

Using water vapor isotopic measurements to evaluate moist and cloud processes in general circulation models

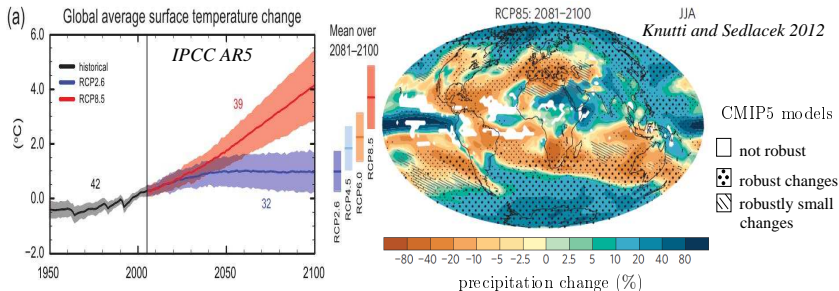
Camille Risi

LMD/IPSL/CNRS

with the contribution of: Sandrine Bony, Valérie Masson-Delmotte, You He, Jean-Lionel Lacour, Boutheina Oueslati, Obbe Tuinenburg, Françoise Vimeux, John Worden

NASA-JPL, October 16, 2014

Moist and cloud processes in climate models

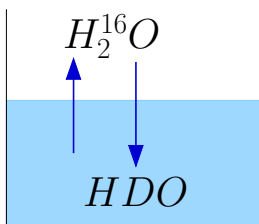


- ▶ Sources of spread? Temperature: cloud feedbacks (*Bony et al 2004, Dufresne et al 2008, Vial et al 2014*)
- ▶ For precip: convective parameterizations, cloud feedbacks? (e.g. *Kang et al 2008, Frierson and Hwang 2012*)

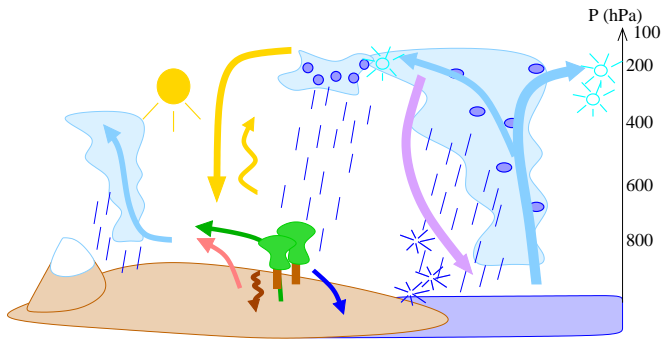
⇒ Need to better evaluate these processes in models

Water isotopes

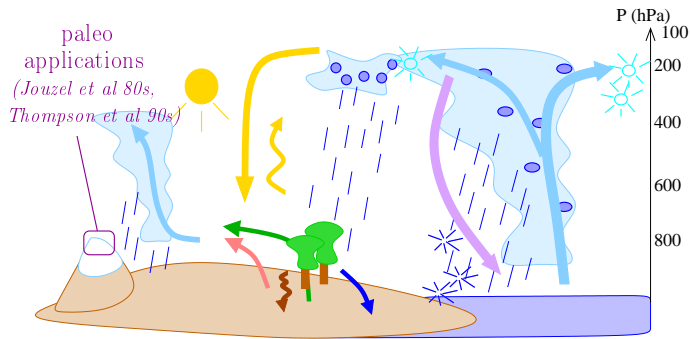
- ▶ $H_2^{16}O$, HDO , $H_2^{18}O$...
- ▶ $\delta D = (HDO/H_2O/R_{sea\ water} - 1) \cdot 1000$ in ‰
- ▶ fractionation during phase changes
⇒ tracers of the water cycle



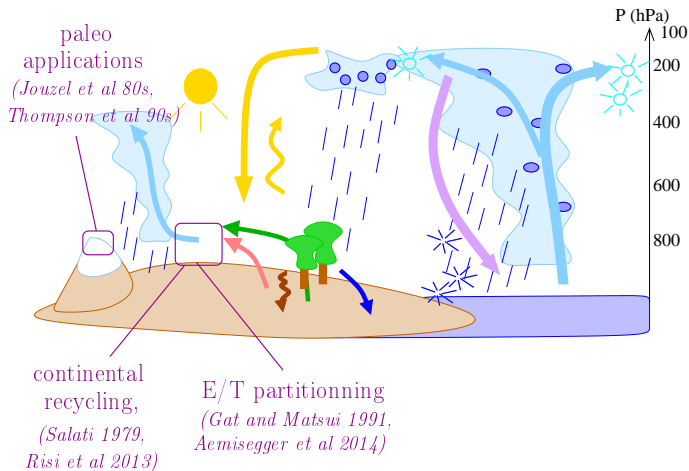
What can we use isotopes for?



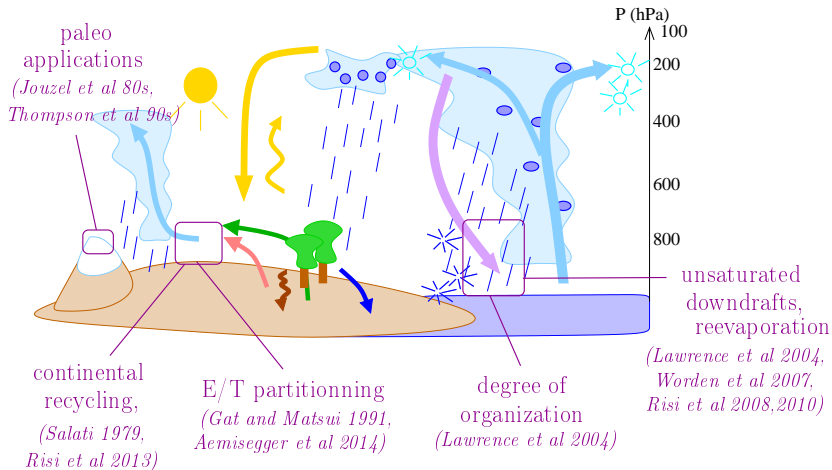
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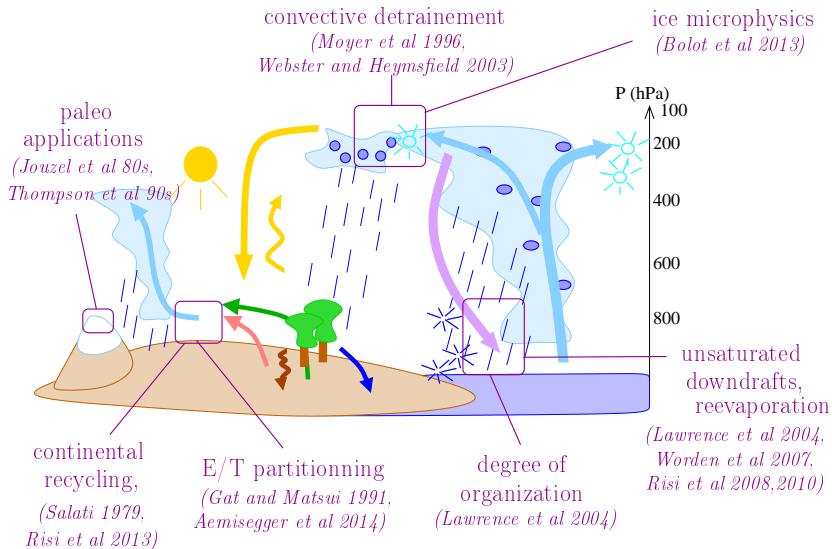
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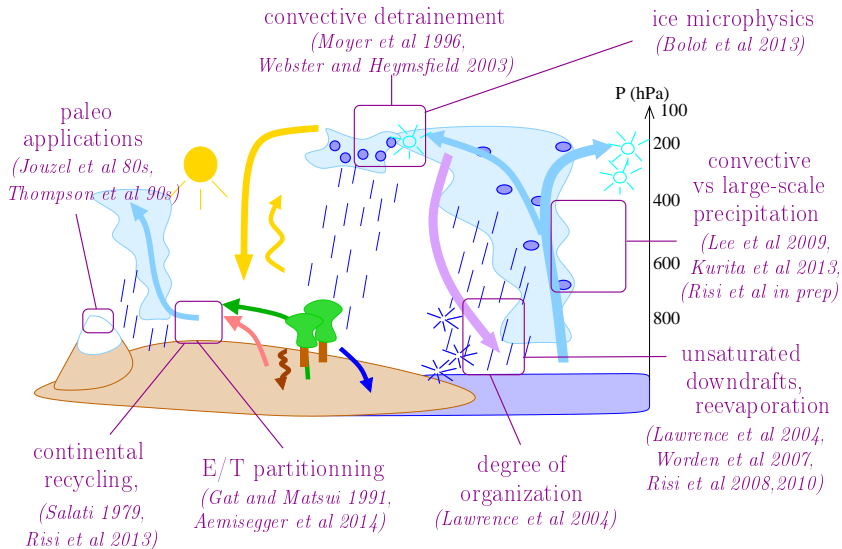
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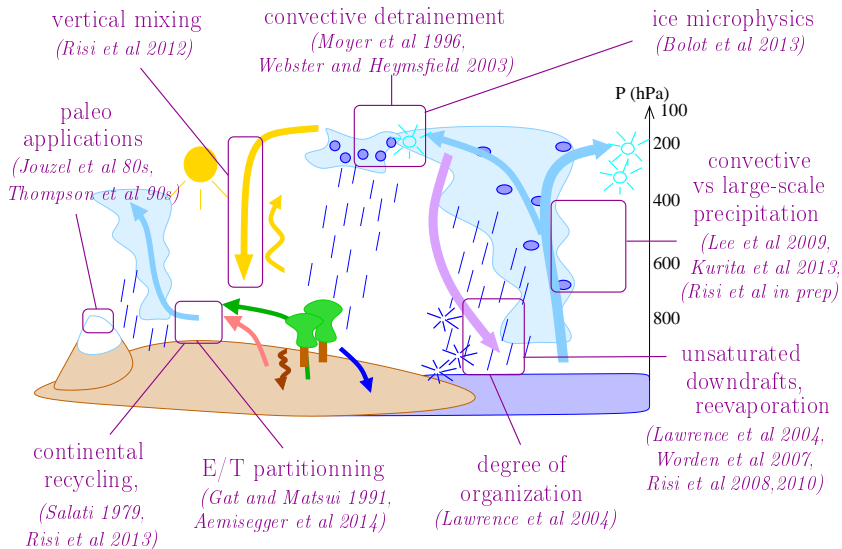
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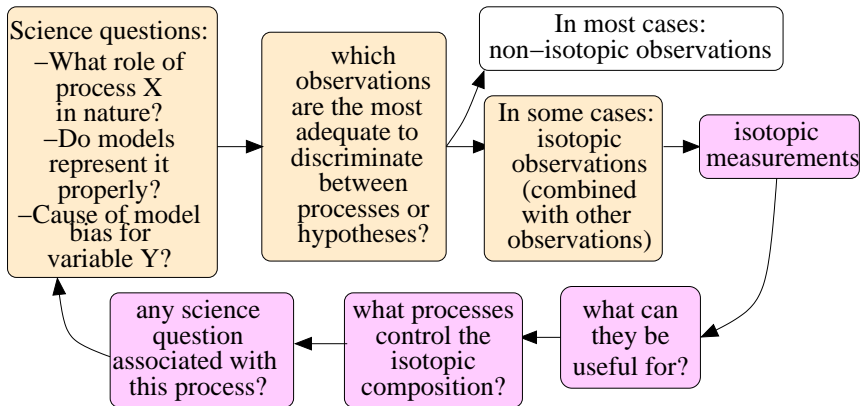
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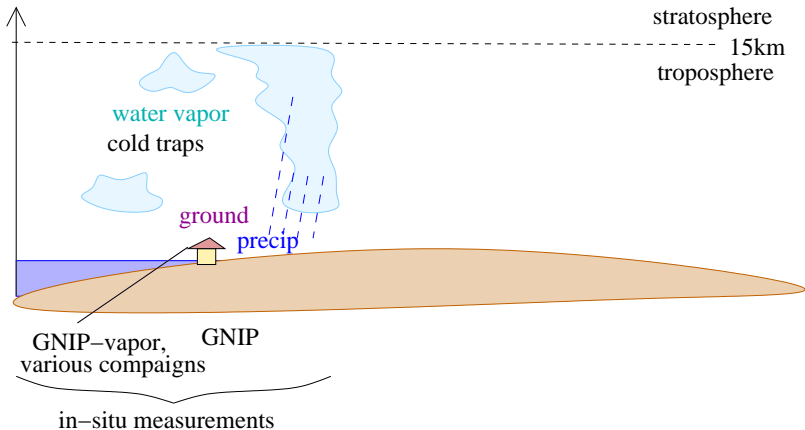
Why do we want isotopes to be useful?



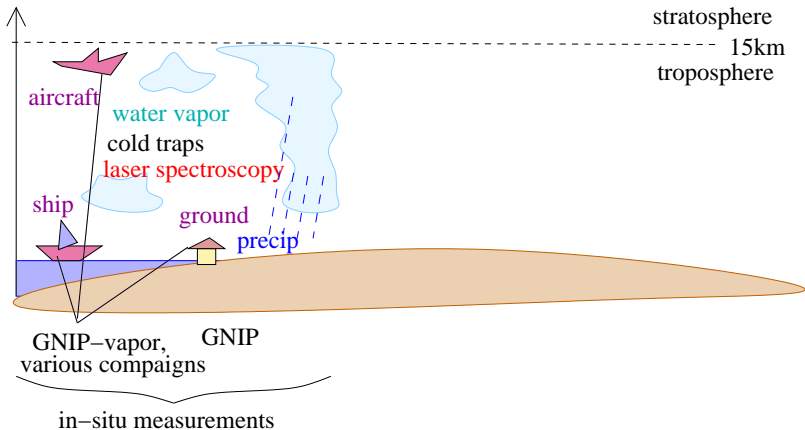
Seminar outline and summary

1. recent developments in water vapor isotopic measurements and models
2. Upper troposphere: δD sensitive to moistening by convective detrainment
3. Mid-troposphere: what does the δD -precipitation link says about the model physics?
 - 3.1 δD reflects shallow vs deep convective mixing and associated large-scale circulation
 - 3.2 δD reflects convective vs large-scale precipitation and associated heating profiles
4. During MJO events: mid-tropospheric δD evolution reflects the relative timing of different cloud types and associated moistening and dehydrating processes

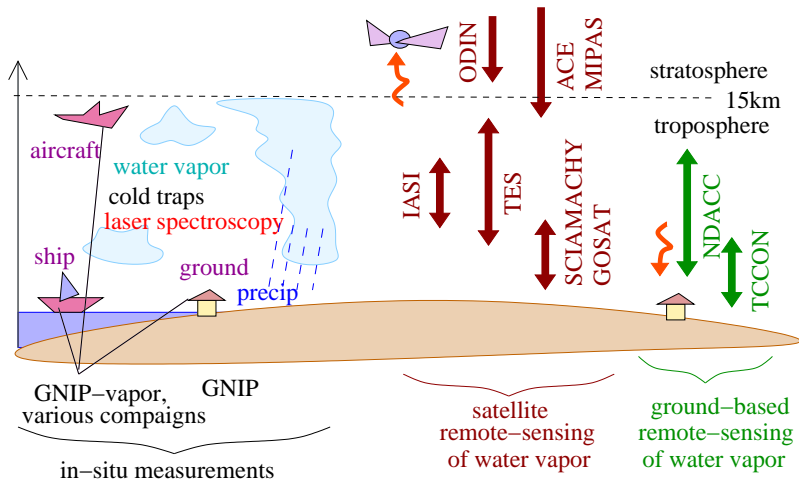
Measurements



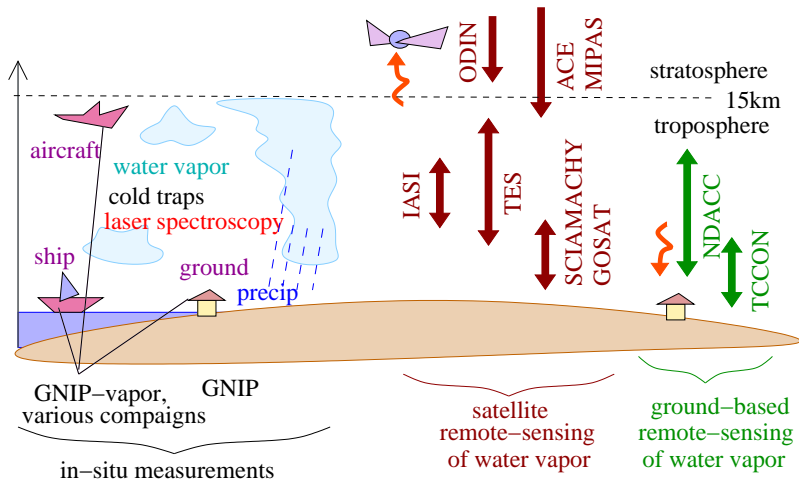
Measurements



Measurements



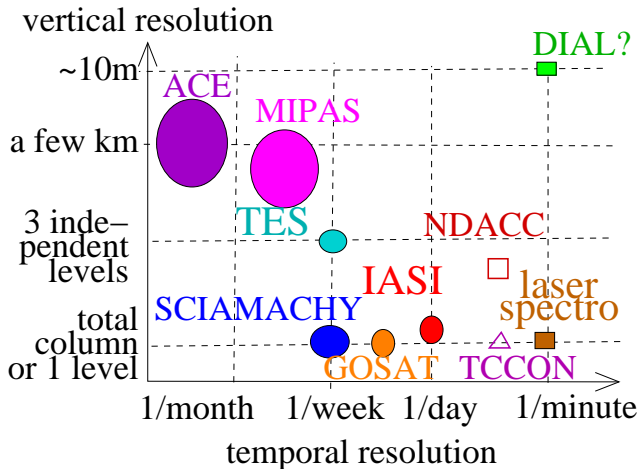
Measurements



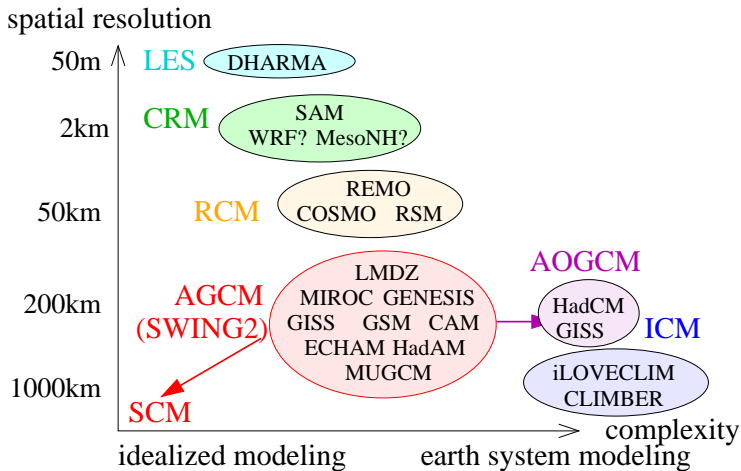
- ▶ TES: weekly vertical profiles of δD
- ▶ IASI: 2 daily global coverage of δD

Diversity of measurements

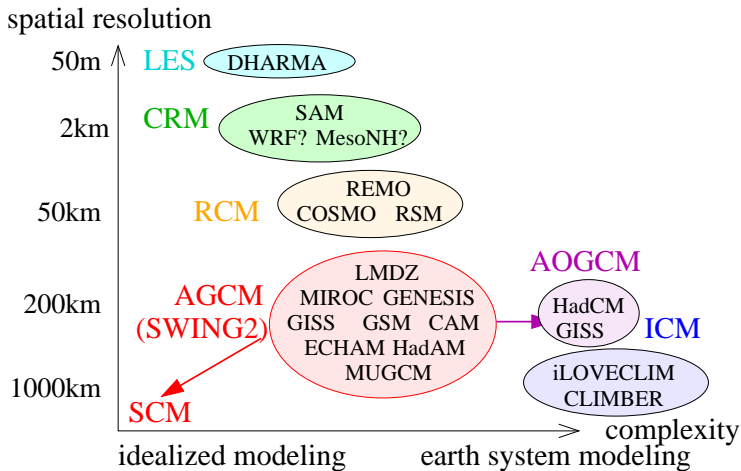
- ▶ different altitudes, temporal resolution, vertical resolution, precision, spatial resolution and coverage...



Numerical models



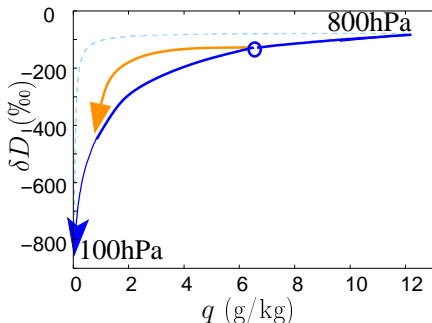
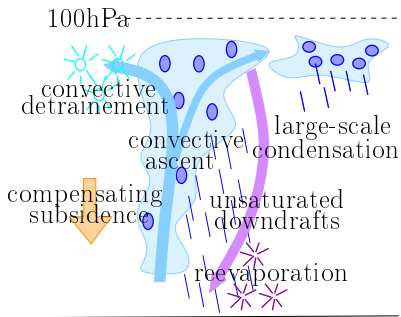
Numerical models



► Here: LMDZ model, SWING2

Theoretical framework: q - δD

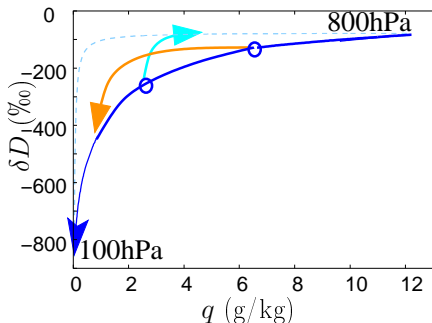
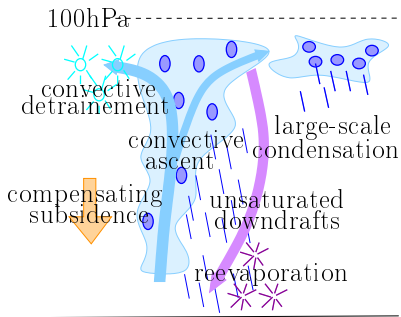
- ▶ Moistening and dehydrating processes (*Worden et al 2007*)



- ▶ large-scale condensation
- ▶ subsidence

Theoretical framework: q - δD

- Moistening and dehydrating processes (*Worden et al 2007*)

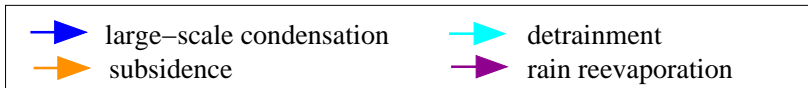
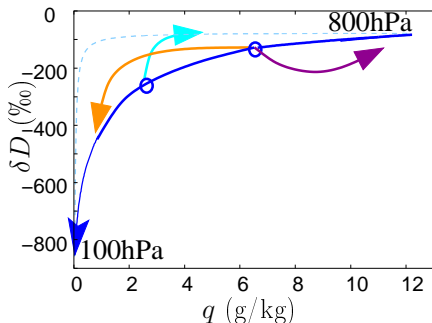
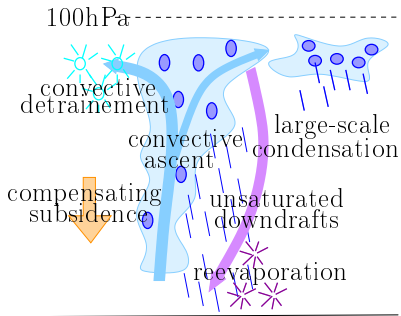


► large-scale condensation
► subsidence

► detrainment

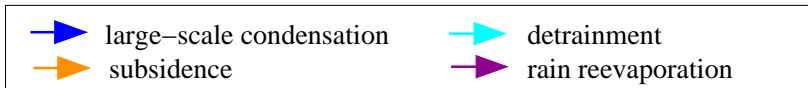
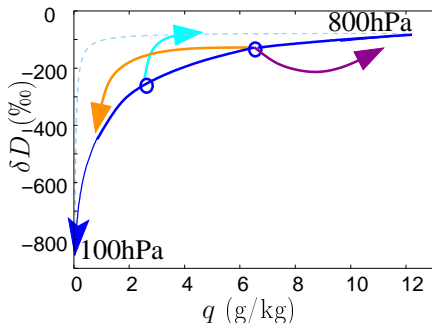
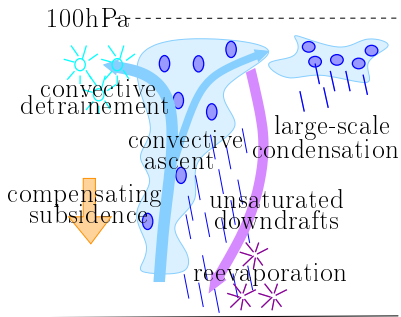
Theoretical framework: q - δD

- Moistening and dehydrating processes (*Worden et al 2007*)



Theoretical framework: q - δD

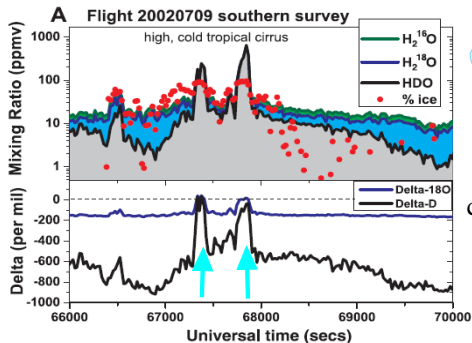
- ▶ Moistening and dehydrating processes (*Worden et al 2007*)



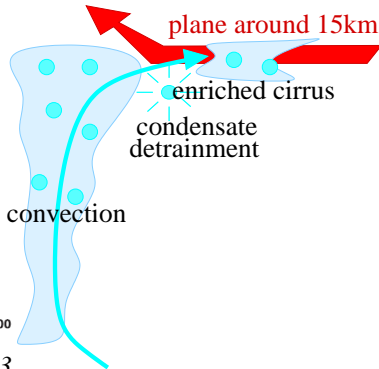
- ▶ limitation: need to bridge gap between this simple framework and numerical modeling

Isotopes in the upper troposphere

- ▶ papers from *Moyer, Kuang, Dessler, Sherwood, Sayres, Hanisco...*

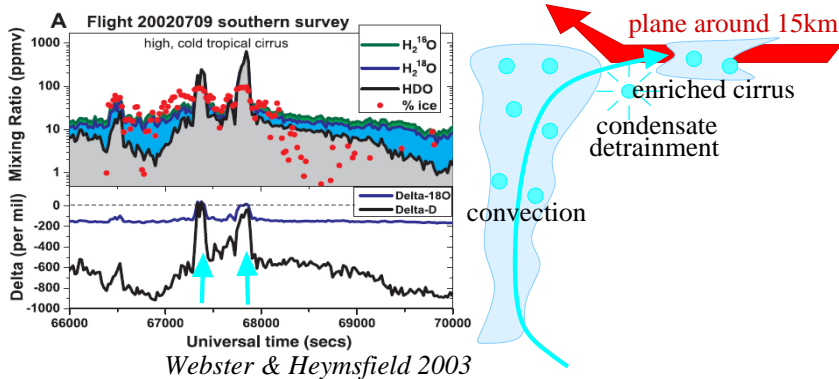


Webster & Heymsfield 2003



Isotopes in the upper troposphere

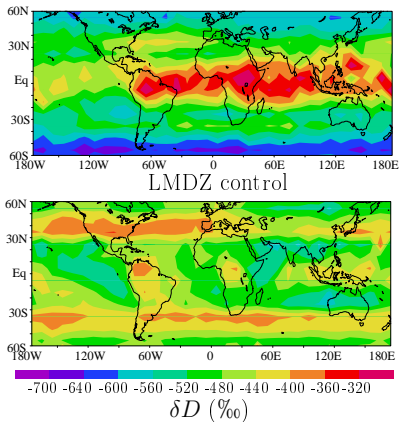
- ▶ papers from *Moyer, Kuang, Dessler, Sherwood, Sayres, Hanisco...*



- ▶ Limitation: isotopes consistent with some convective injection of water through the tropopause layer. But how to make quantitative estimations?

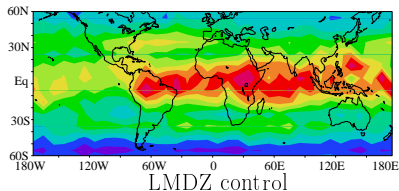
Convective detrainment in upper troposphere

MIPAS data at 200hPa, annual

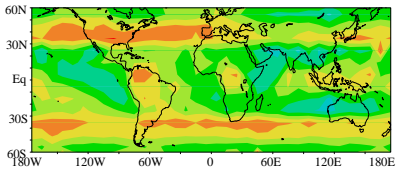


Convective detrainment in upper troposphere

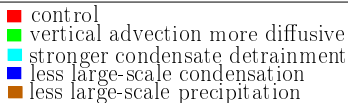
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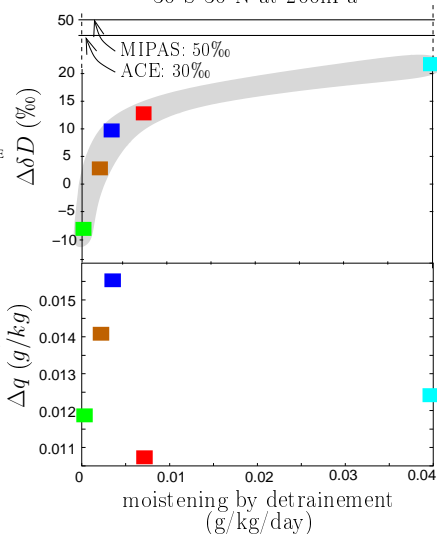
LMDZ control



ΔD (‰)

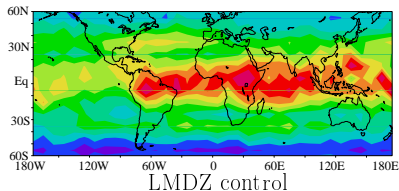


Difference 15°S-15°N minus
30°S-30°N at 200hPa

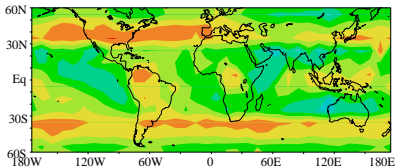


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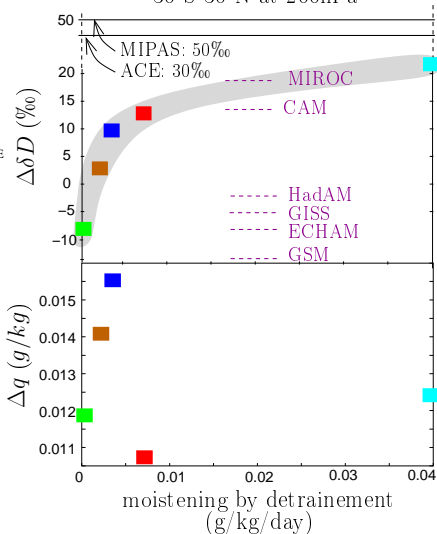
LMDZ control



ΔD (‰)

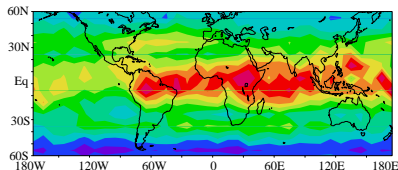
- control
- vertical advection more diffusive
- stronger condensate detrainment
- less large-scale condensation
- less large-scale precipitation

Difference 15°S-15°N minus 30°S-30°N at 200hPa

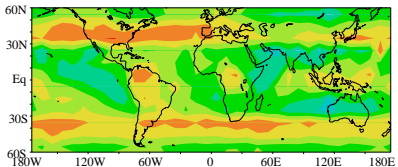


Convective detrainment in upper troposphere

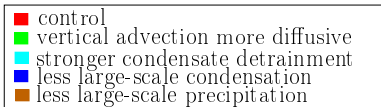
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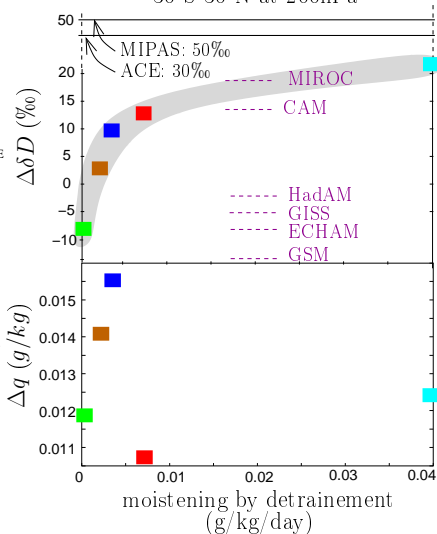
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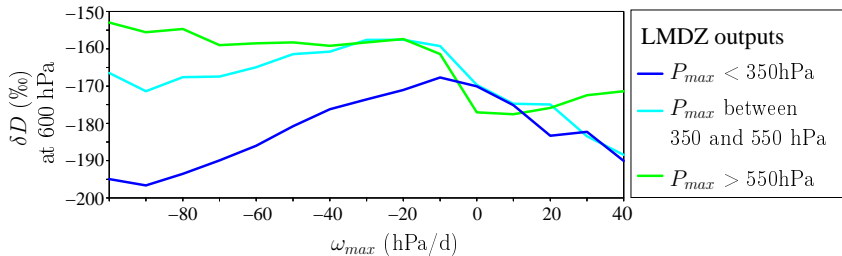


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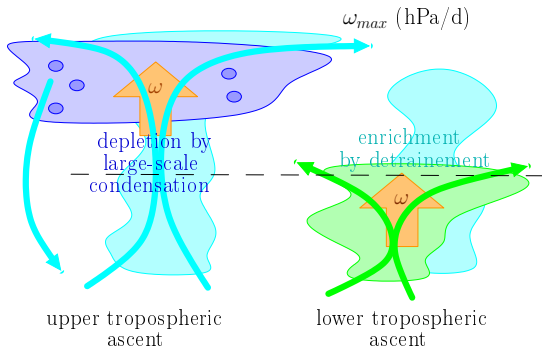
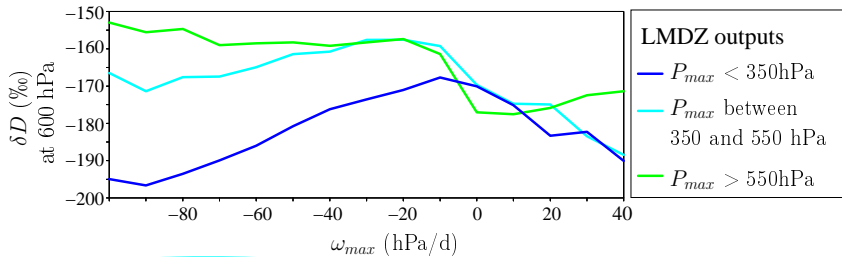


► what is missing? Correlation detrainment ↔ updraft speed?

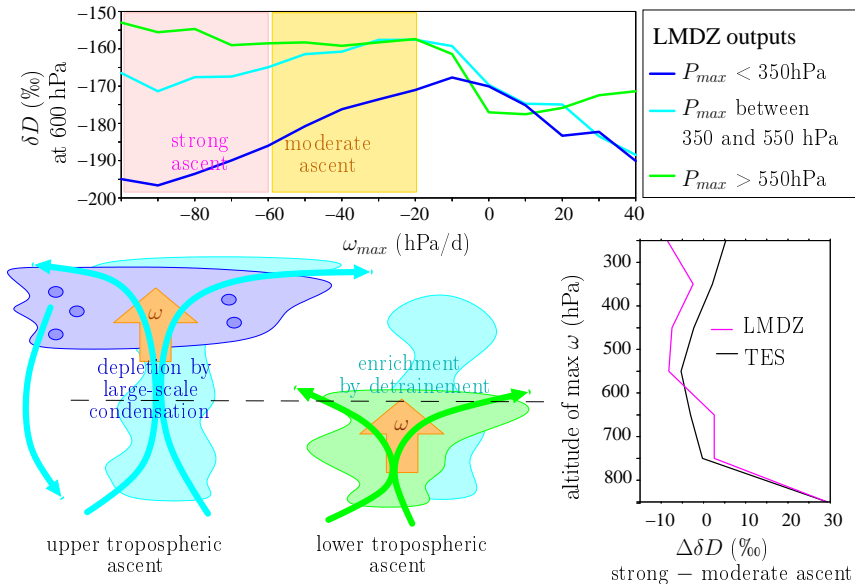
Sensitivity to large-scale velocity profile



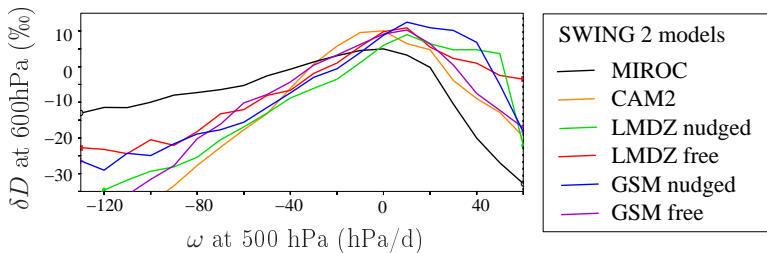
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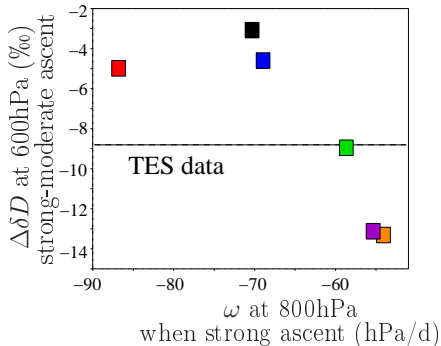
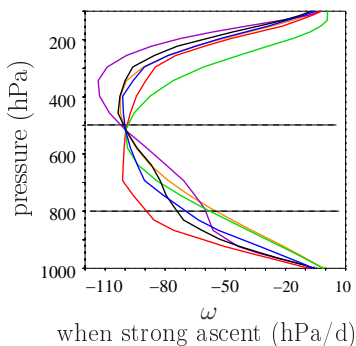
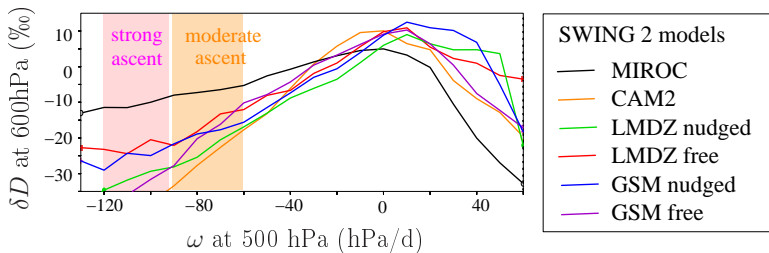
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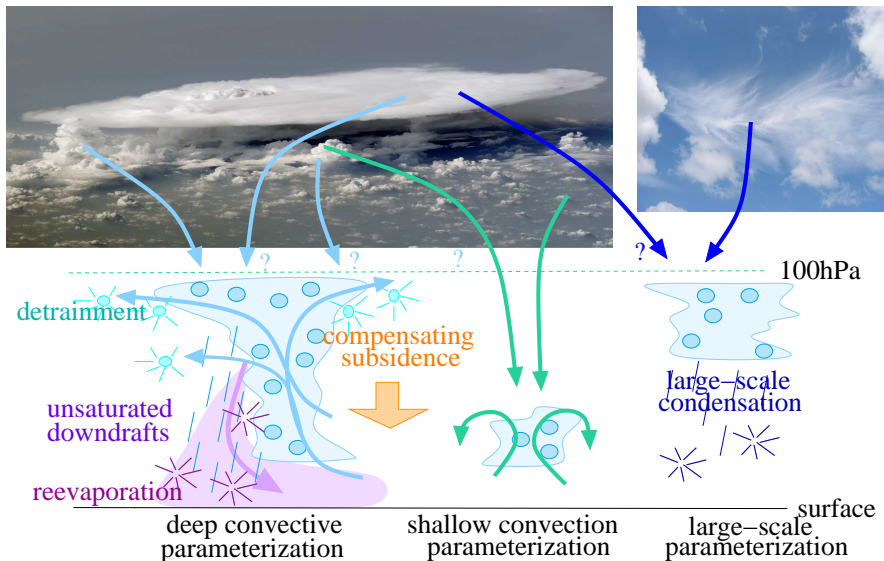
And across models?



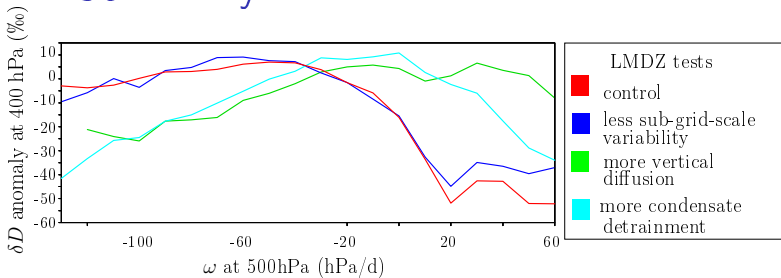
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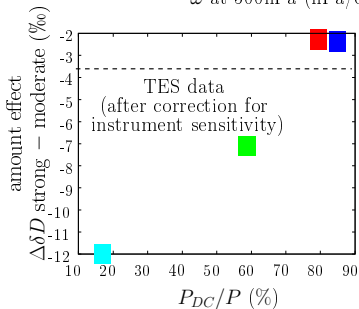
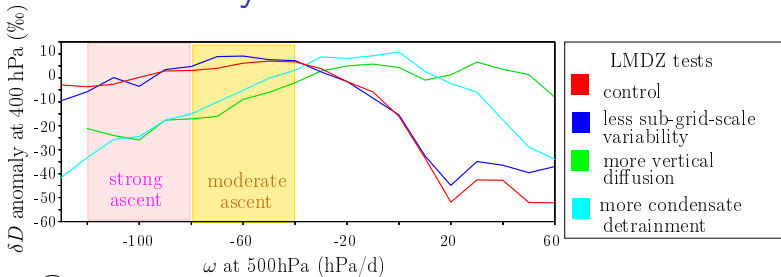
3) Convection vs large-scale schemes



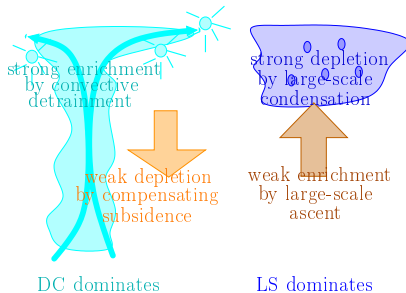
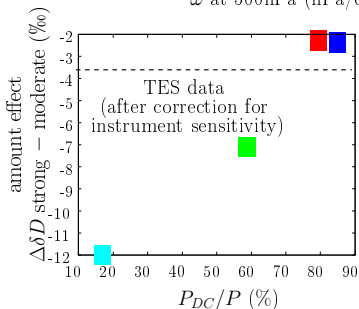
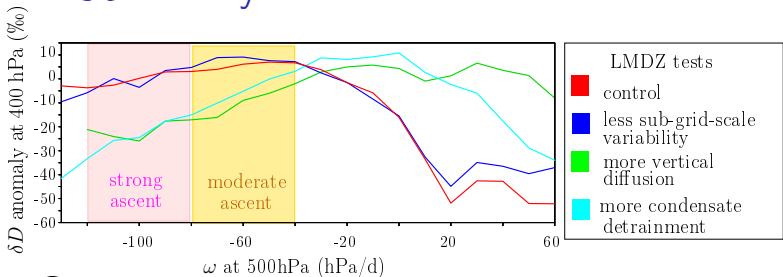
Sensitivity tests in LMDZ



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Sensitivity tests in LMDZ



Implications

- ▶ precipitating events deplete the tropospheric vapor all the more as it is associated with large-scale precipitation
- ▶ conv vs large-scale precipitation = arbitrary choice specific to each model, but with consequences on:
 - ▶ latent heating profiles \Rightarrow large-scale circulation
 - ▶ cloudiness
 - ▶ water vapor, chemical and aerosol transport
 - ▶ intra-seasonal variability (*Kim et al 2012*)

\Rightarrow use water isotopes quantitatively to evaluate conv vs large-scale precip partitioning and underlying heating profiles?

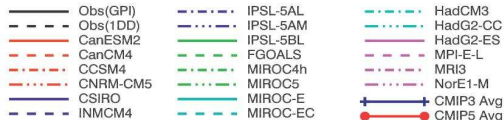
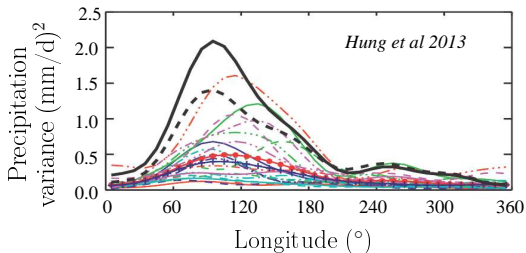
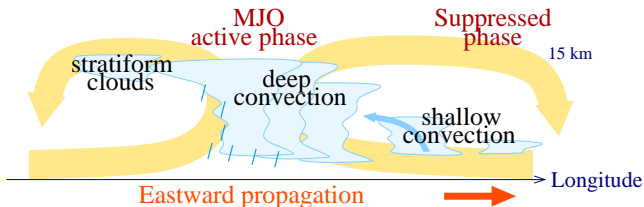
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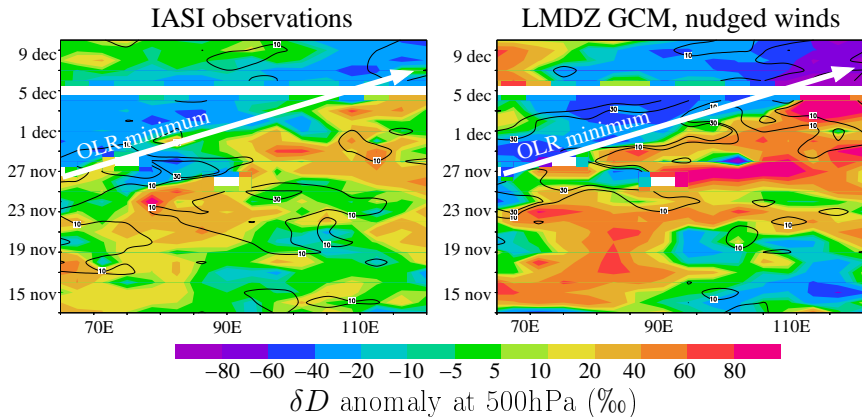
\Rightarrow use water isotopes quantitatively to evaluate conv vs large-scale precip partitioning and underlying heating profiles?

- ▶ work in progress with SCM
- ▶ Obbe Tuinenbourg's work: use water isotopes to evaluate sequence of cloud processes during the Madden-Julian Oscillation (MJO)?

4) MJO so difficult to simulate?



Cindy Dynamo campaign case

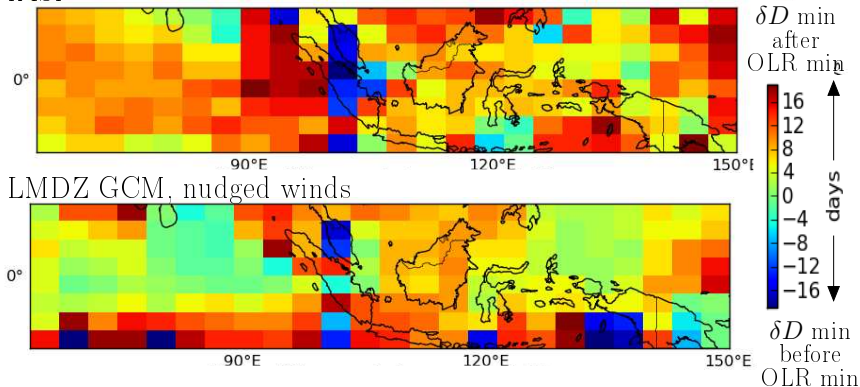


- ▶ Observed q max 0-1 days before OLR min
- ▶ Observed δD min 3 days after OLR min
- ▶ LMDZ captures this lag for this case

Statistical analysis for 2006-2007

Phasing of δD min at 500hPa vs OLR min

IASI

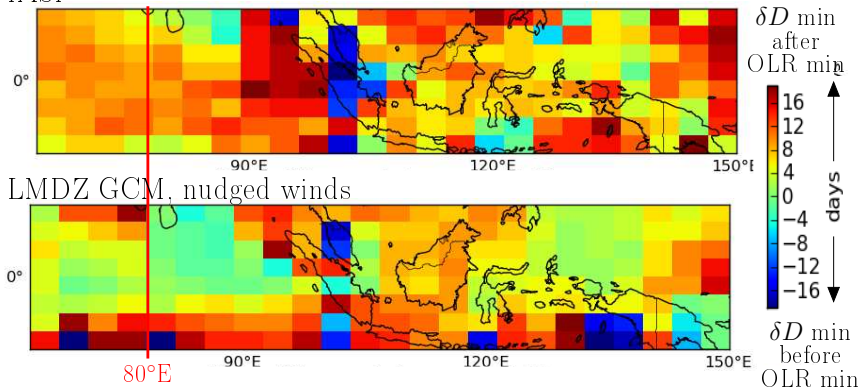


- ▶ Observed δD min lags OLR min in Indian Ocean
- ▶ More complicated over Maritime Continent
- ▶ LMDZ δD to in phase with OLR

Statistical analysis for 2006-2007

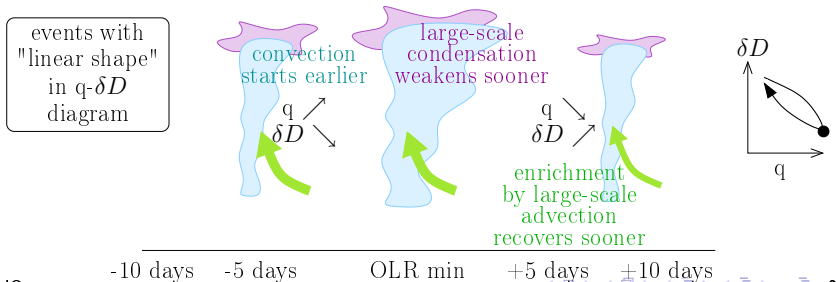
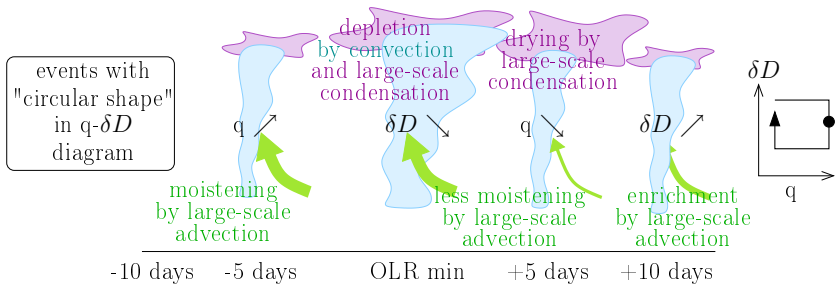
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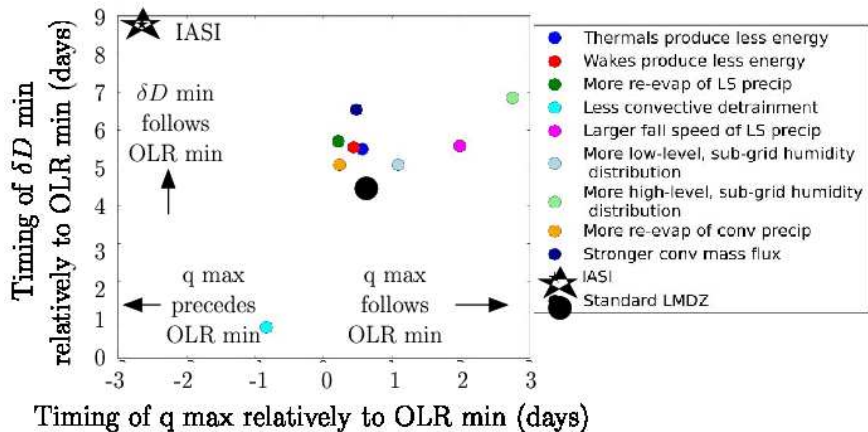
What determines $q - \delta D$ shape in LMDZ?



Preliminary summary on $q - \delta D$ cycles during the MJO

- ▶ Observed “circular shape” over Indian Ocean consistent with cloud evolution shallow \rightarrow deep \rightarrow stratiform
- ▶ What happens over the Maritime Continent?
- ▶ Still lot of work to fully understand both data and model behavior
- ▶ LMDZ too in phase:
 - ▶ convection triggers too soon?
 - ▶ Large-scale condensation not maintained long enough?
 - ▶ Large-scale advective enrichment recovers too soon?
- ▶ $q - \delta D$ useful for model evaluation? \Rightarrow work in progress: analyze sensitivity tests

Sensitivity tests with LMDZ



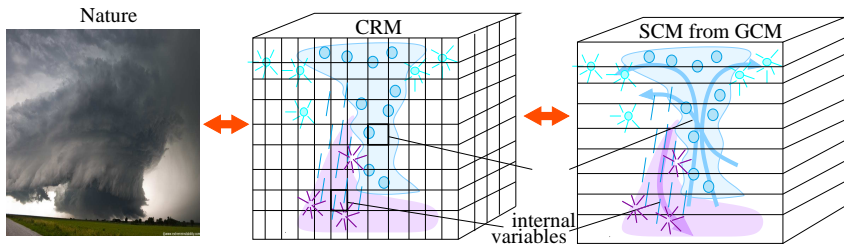
- ▶ $q - \delta D$ shape sensitive to convection/cloud parameters
- ▶ How to get closer to observations?

Summary on $q - \delta D$ cycles during the MJO

- ▶ $q - \delta D$ cycles during MJO: informs about the relative timing of shallow convection, deep convection, large-scale condensation and large-scale advection
- ▶ Potentially useful for model evaluation
- ▶ Still lot of work to fully understand both data and model behavior
- ▶ Exploit better the Cindy Dynamo campaign data?

General perspectives

- ▶ Bridge gap between simple theoretical $q - \delta D$ framework and more complex modeling
- ▶ Better exploit model hierarchy: GCM vs SCM, with large-scale circulation as forcing or as a response (WTG).
- ▶ intercompare GCMs: add daily in SWING2?
- ▶ CRM/LES to study processes and to compare more easily to observations and to SCMs (conditional sampling)



- ▶ longer term: combine q , δD + chemical tracers: CO , O_3 , ^{10}Be
⇒ better characterize fluxes