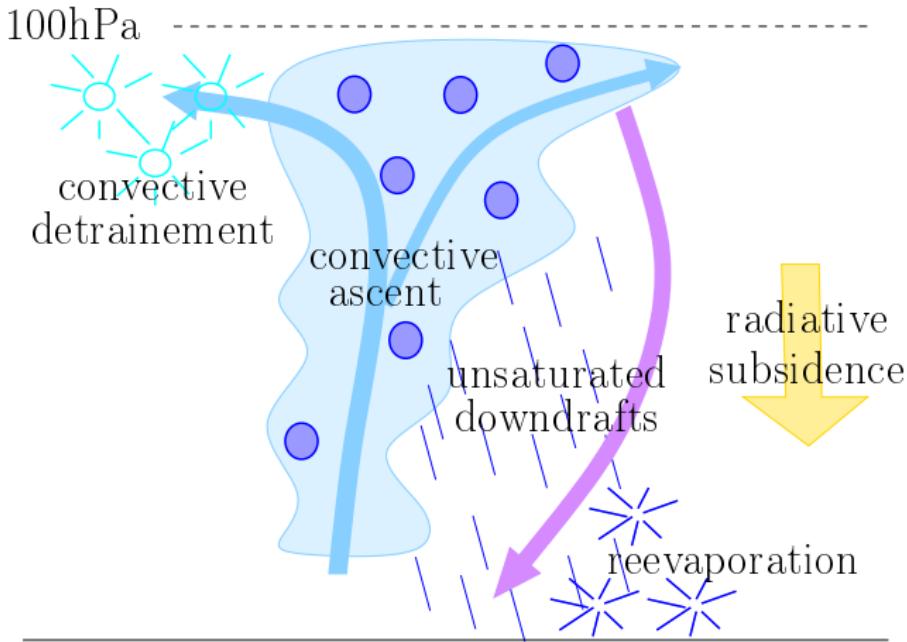
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- ▶ Excessive vertical diffusion during water vapor transport is a widespread cause of moist bias in climate models
- ▶ Understanding this reason is all the more important as humidity change projections depends on the reason for the moist bias
- ▶ Consequences on climate change complicated by dynamical and cloud feedbacks

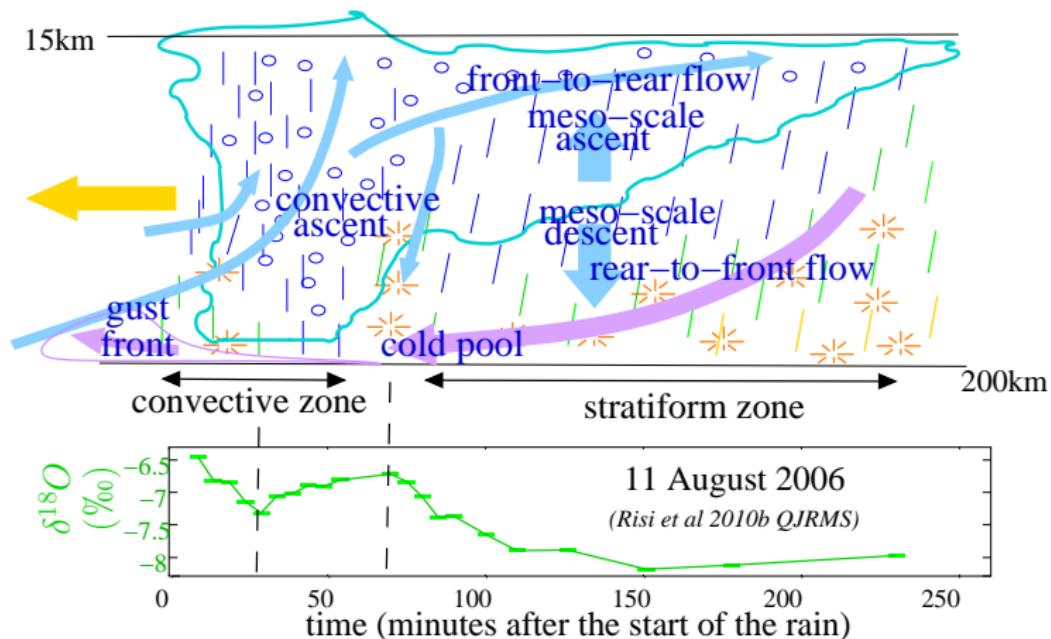
## 2) Convective processes

- ▶ microphysical processes? (*Emanuel and Pierrehumbert 1996*)



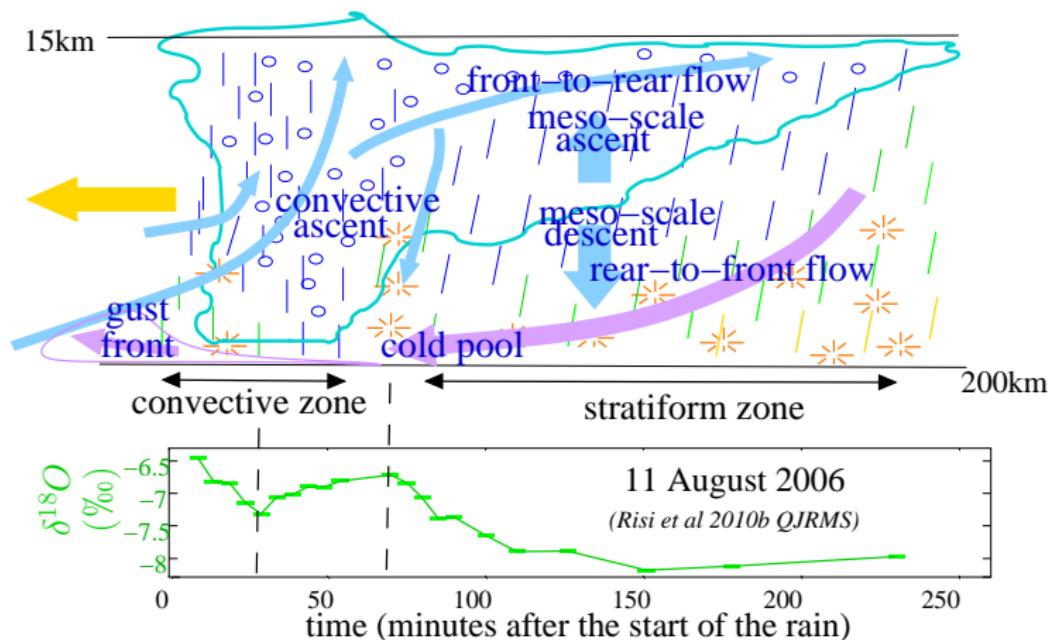
## Processes along squall lines

- ▶ rain sampled every 5 mins in Niamey during AMMA campaign



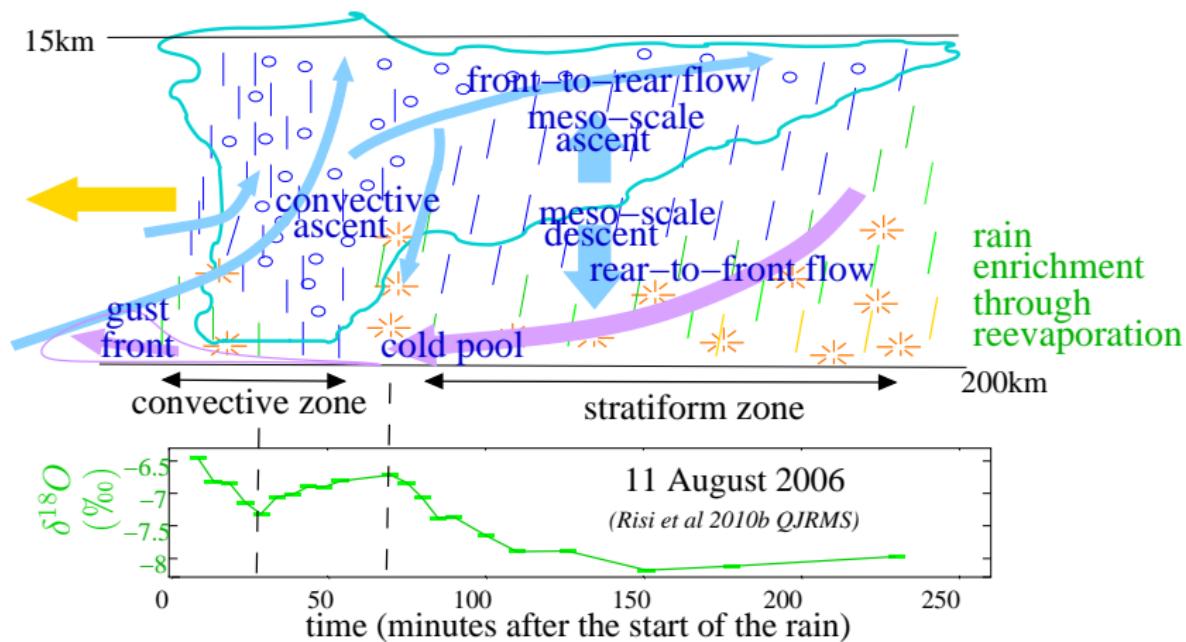
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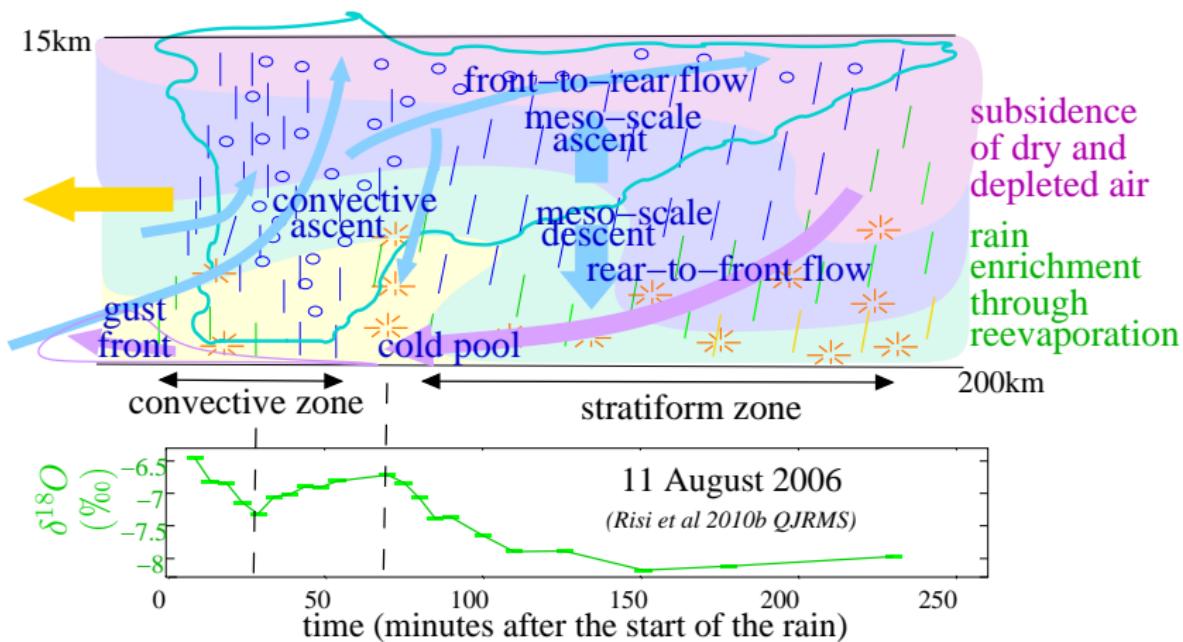
## Processes along squall lines

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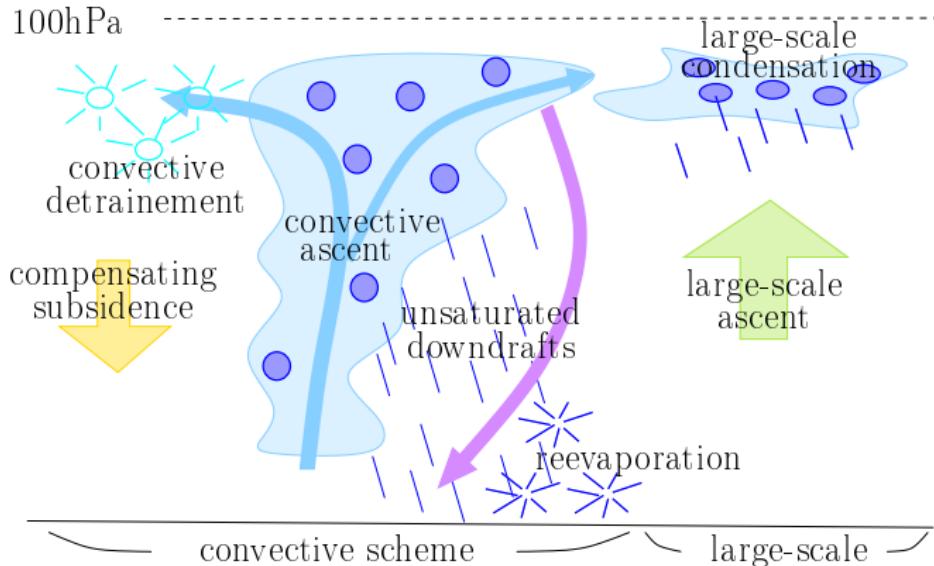


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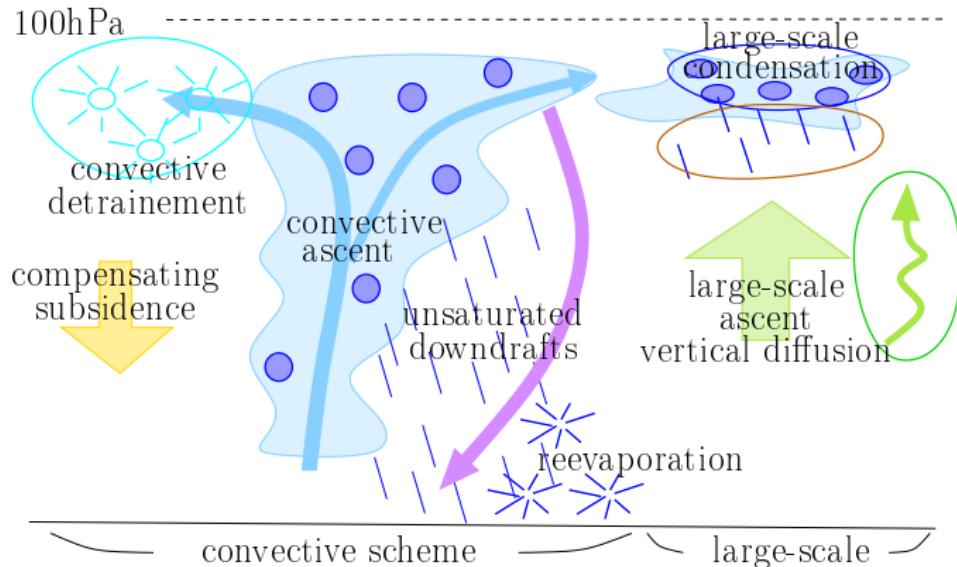
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## Convective/large-scale fluxes



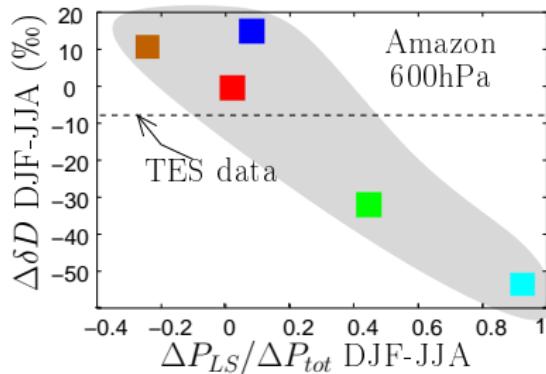
# Convective/large-scale fluxes



Sensitivity tests with LMDZ:

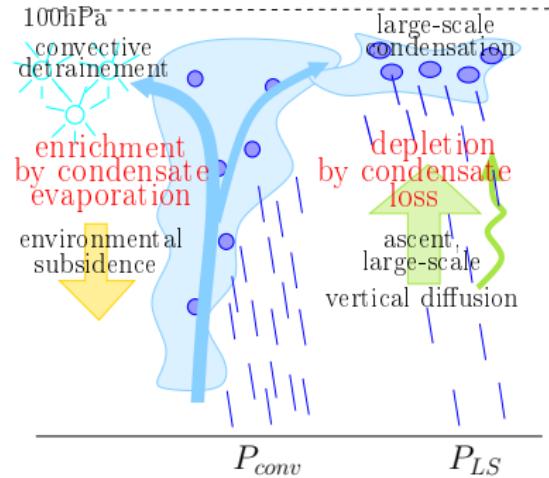
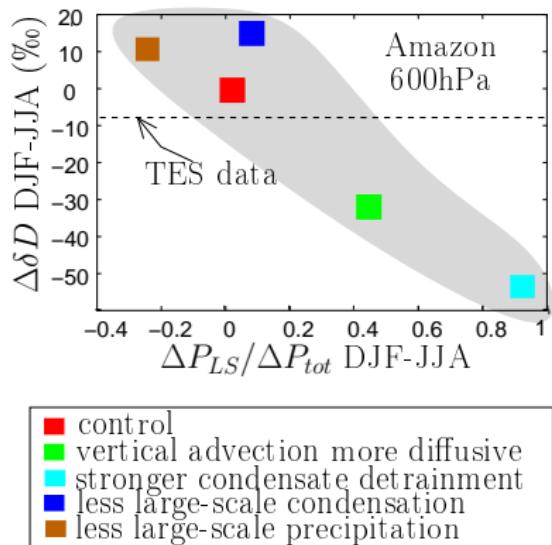
- control: AR4
- more diffusive vertical advection
- stronger condensate detrainement
- less large-scale condensation
- less large-scale precipitation

# Convective contribution to water budget

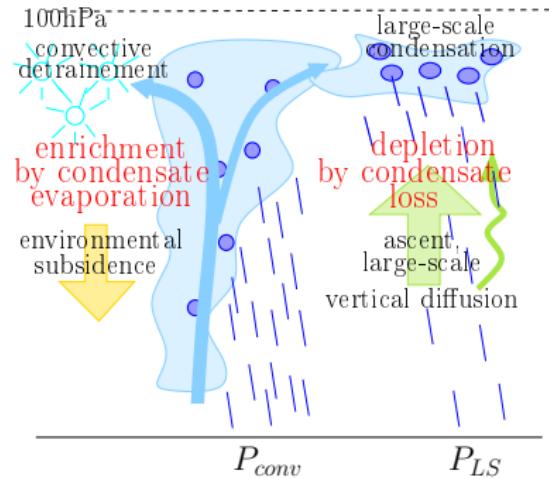
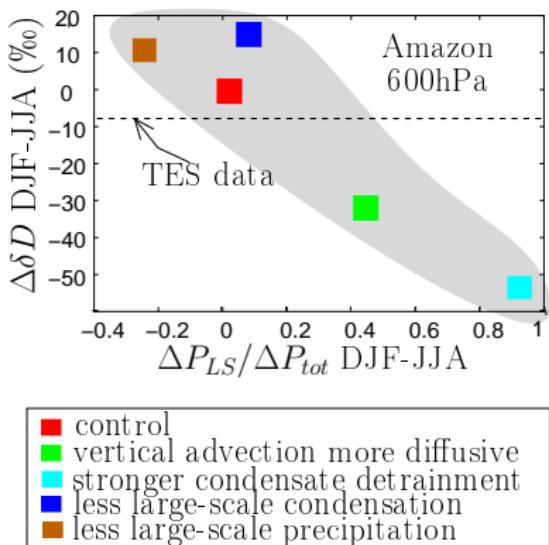


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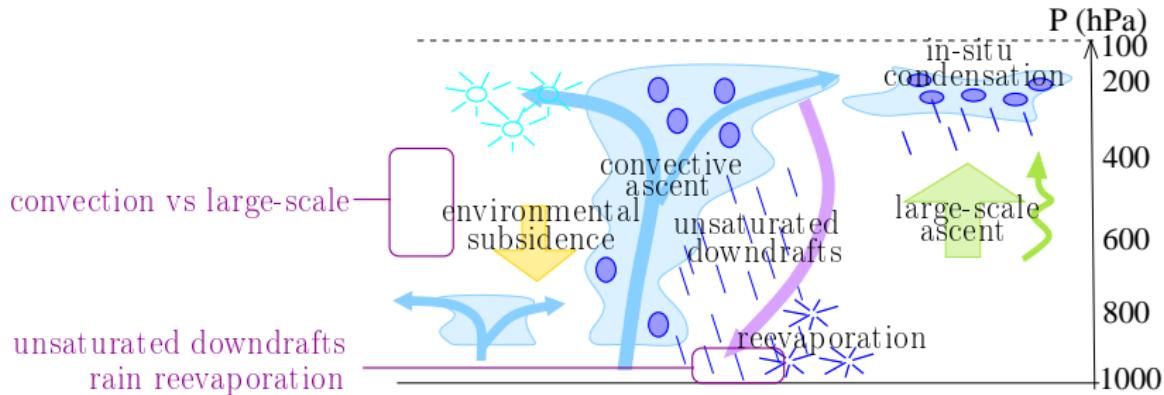


## Convective contribution to water budget

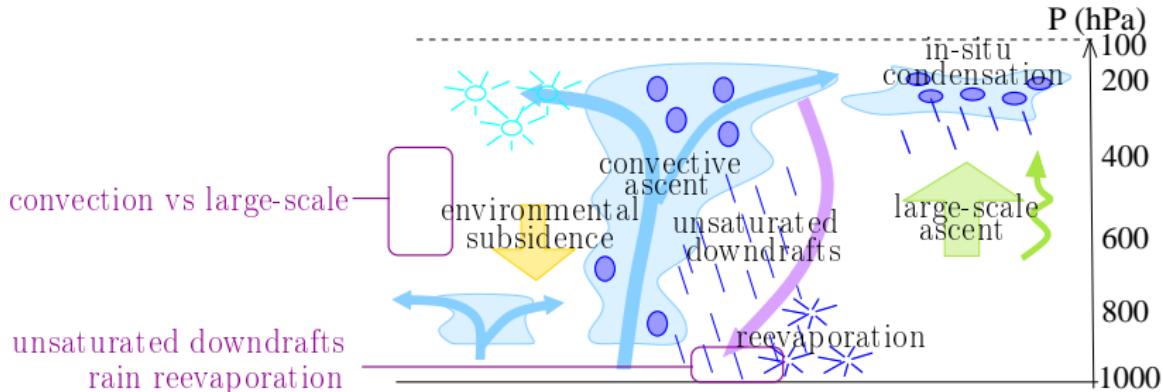


- ▶  $P_{LS}/P_{tot}$  ill-defined quantity, but influences cloudiness, intra-seasonal variability, chemical tracer transport

## Summary on convection

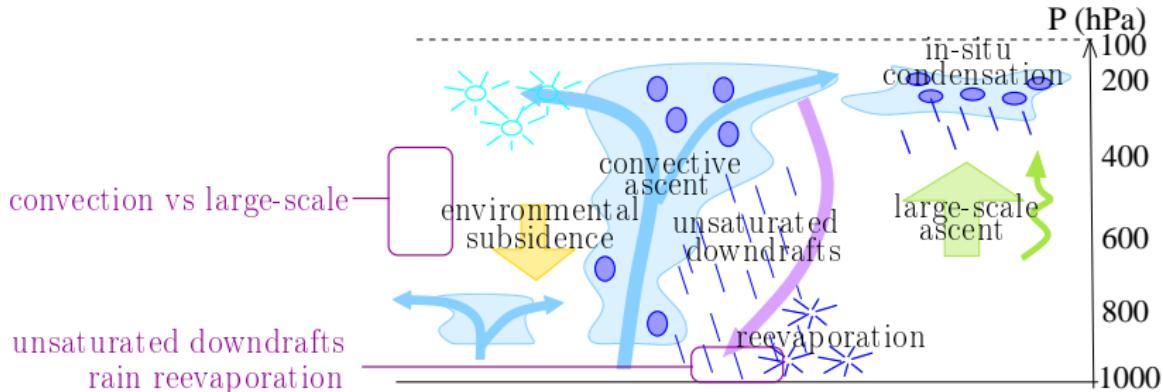


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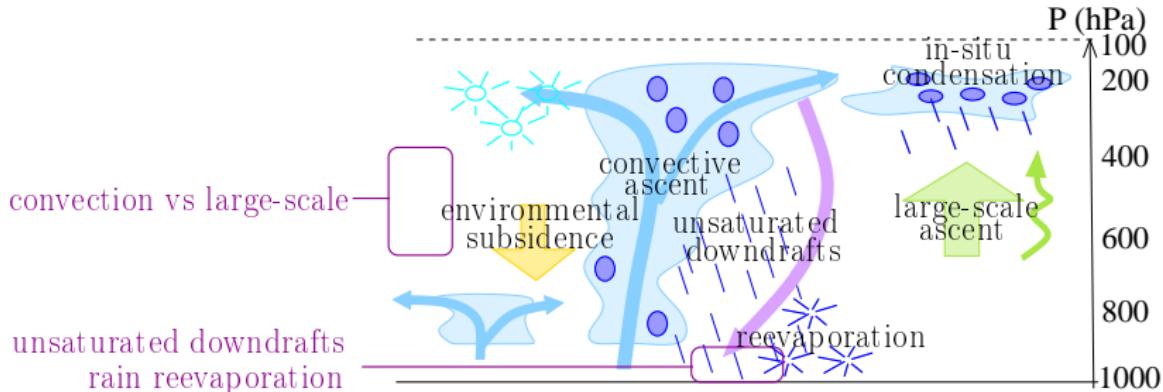
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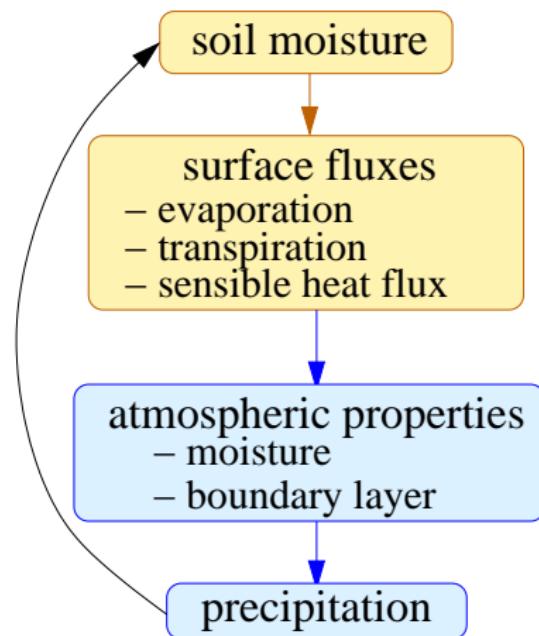
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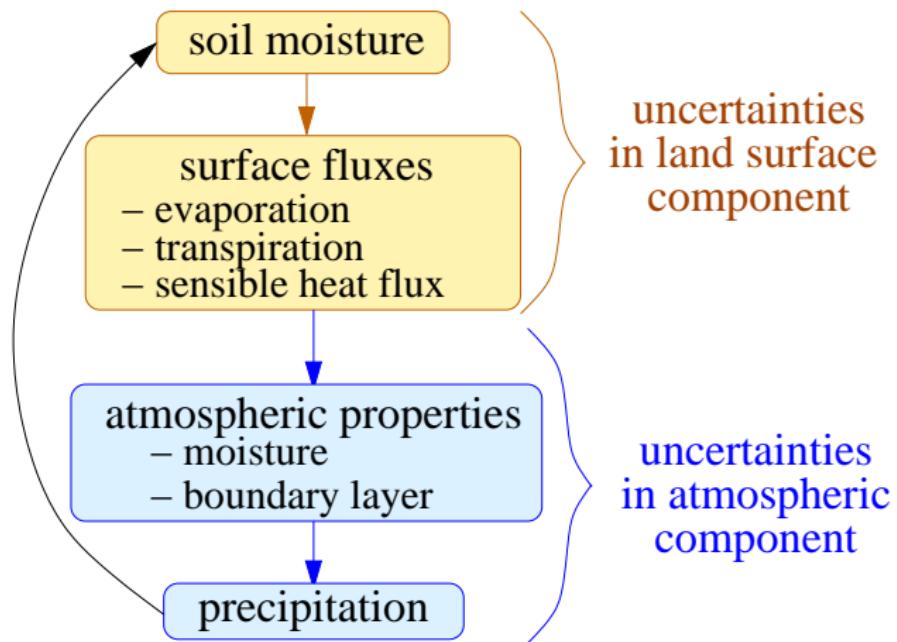
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  - ▶ impact of misrepresentation of convective processes on precipitation changes?

### 3) Land atmosphere feedbacks



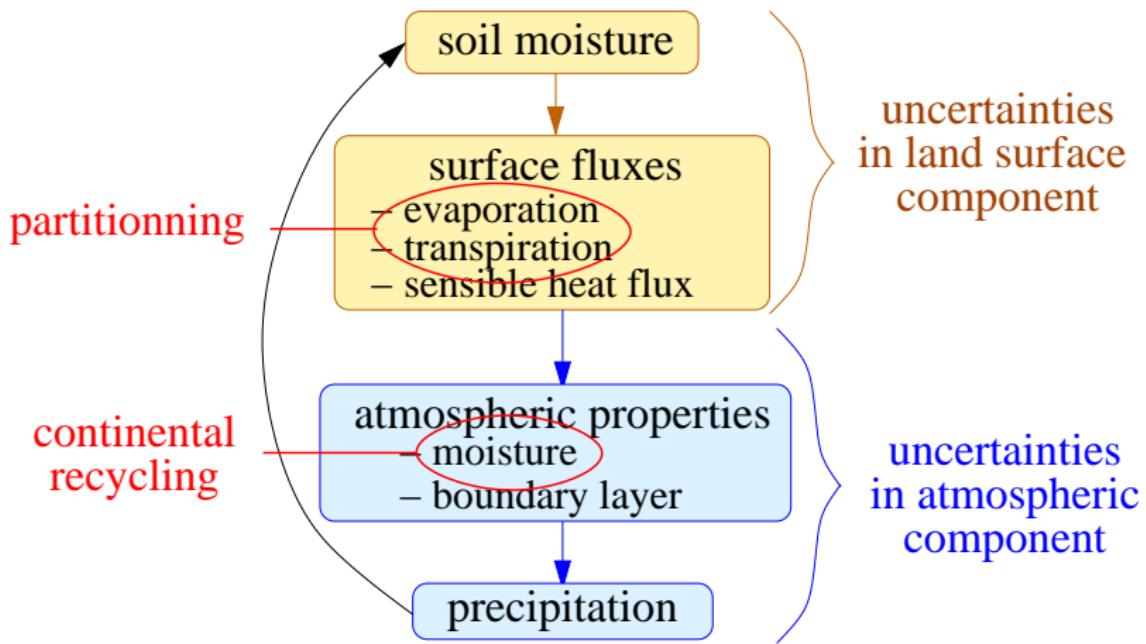
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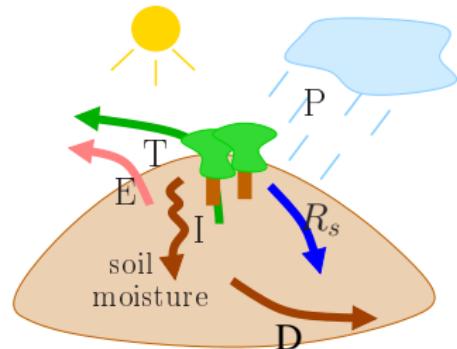
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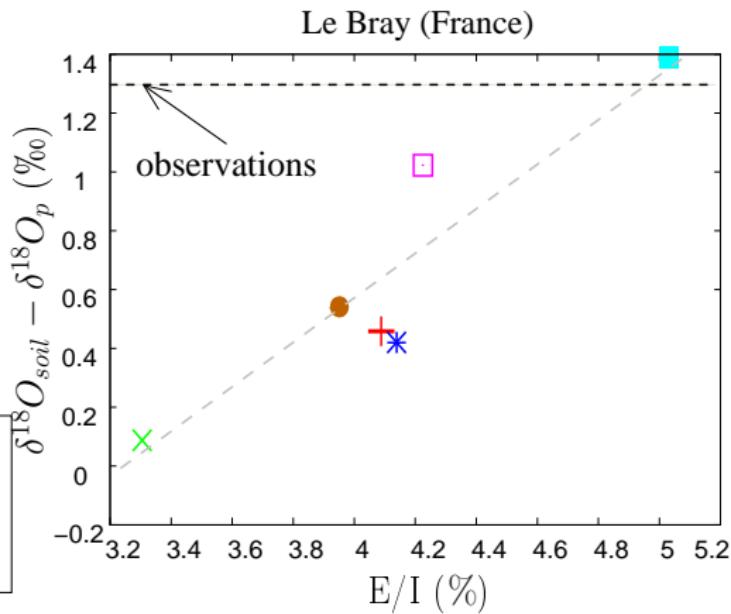


# Evapo-transpiration partitioning

- ▶ ORCHIDEE-iso land surface model (*Risi et al in rev,a*)

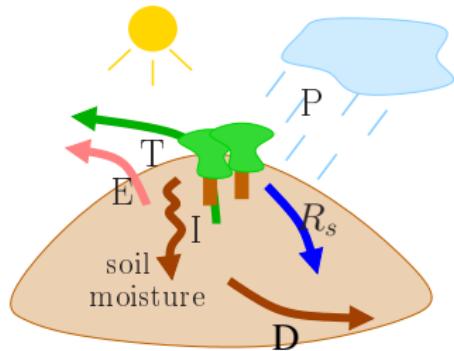


- + control
- ✖ stomatal resistance /5
- no drainage, only surface runoff
- \* soil capacity /2
- less vegetation cover
- root extraction depth /4

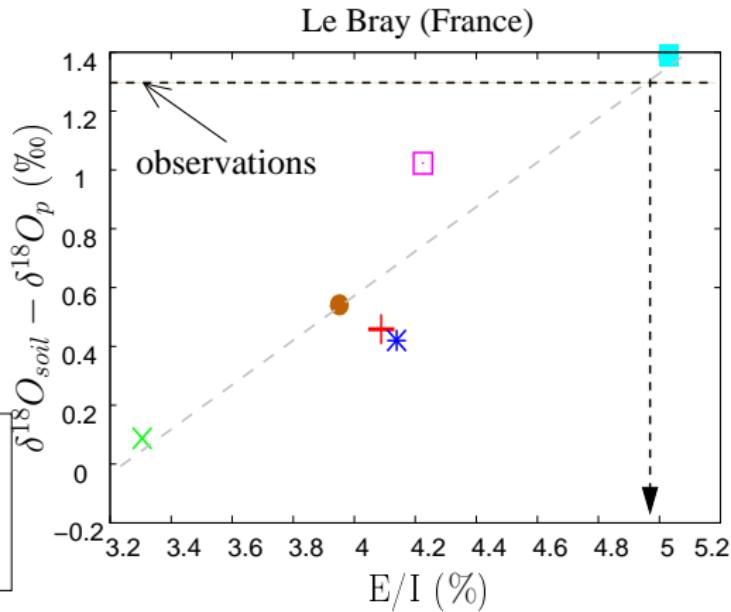


## Evapo-transpiration partitioning

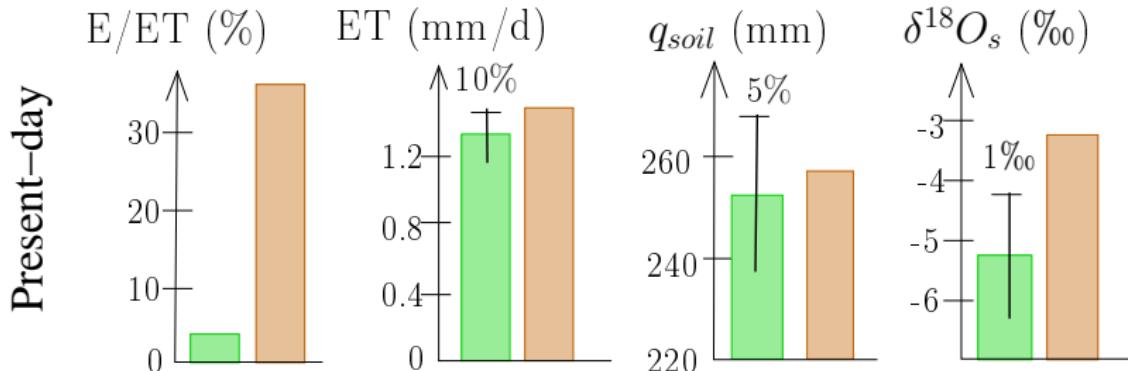
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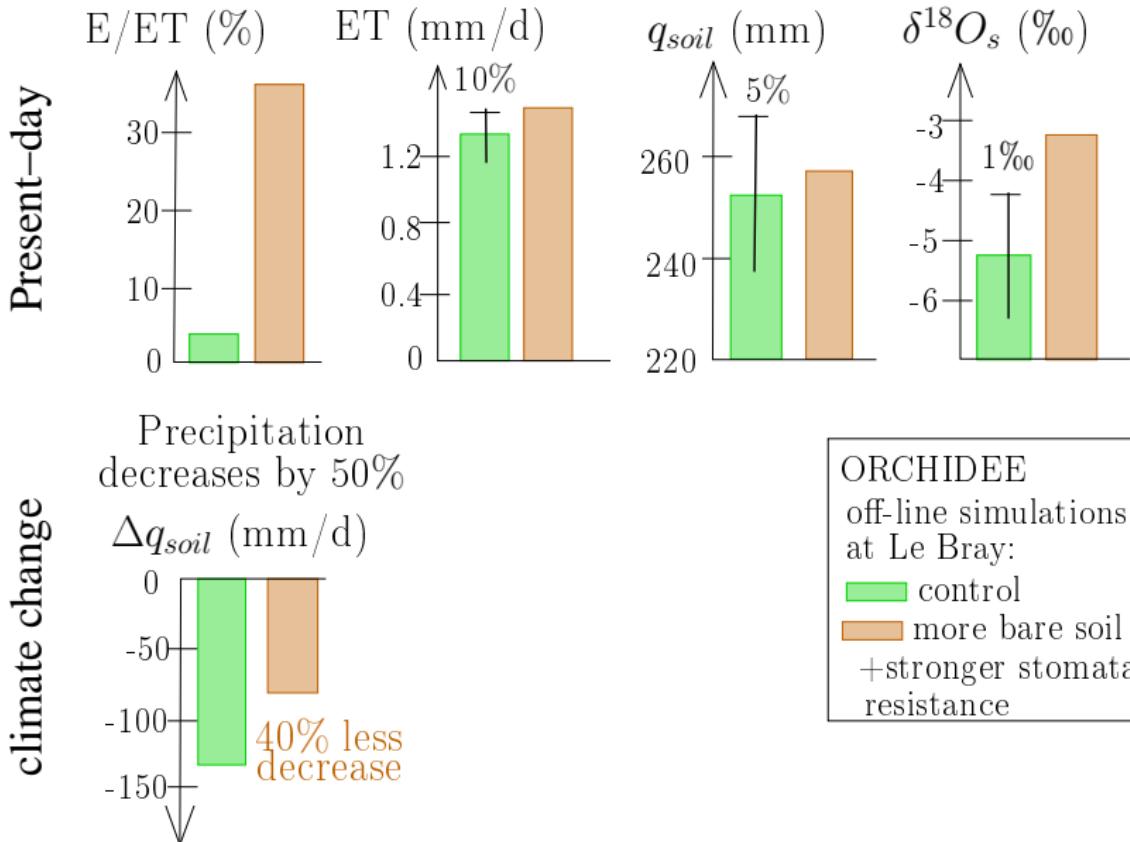


# Does it matter for climate change?

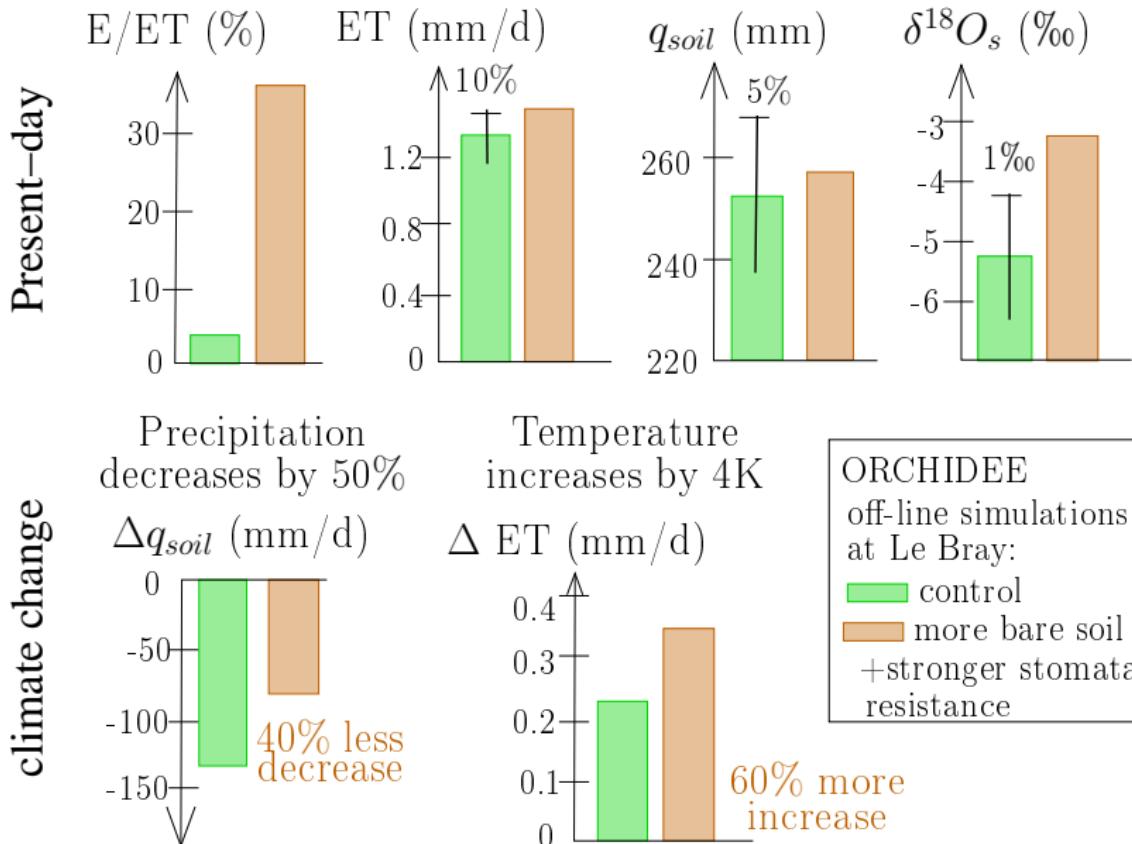


ORCHIDEE  
off-line simulations  
at Le Bray:  
 control  
 more bare soil  
+stronger stomatal  
resistance

# Does it matter for climate change?

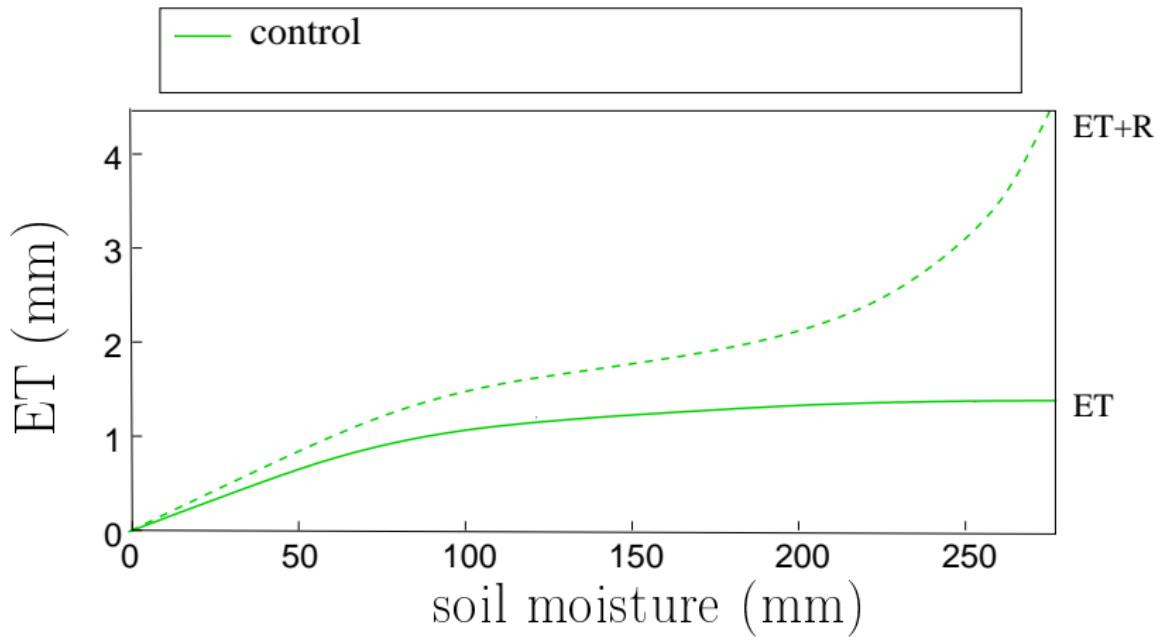


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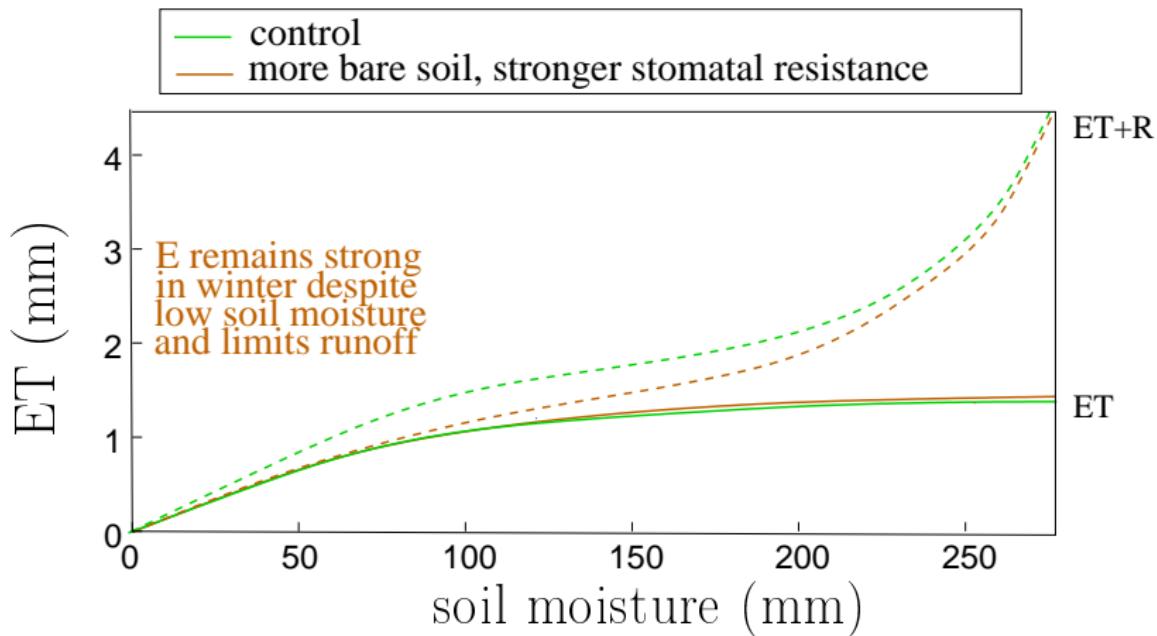
## Impact on response to precipitation

Functional relationships (*Koster and Milly 1996*)



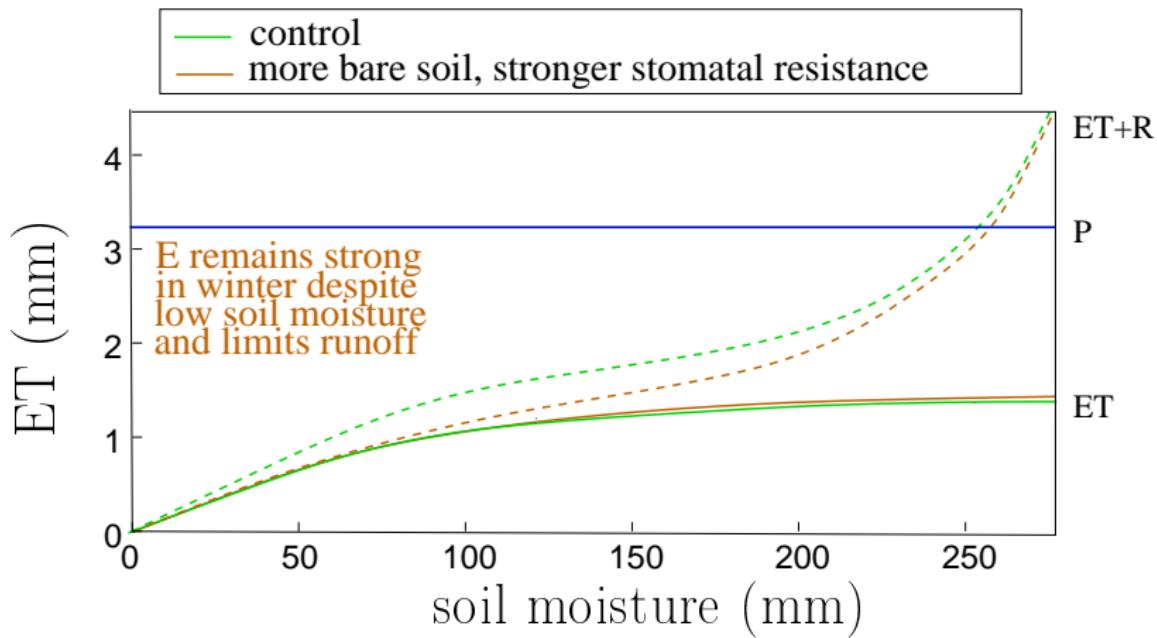
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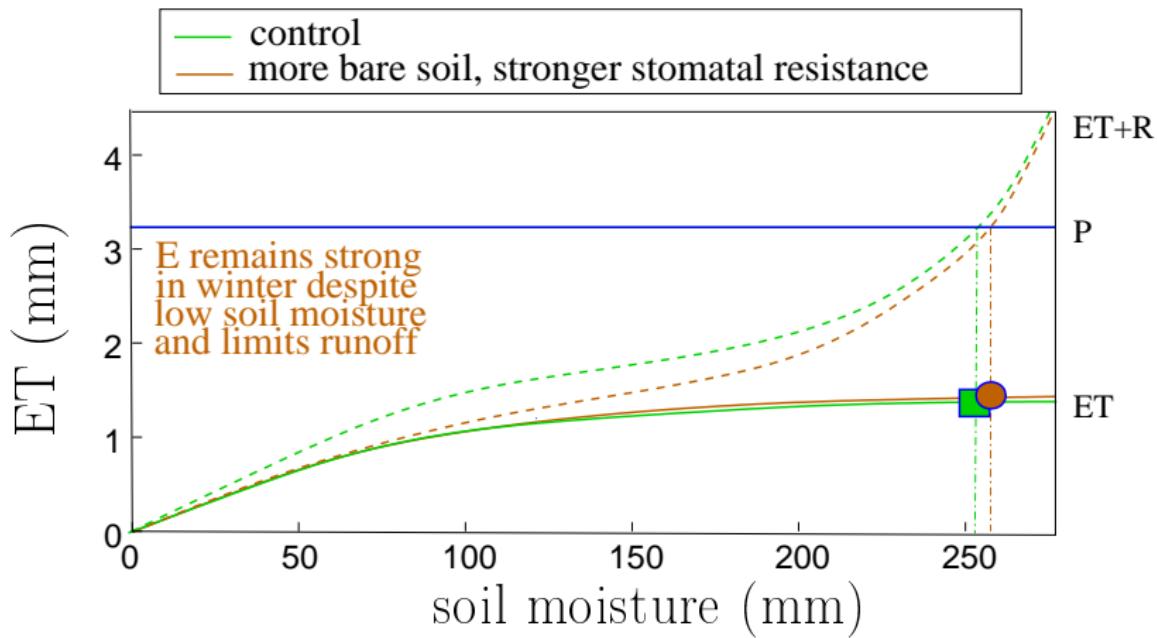
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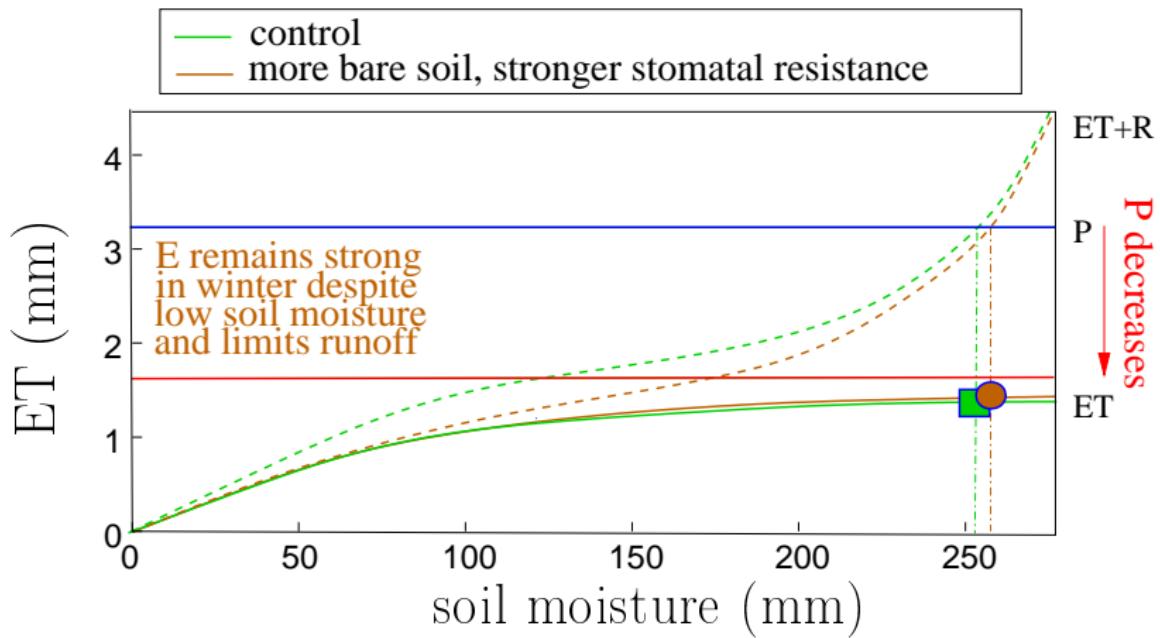
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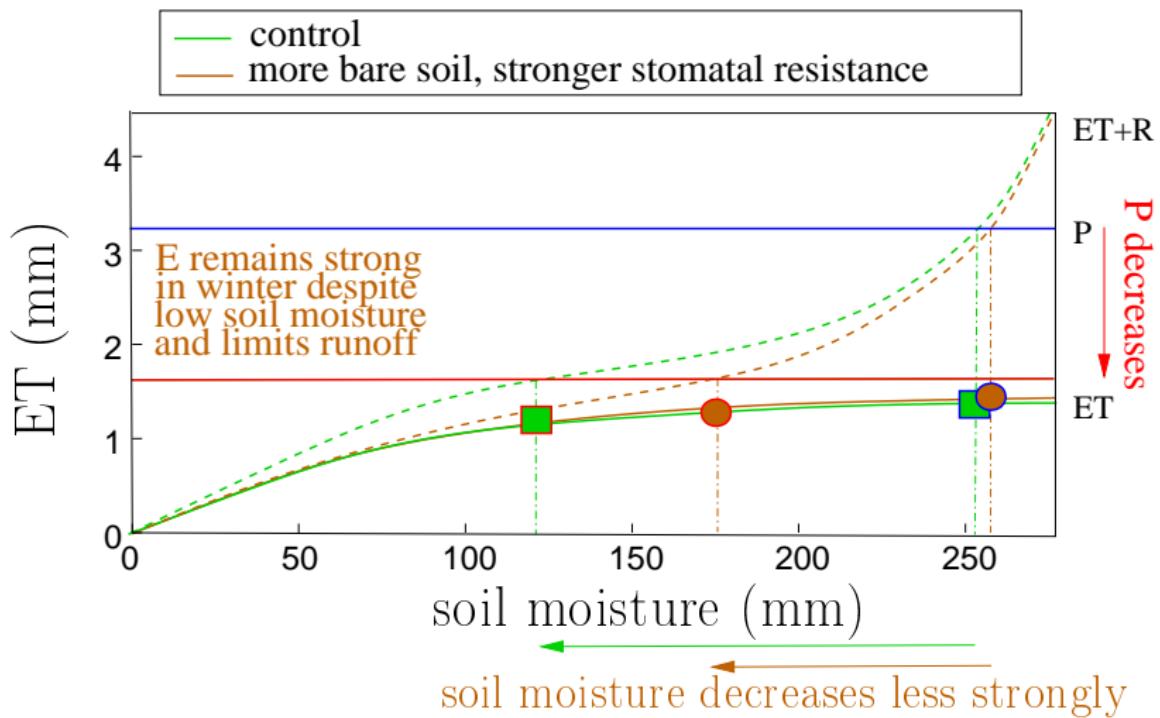
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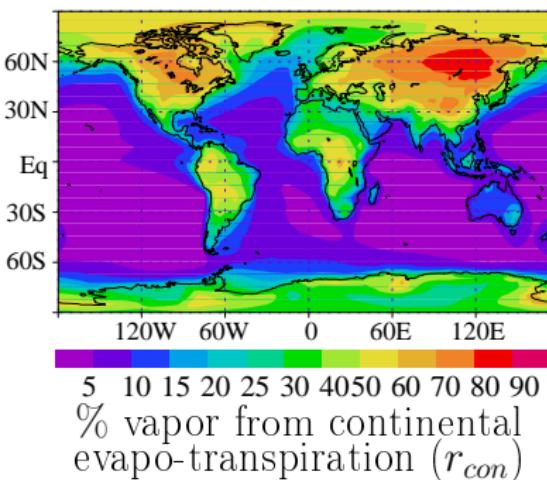
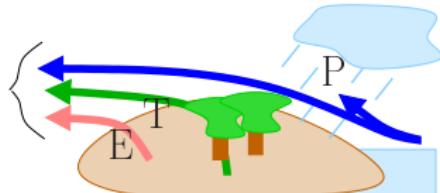
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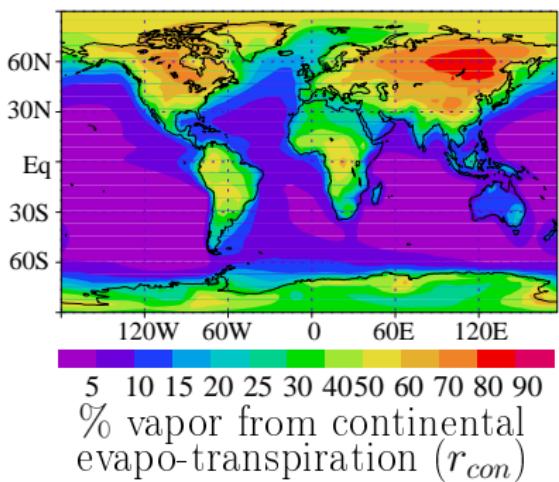
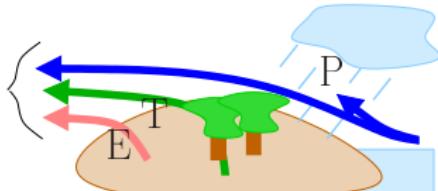
# Isotopic signature of evaporative origin

Water tagging:

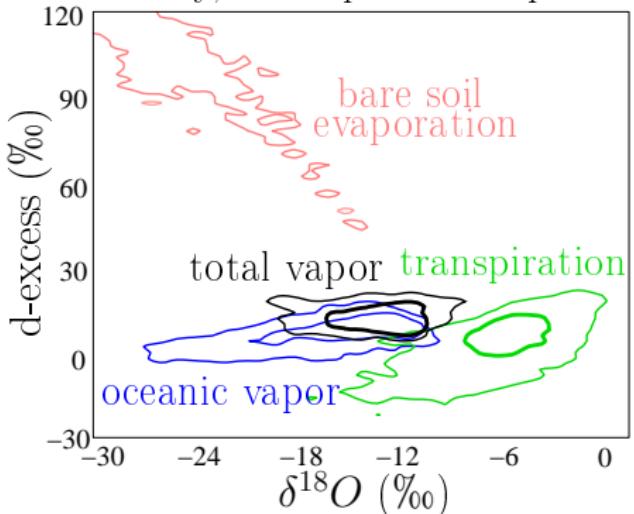


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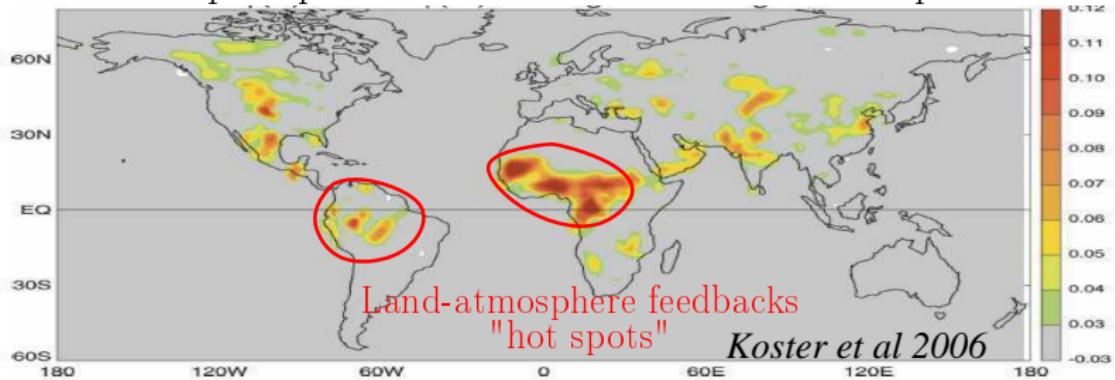


### PDF of vapor composition monthly, all tropical land points



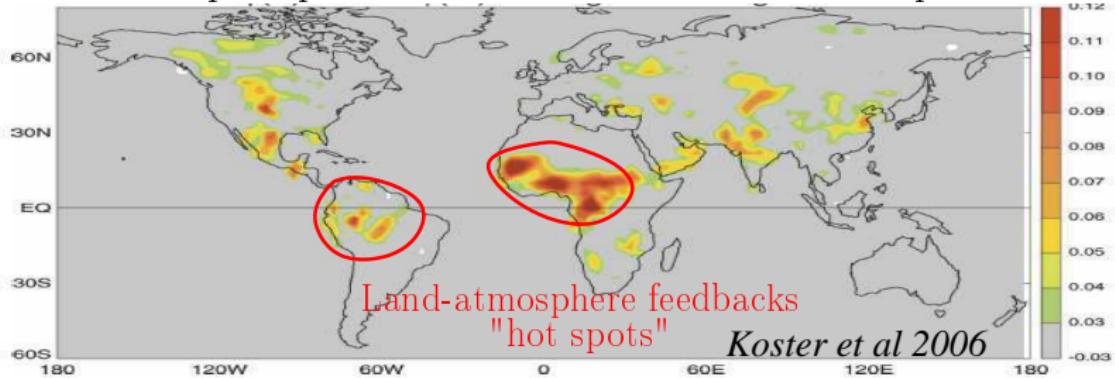
## Water isotopes and continental recycling

decrease in precip variance when soil moisture is prescribed

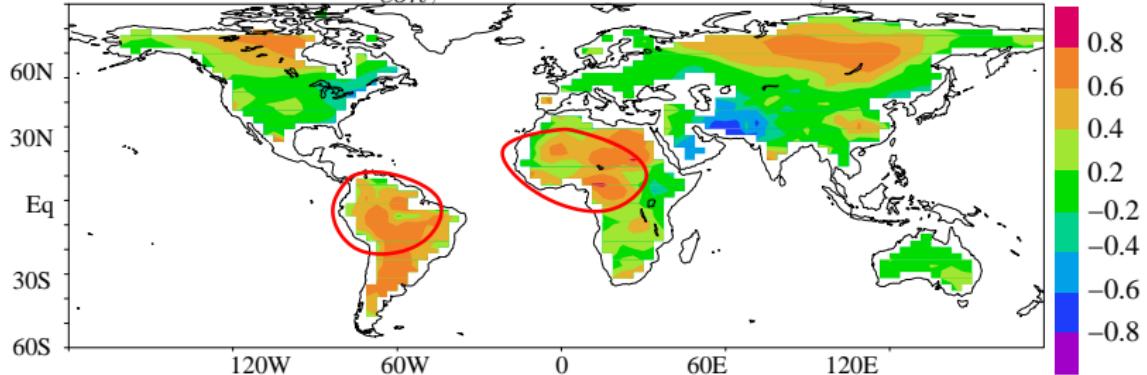


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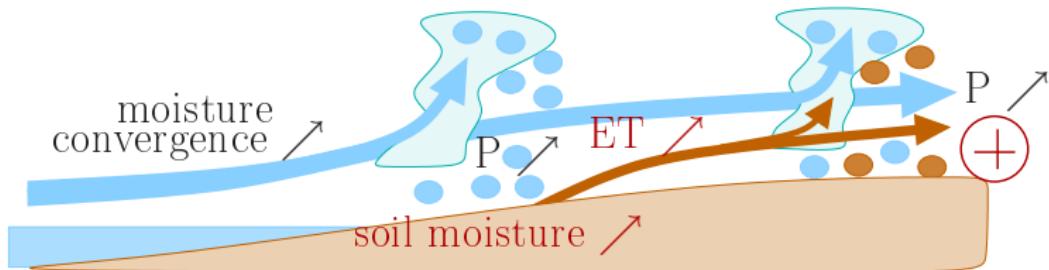
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correlation  $\delta^{18}\text{O}$  -  $r_{con}$ , intra-seasonal scale, annual mean

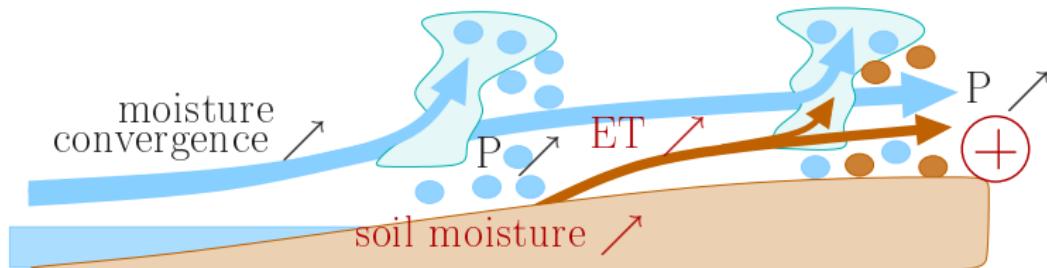


# Diagnosing land-atmosphere feedbacks

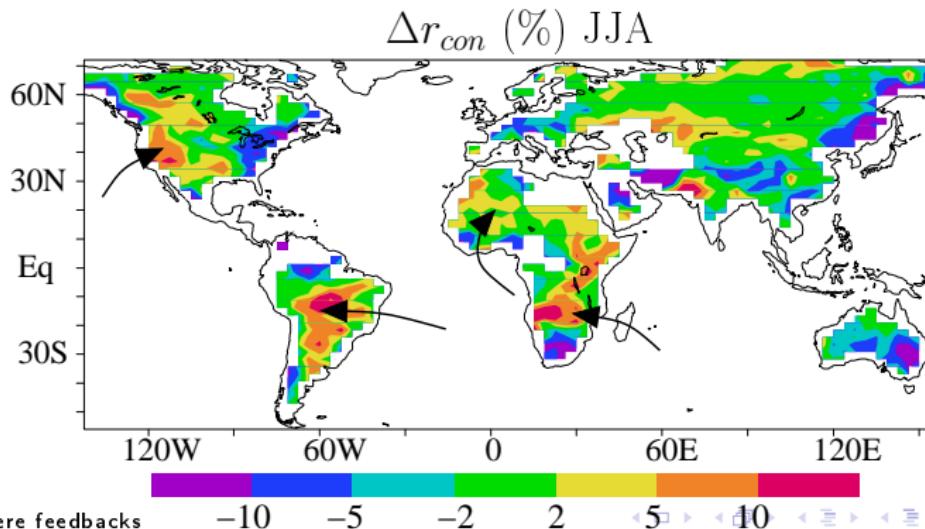


strong precipitation composite minus seasonal average:

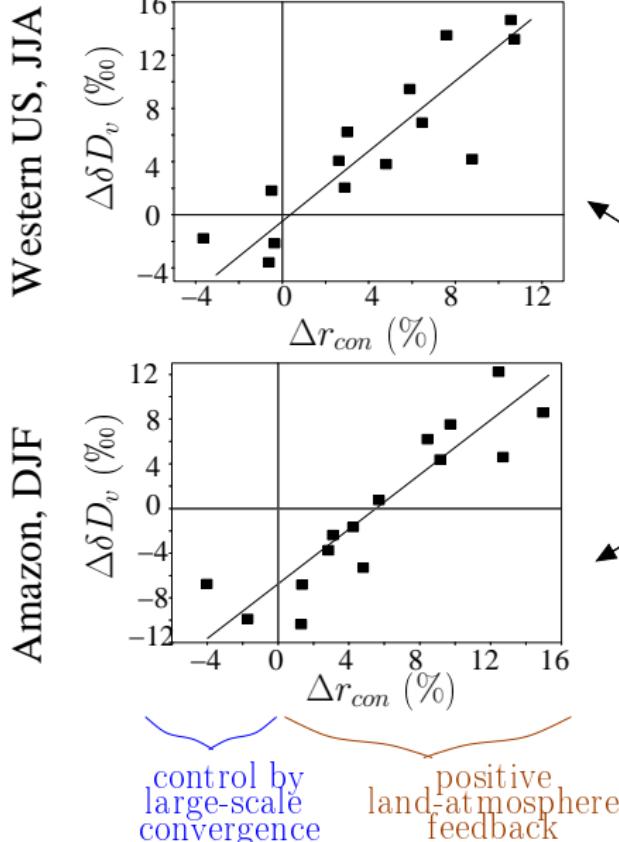
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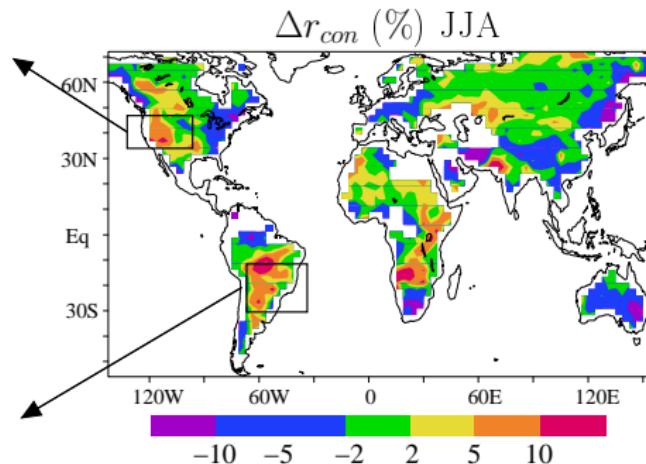
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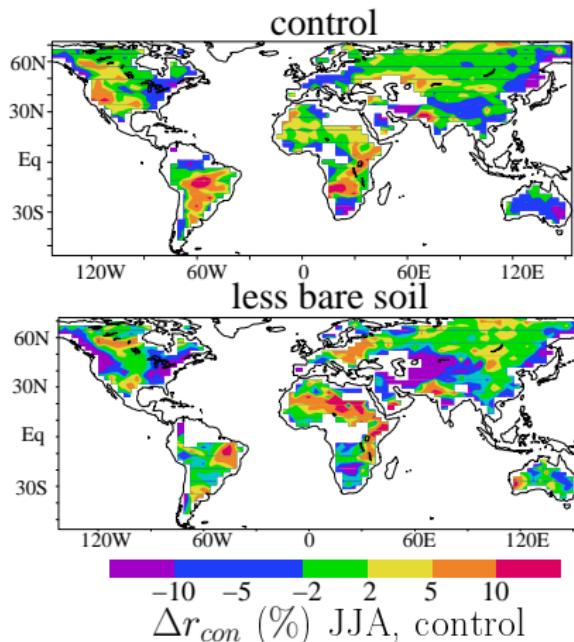
## Isotopic signature of feedbacks



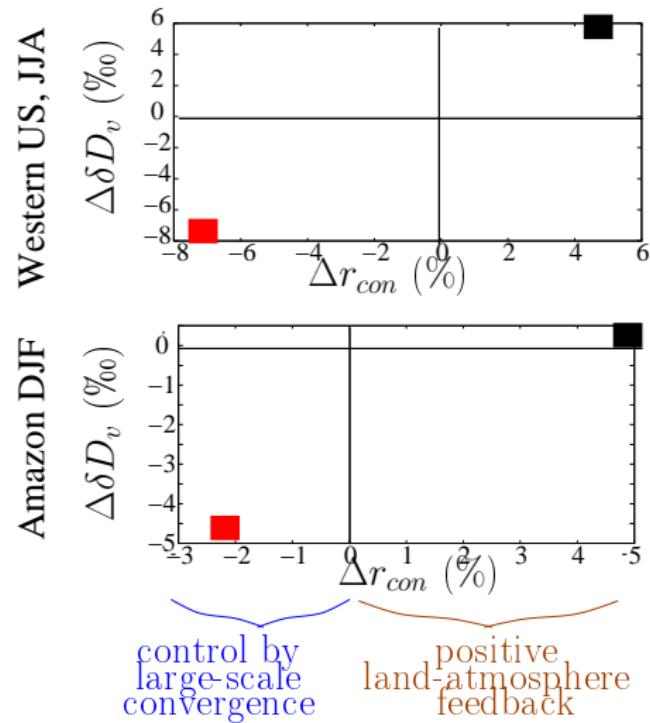
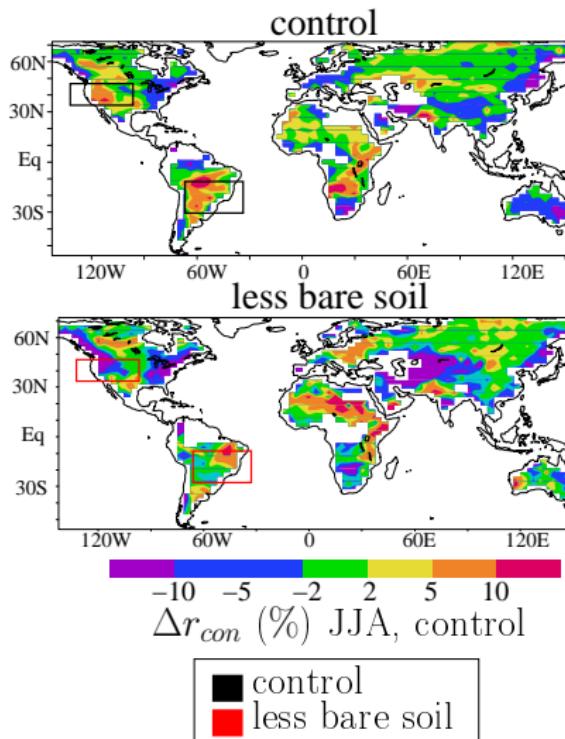
### Strong precipitation composite minus seasonal average:



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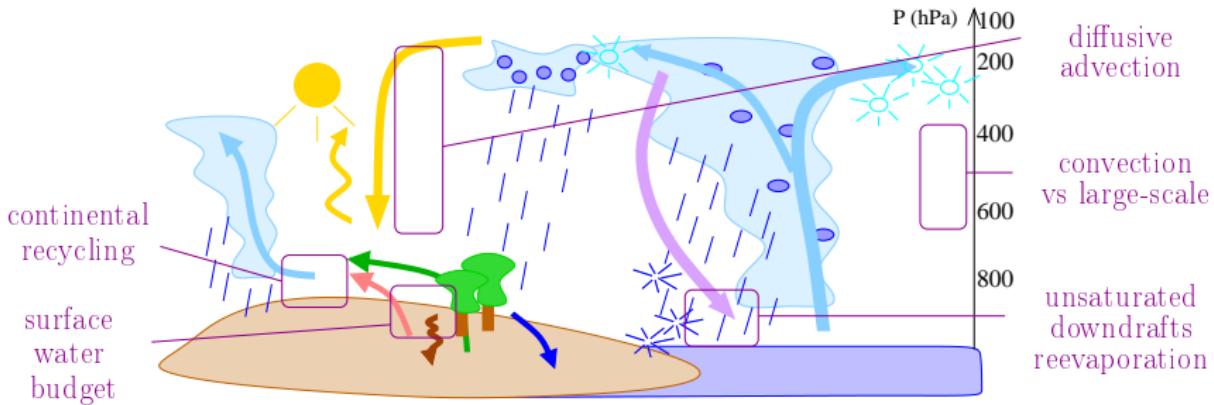
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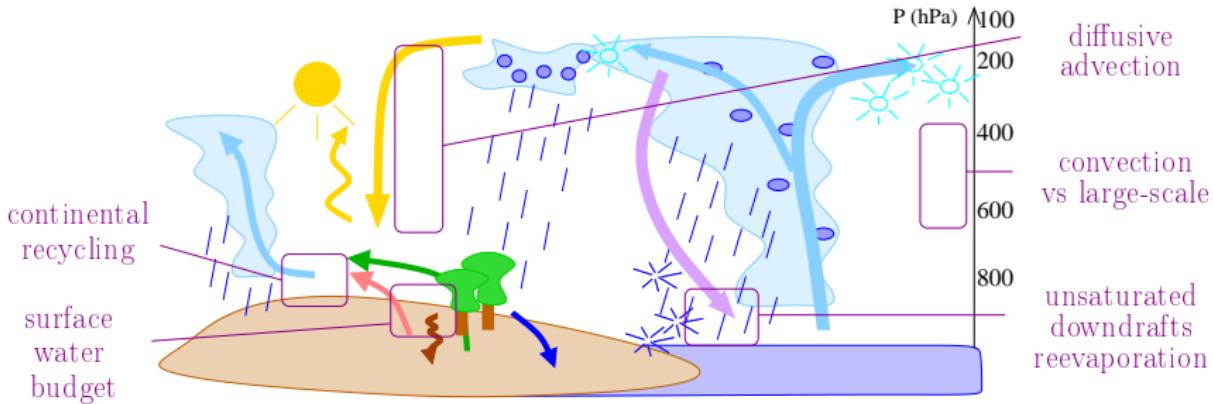
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  - ▶ Relevance for hydrological projections?
    - ▶ Do some processes determine behavior at intra-seasonal scales, and in context of
      - global warming
      - land use change (deforestation, irrigation)

## Conclusion

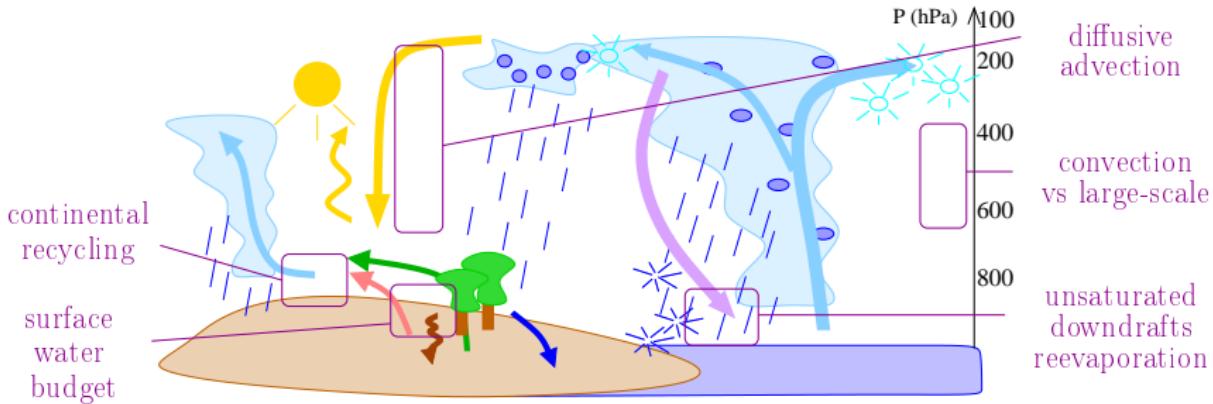


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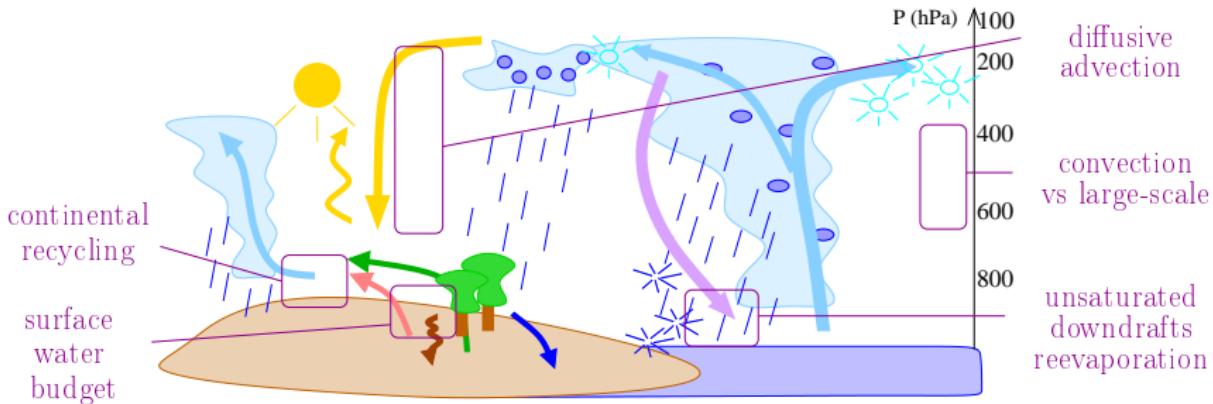
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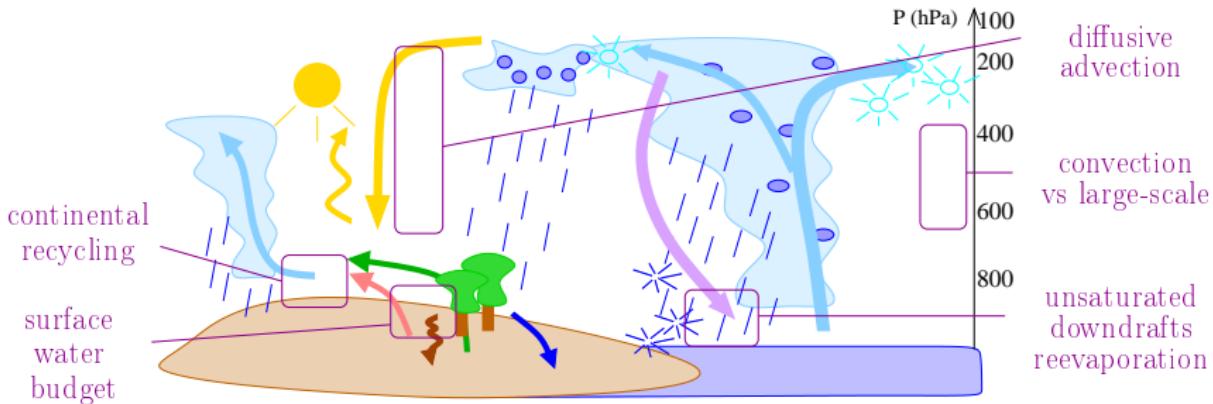
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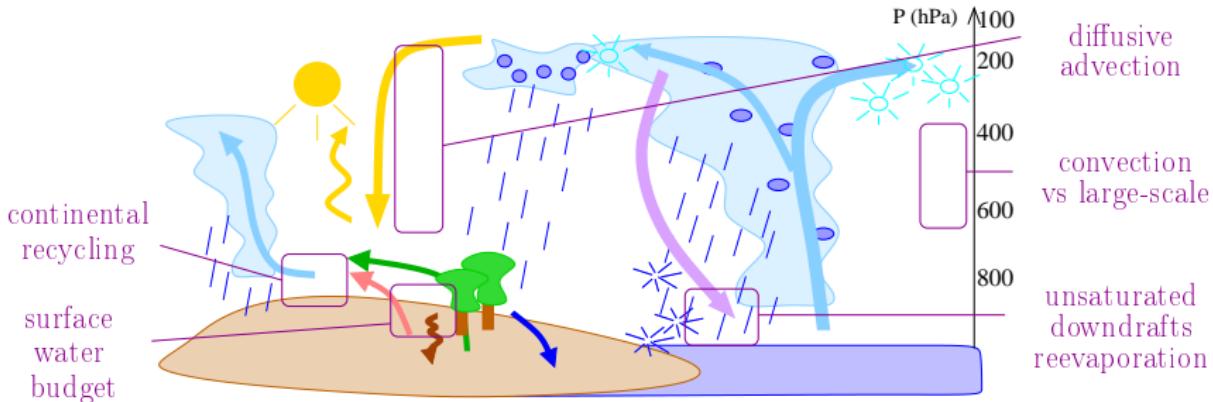
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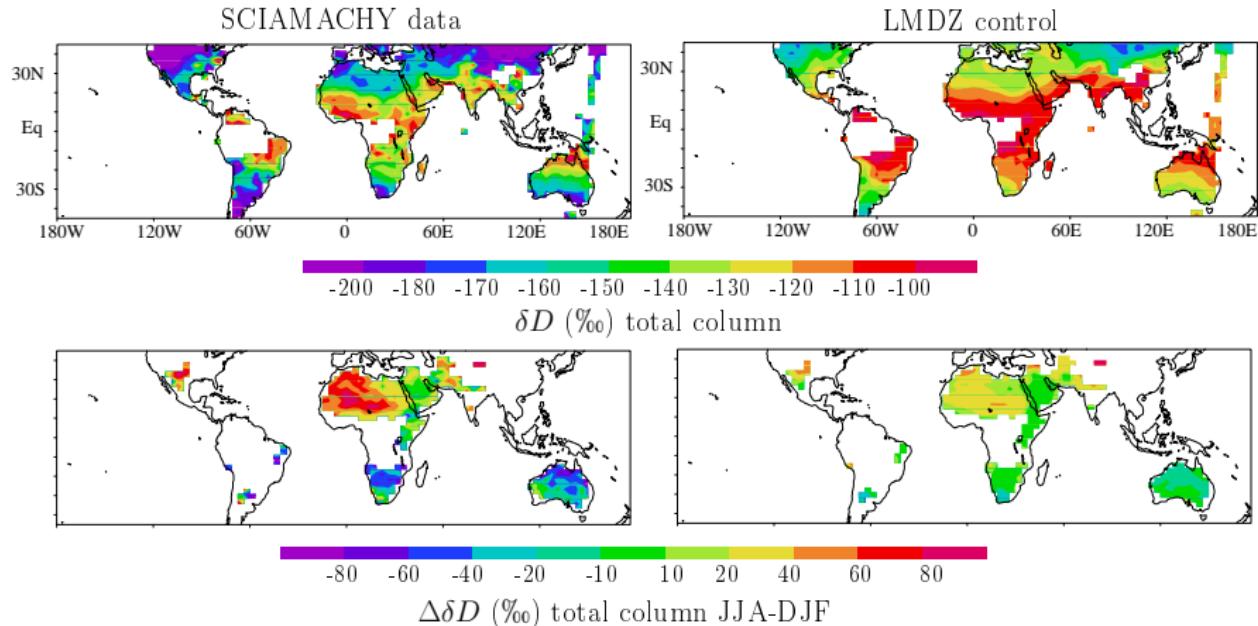
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    - ▶ process/feedbacks studies comparing models behavior for present climate and for projections

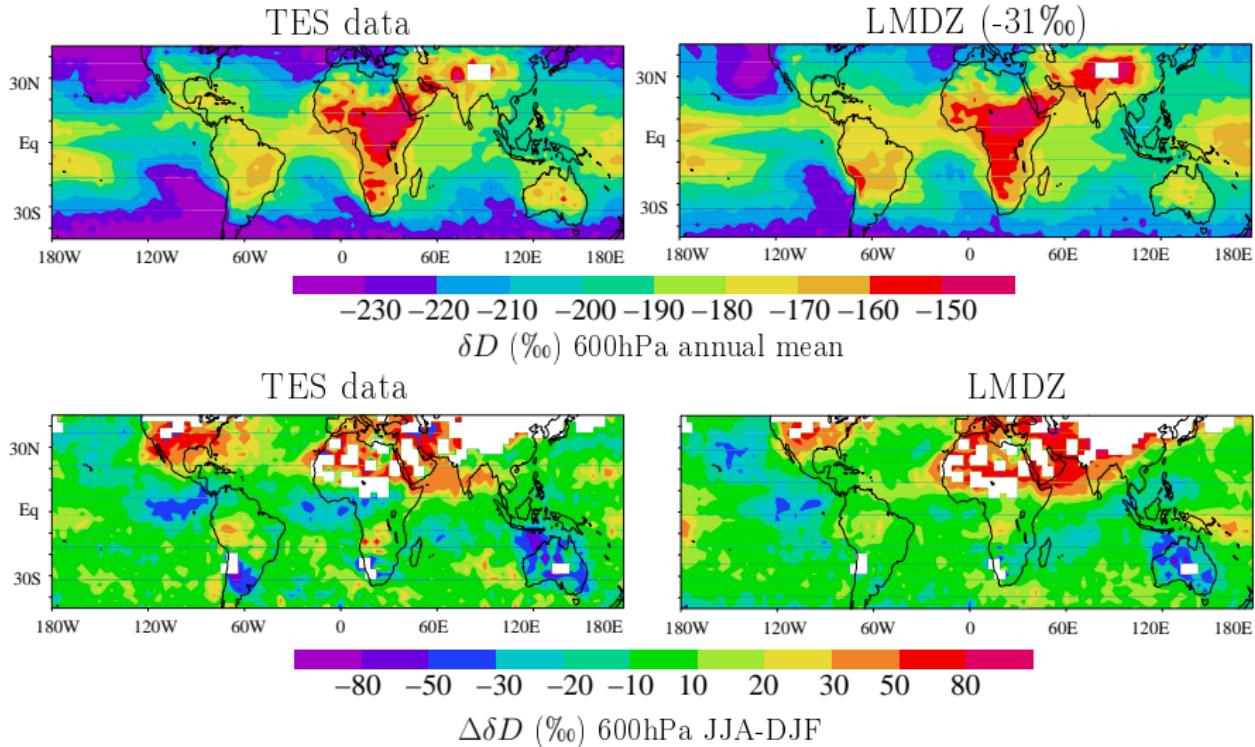
# Supplementary material

# Evaluation against SCIAMACHY



Risi et al in rev,b

# Evaluation against TES



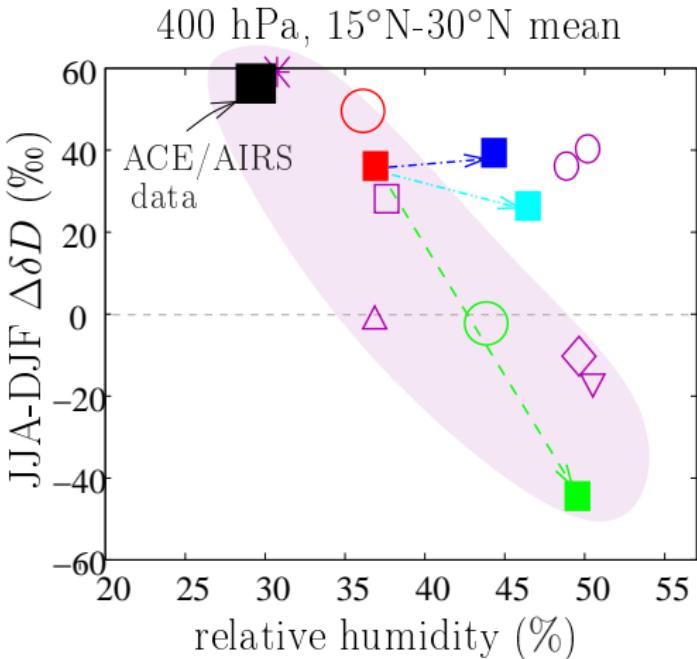
## What causes the moist bias?

Sensitivity tests:  
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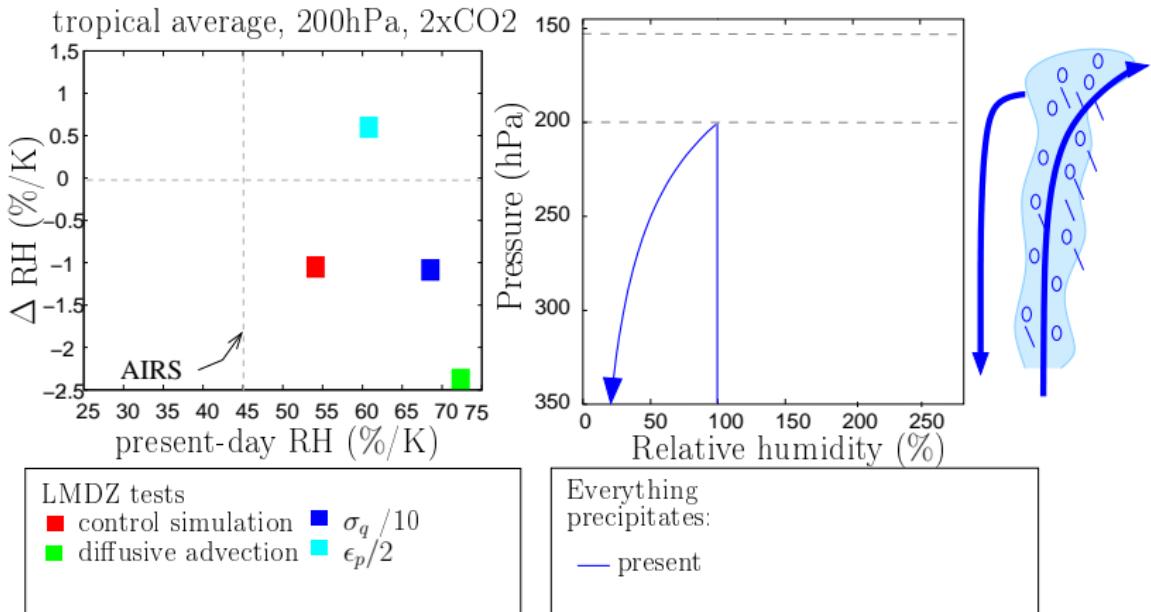
- Control
  - Excessively diffusive vertical advection
  - Excessive condensate detrainement
  - Insufficient in-situ condensation
  - vertical resolution

SWING2 models:

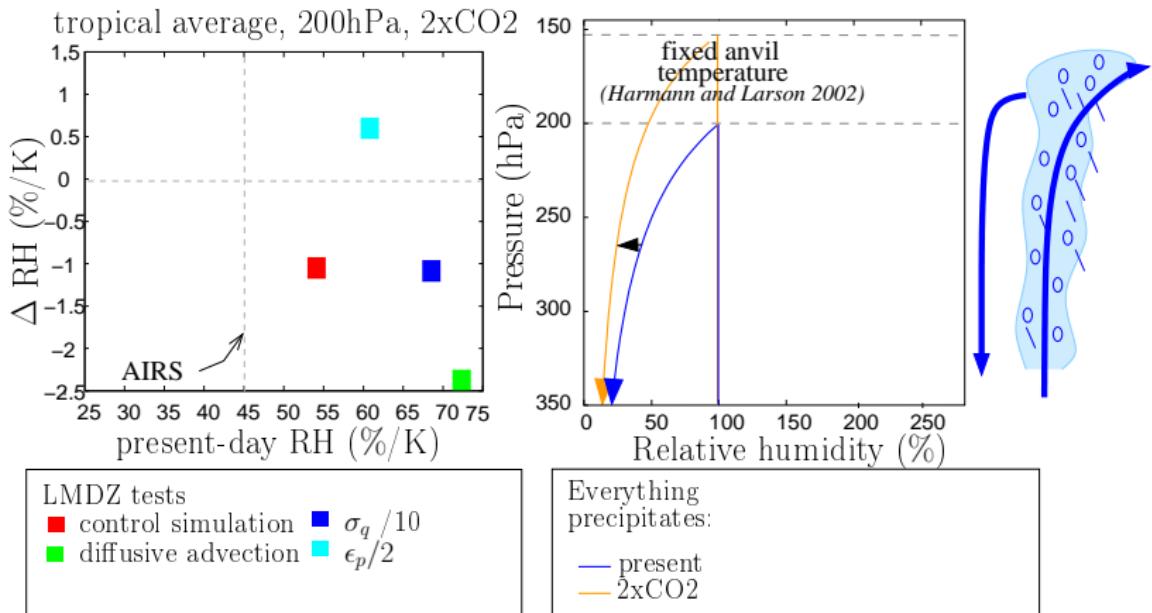
- ECHAM    CAM2  
 MIROC    GISS  
 HadAM    GSM



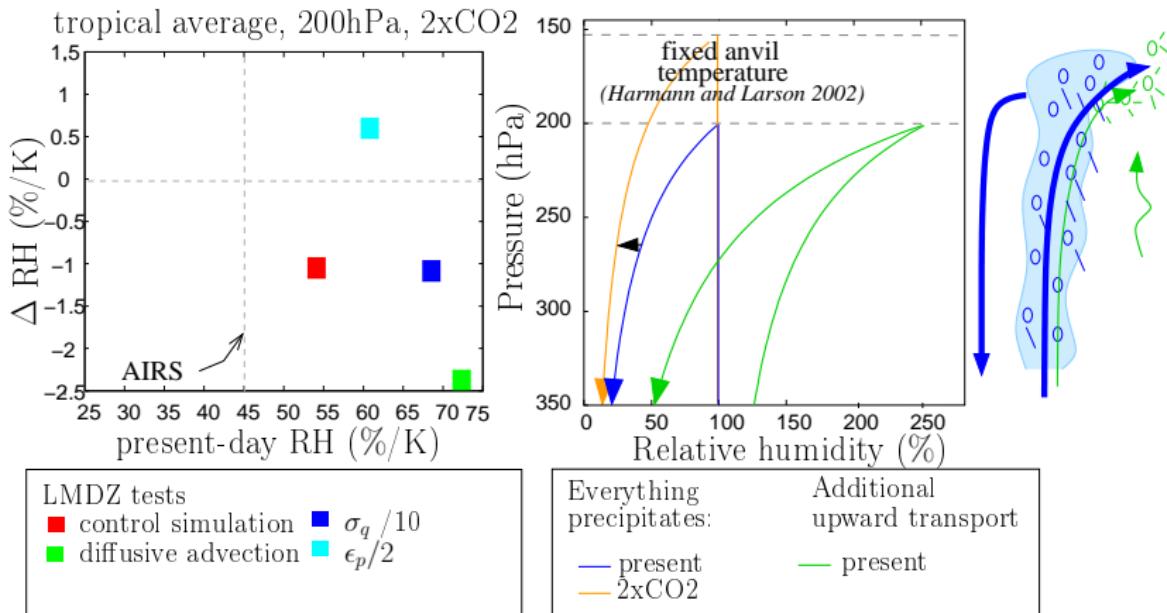
## Consequences on projections



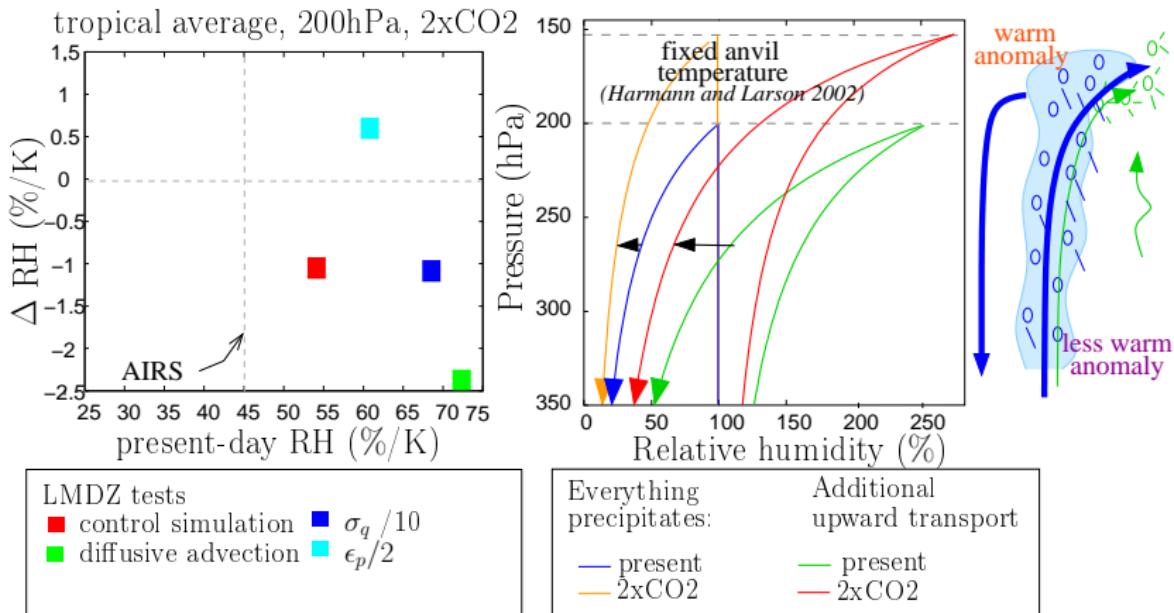
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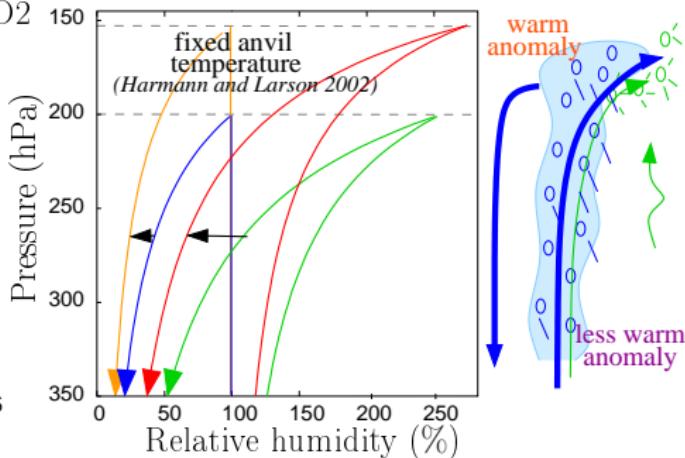
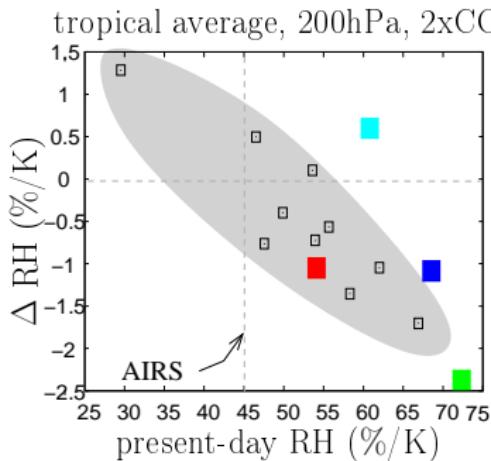
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LMDZ tests

- control simulation    ■  $\sigma_q / 10$
- diffusive advection    ■  $\epsilon_p / 2$
- CMIP3 models

Everything  
precipitates:

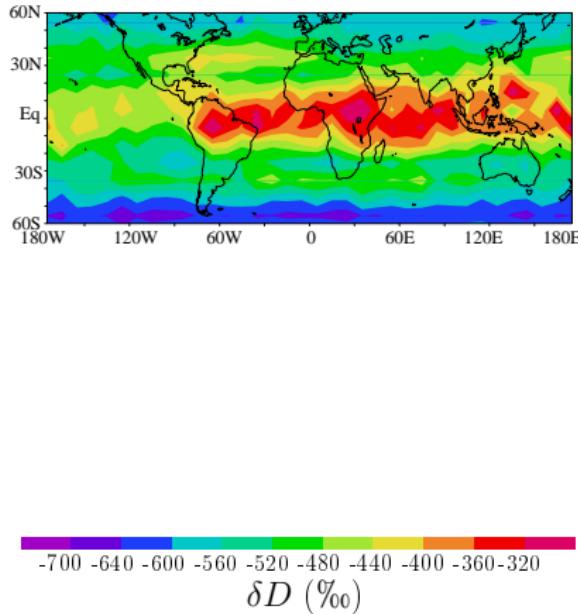
- present
- 2xCO<sub>2</sub>

Additional  
upward transport

- present
- 2xCO<sub>2</sub>

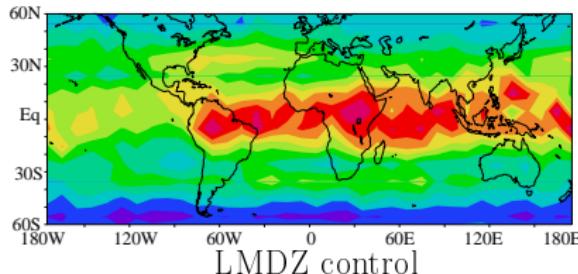
## Upper troposphere detrainment

### MIPAS data at 200hPa, annual

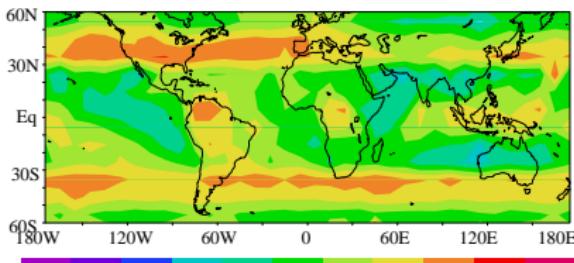


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MIPAS data at 200hPa, annual



LMDZ control

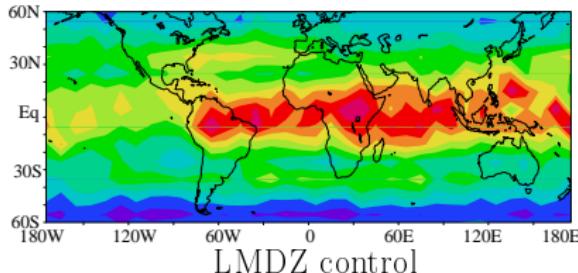


-700 -640 -600 -560 -520 -480 -440 -400 -360 -320

$\delta D$  (%)

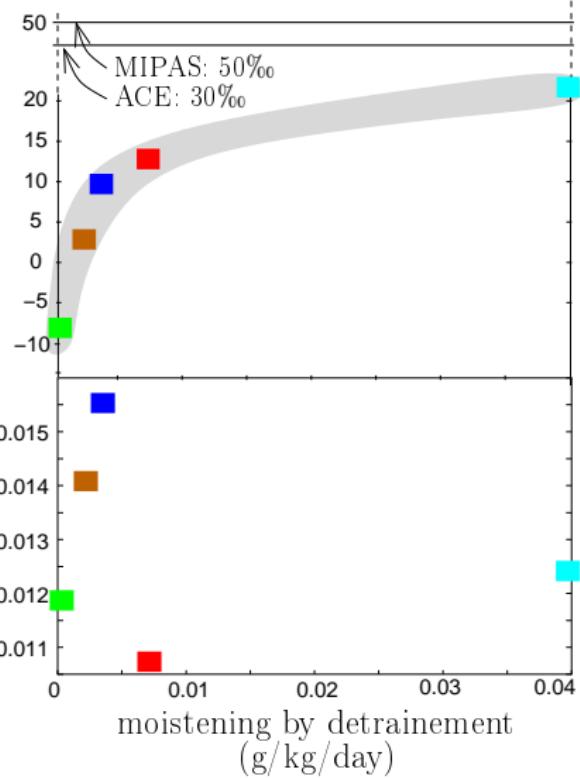
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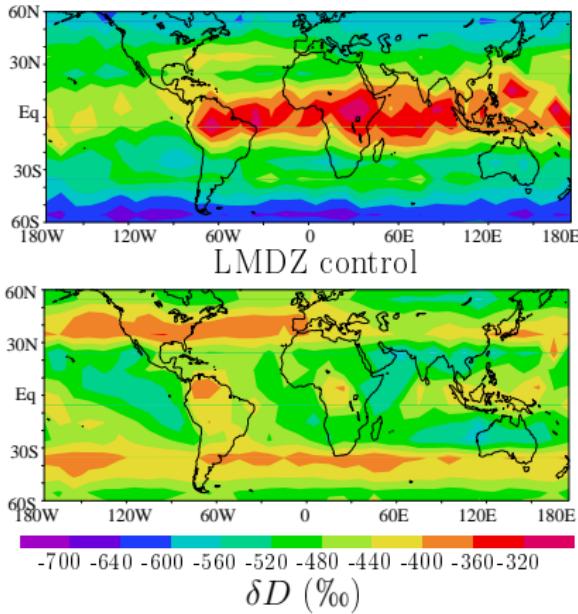
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Difference 15°S-15°N minus  
30°S-30°N at 200hPa



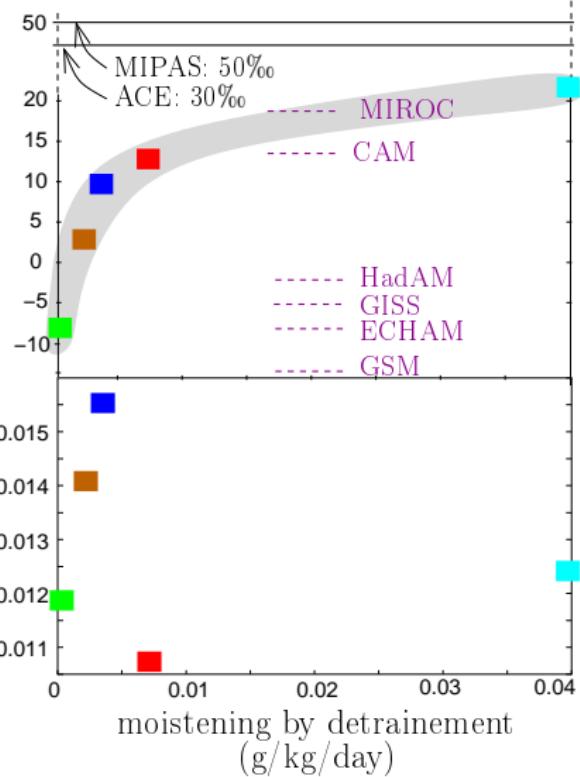
## Upper troposphere detrainment

## MIPAS data at 200hPa, annual

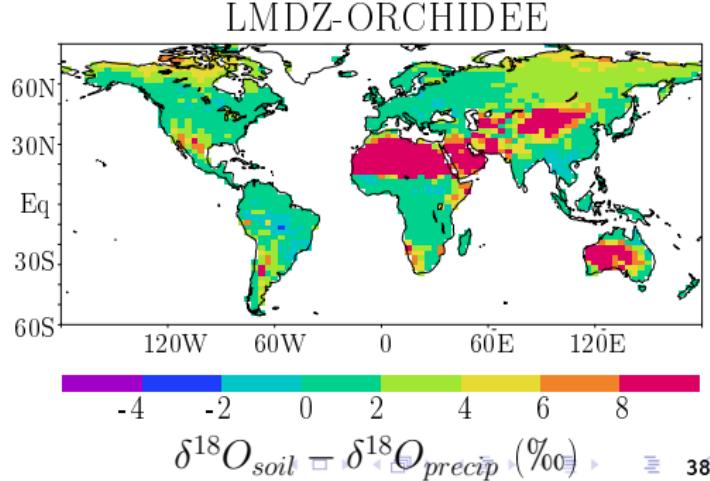
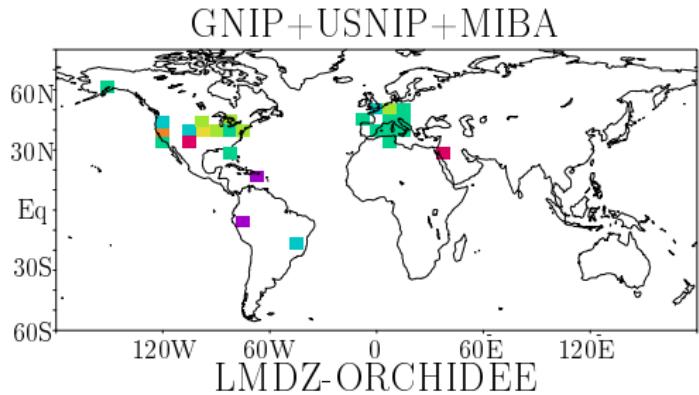
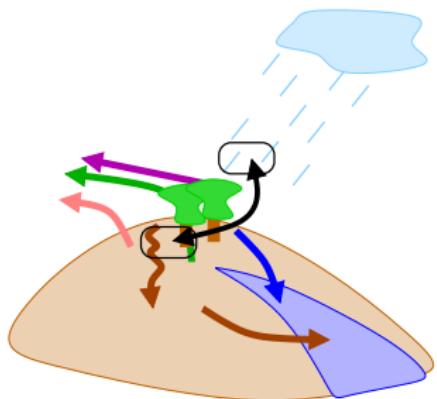


- control
  - vertical advection more diffusive
  - stronger condensate detrainment
  - less large-scale condensation
  - less large-scale precipitation

Difference 15°S-15°N minus  
30°S-30°N at 200hPa



# Soil water isotopes



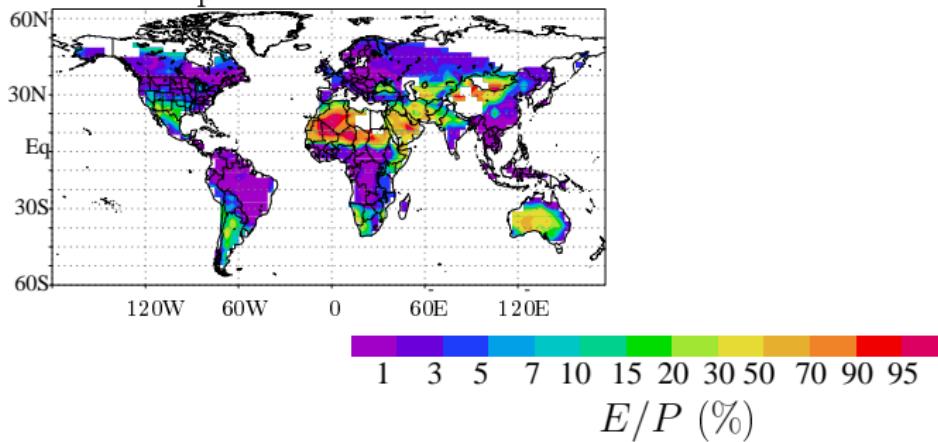
# Estimating evapotranspiration partitioning

$$\frac{\delta^{18}\text{O}_{\text{soil}} - \delta^{18}\text{O}_p}{\delta^{18}\text{O}_v}$$

RH, T



estimated from simulated  
isotopic "measurements"



# Estimating evapotranspiration partitioning

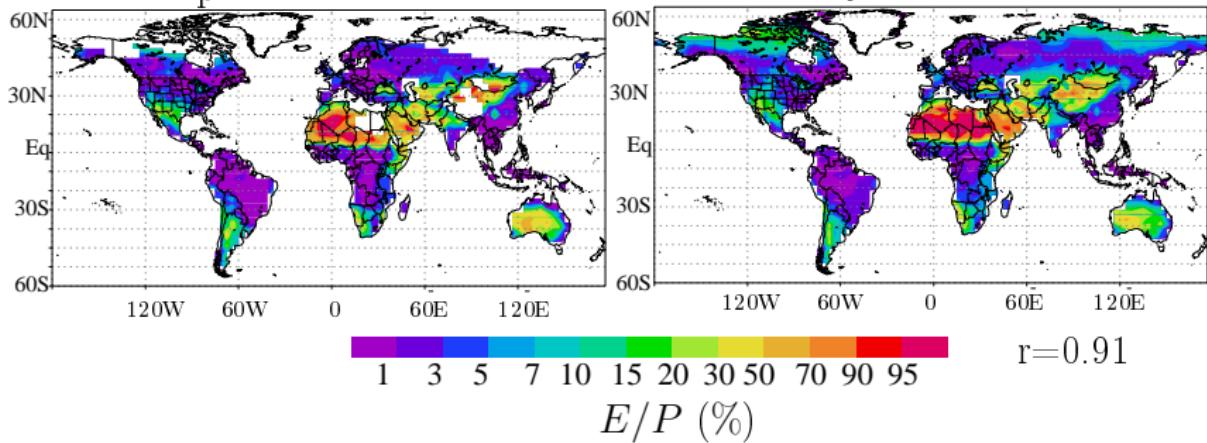
$$\frac{\delta^{18}\text{O}_{\text{soil}} - \delta^{18}\text{O}_p}{\delta^{18}\text{O}_v}$$

RH, T

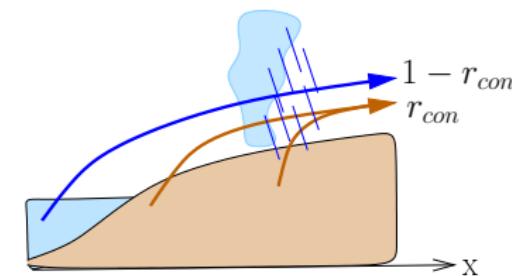


estimated from simulated  
isotopic "measurements"

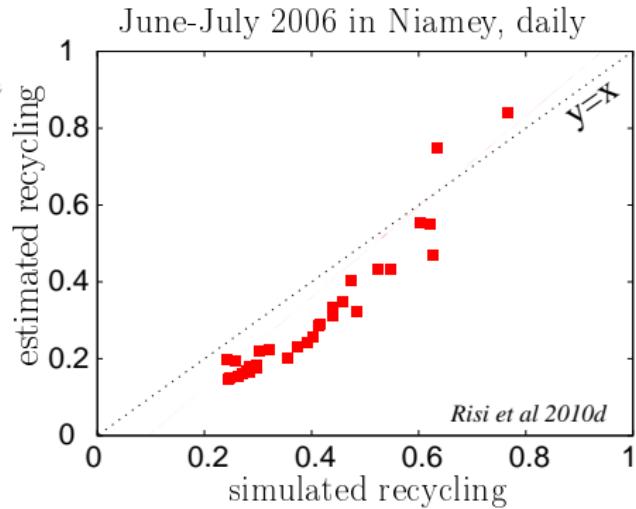
simulated by LMDZ-ORCHIDEE



# Estimating continental recycling

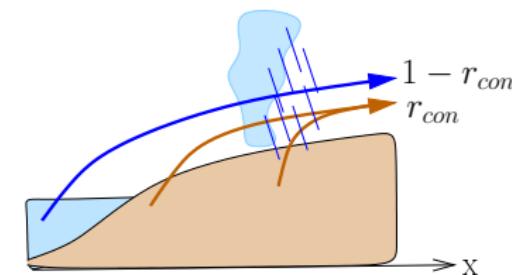


- $\frac{d\delta_{voce}}{dx}$  known

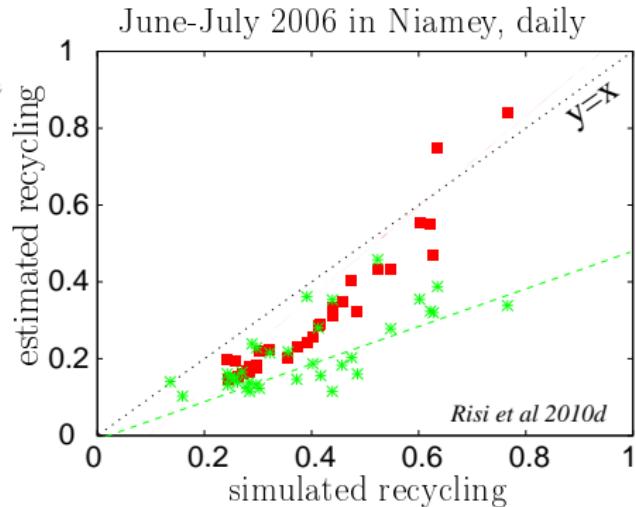


$$d \left( \frac{r_{con}}{1 - r_{con}} \right) / dx = \frac{d\delta_v/dx - d\delta_{voce}/dx}{\delta_p - \delta_v}$$

# Estimating continental recycling

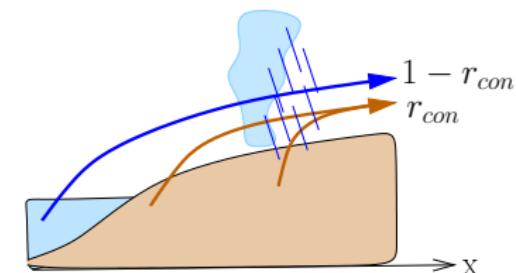


- $\frac{d\delta_{voc}}{dx}$  known
- $\frac{d\delta_v}{dx}$  depends linearly on precipitation

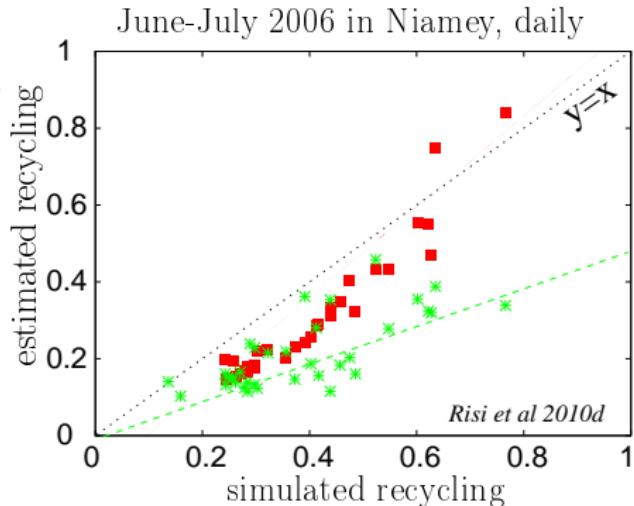


$$d \left( \frac{r_{con}}{1 - r_{con}} \right) / dx = \frac{d\delta_v/dx - d\delta_{voc}/dx}{\delta_p - \delta_v}$$

# Estimating continental recycling



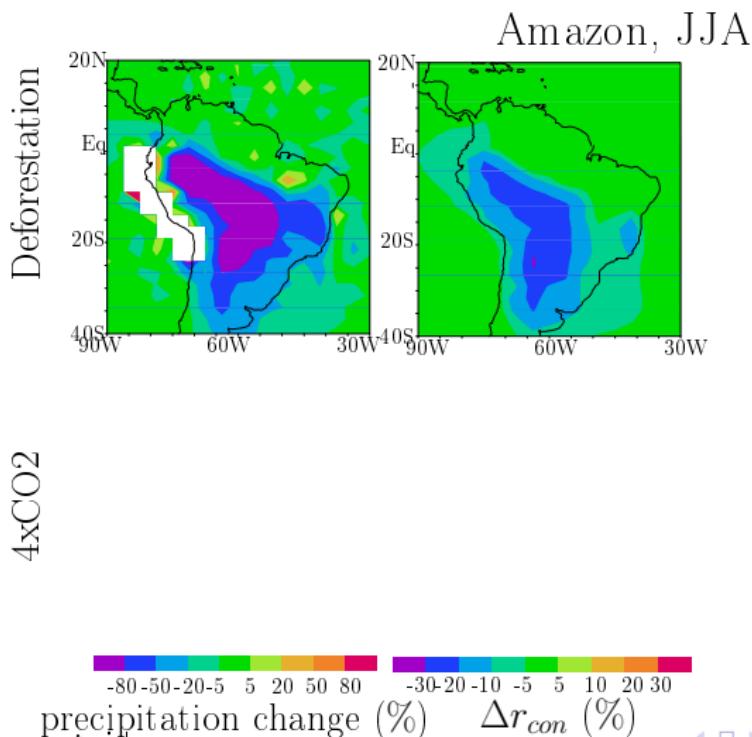
- $\frac{d\delta_{voce}}{dx}$  known
- $\frac{d\delta_{voce}}{dx}$  depends linearly on precipitation



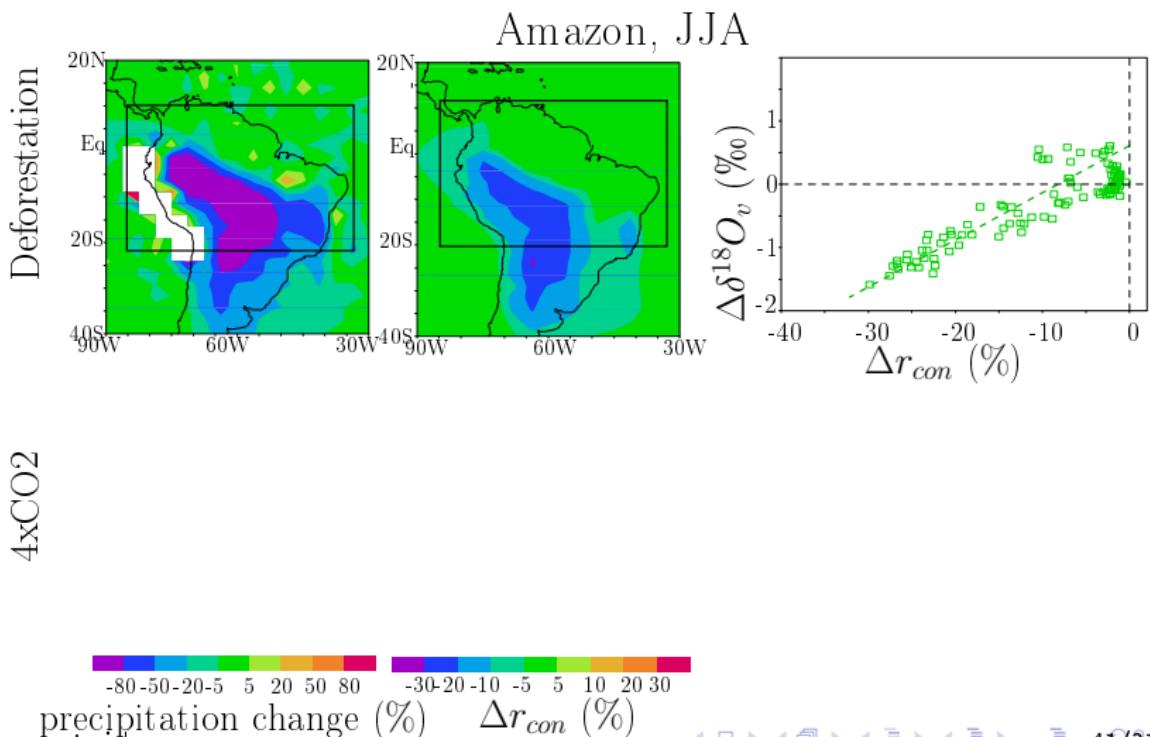
$$d \left( \frac{r_{con}}{1 - r_{con}} \right) / dx = \frac{d\delta_v/dx - d\delta_{voce}/dx}{\delta_p - \delta_v}$$

- ▶ Main limitation in using vapor isotopic measurements for continental recycling: understanding atmospheric controls

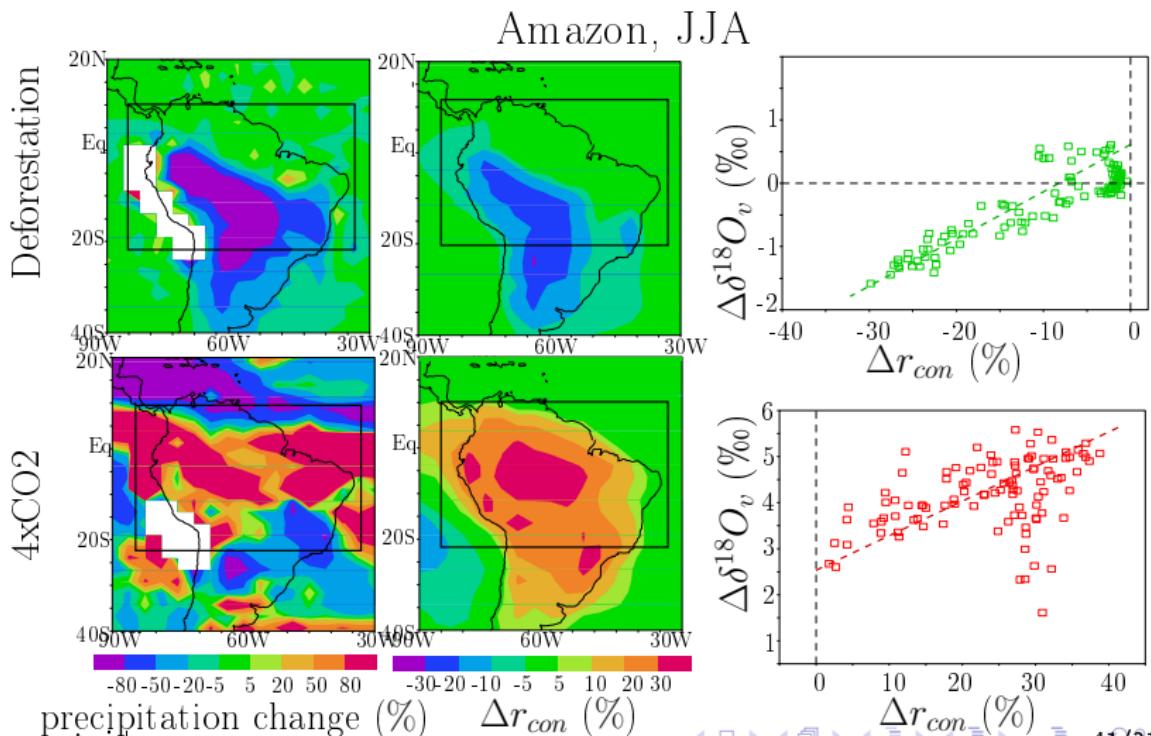
# Monitoring land-atmosphere feedbacks related to land use change or global warming



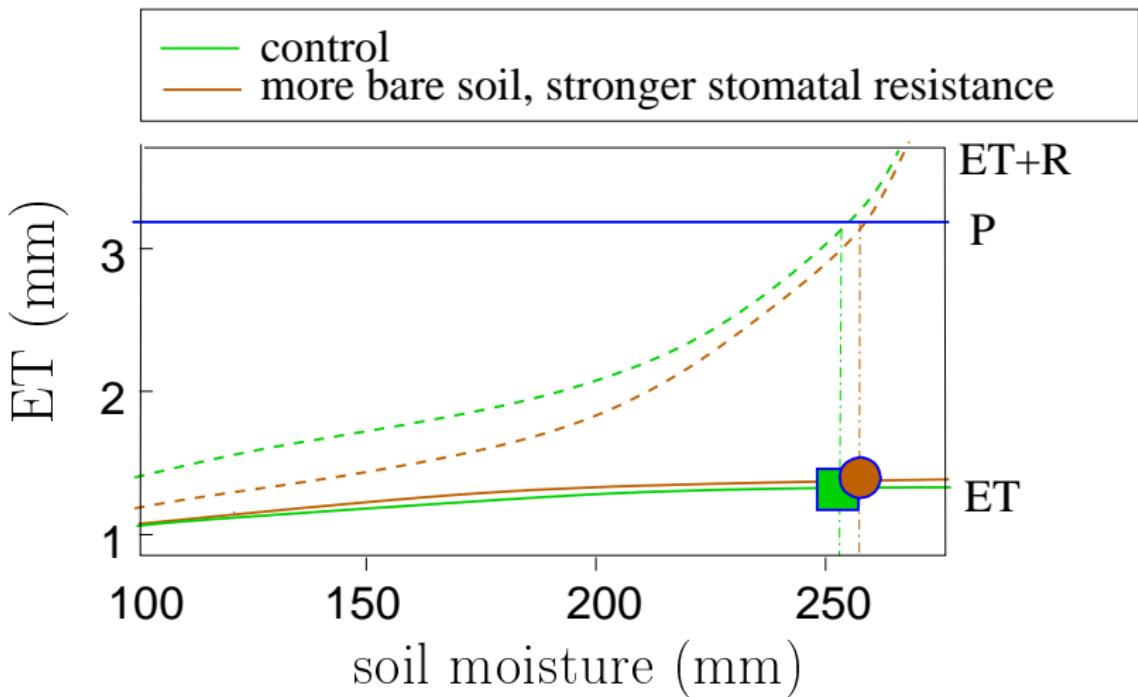
# Monitoring land-atmosphere feedbacks related to land use change or global warming



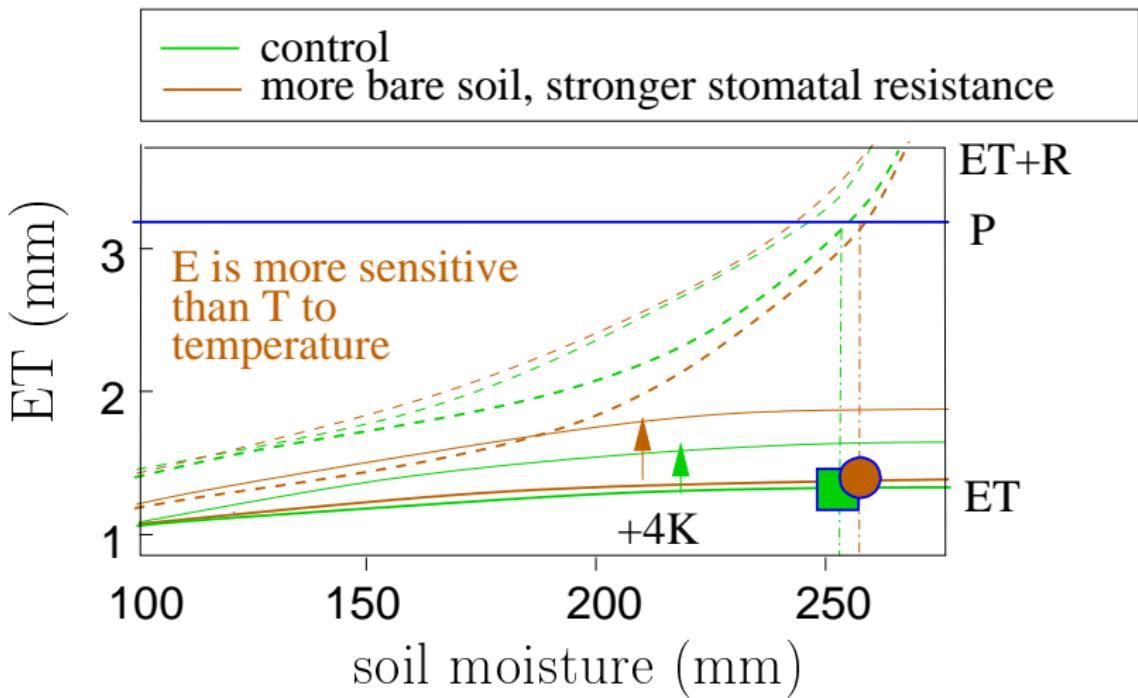
# Monitoring land-atmosphere feedbacks related to land use change or global warming



## Impact on response to temperature



## Impact on response to temperature



## Impact on response to temperature

