

Test d'advection d'un pic Gaussien

— Distribution initiale

— Solution exacte (translation)

U constant, nombre de Courant $U \, dt/dx = 0,2$

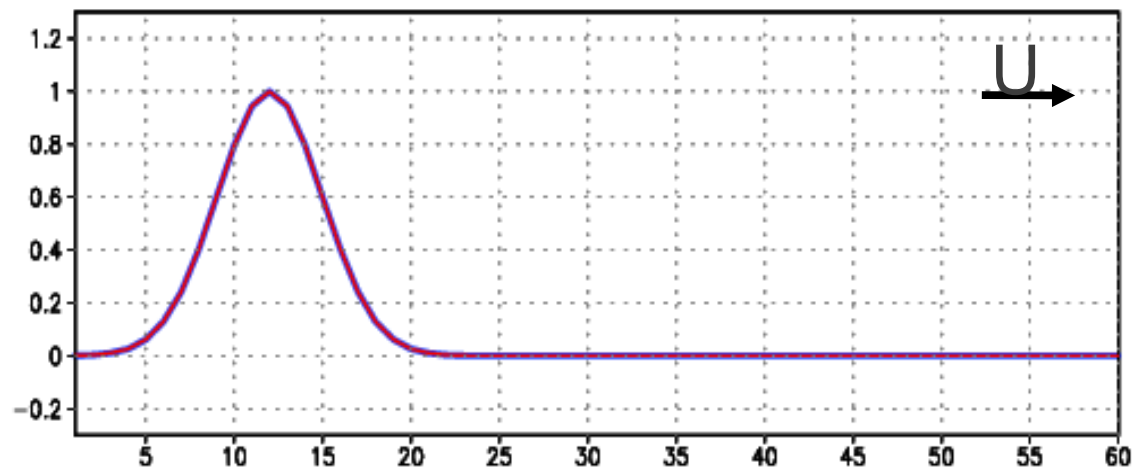
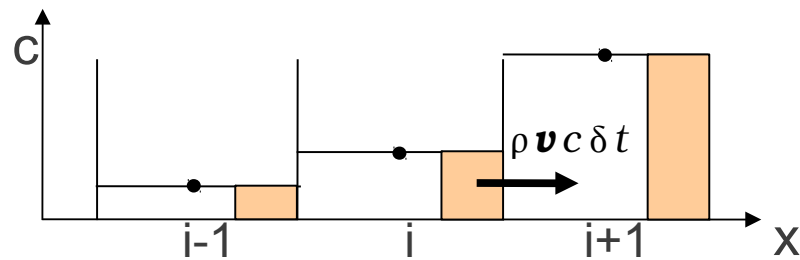


Schéma amont (Godunov, 1952)



Test d'advection d'un pic Gaussien

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— Calcul avec schéma d'advection

U constant, nombre de Courant $U dt/dx = 0,2$

Advection avec le schéma amont ou semi-lagrangien d'ordre 1

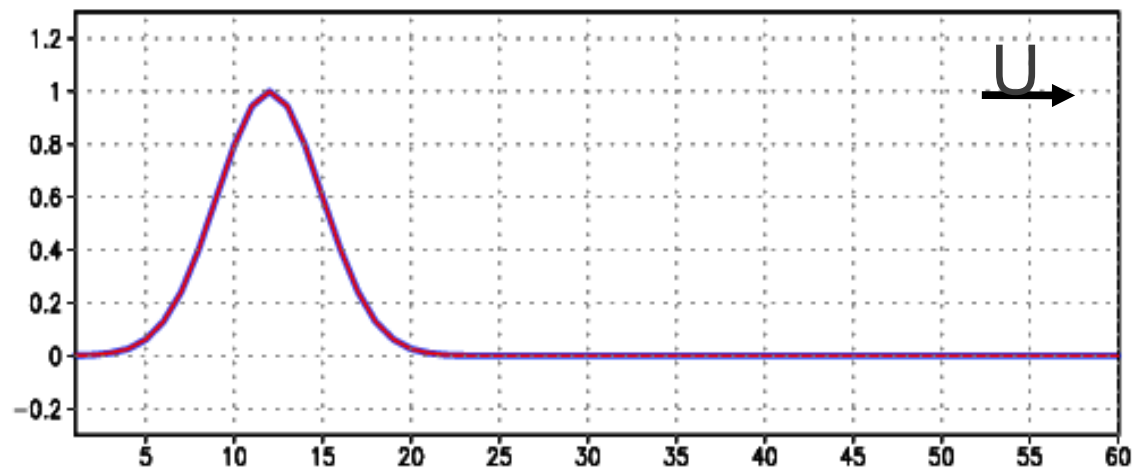
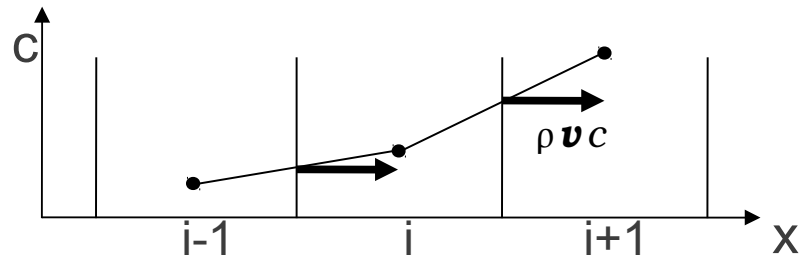


Schéma diffusif et peu précis

Différences finies centrées (second ordre)



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Advection avec le schéma en différences finies centré d'ordre 2

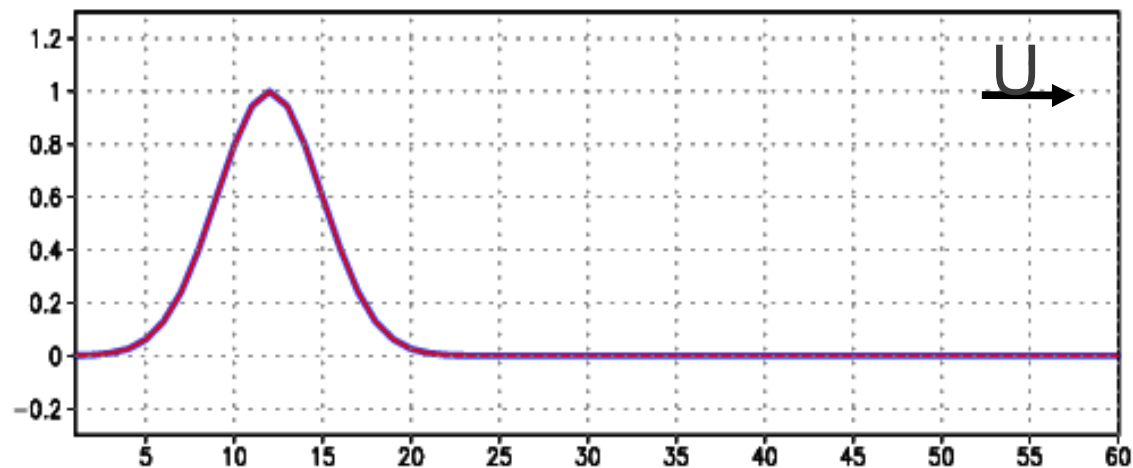
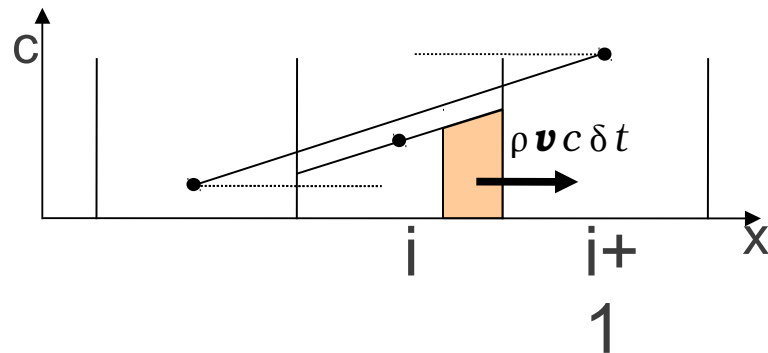


Schéma précis mais création de valeur négative et instabilité

Schéma I de Van Leer (1977)



Test d'advection d'un pic Gaussien

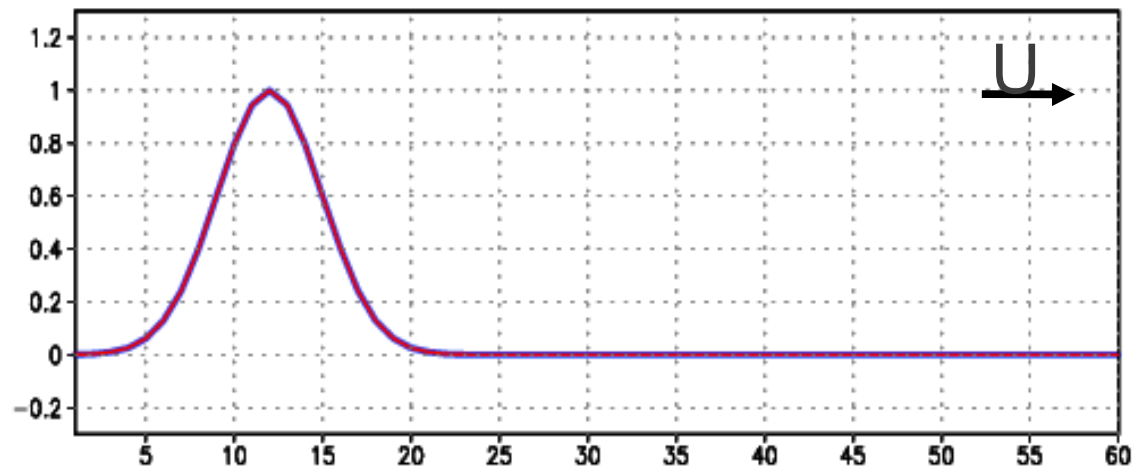
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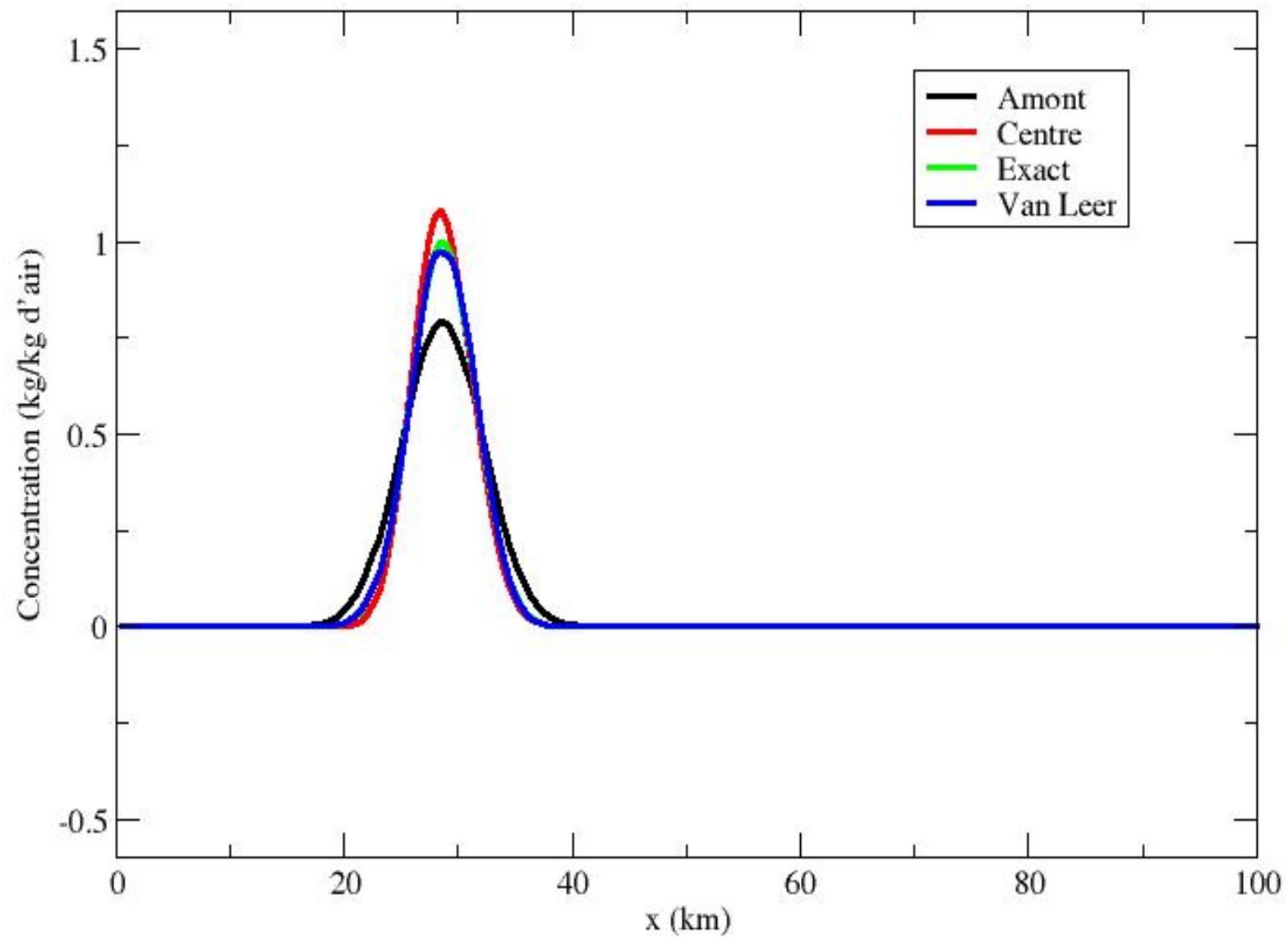
— Solution exacte (translation)

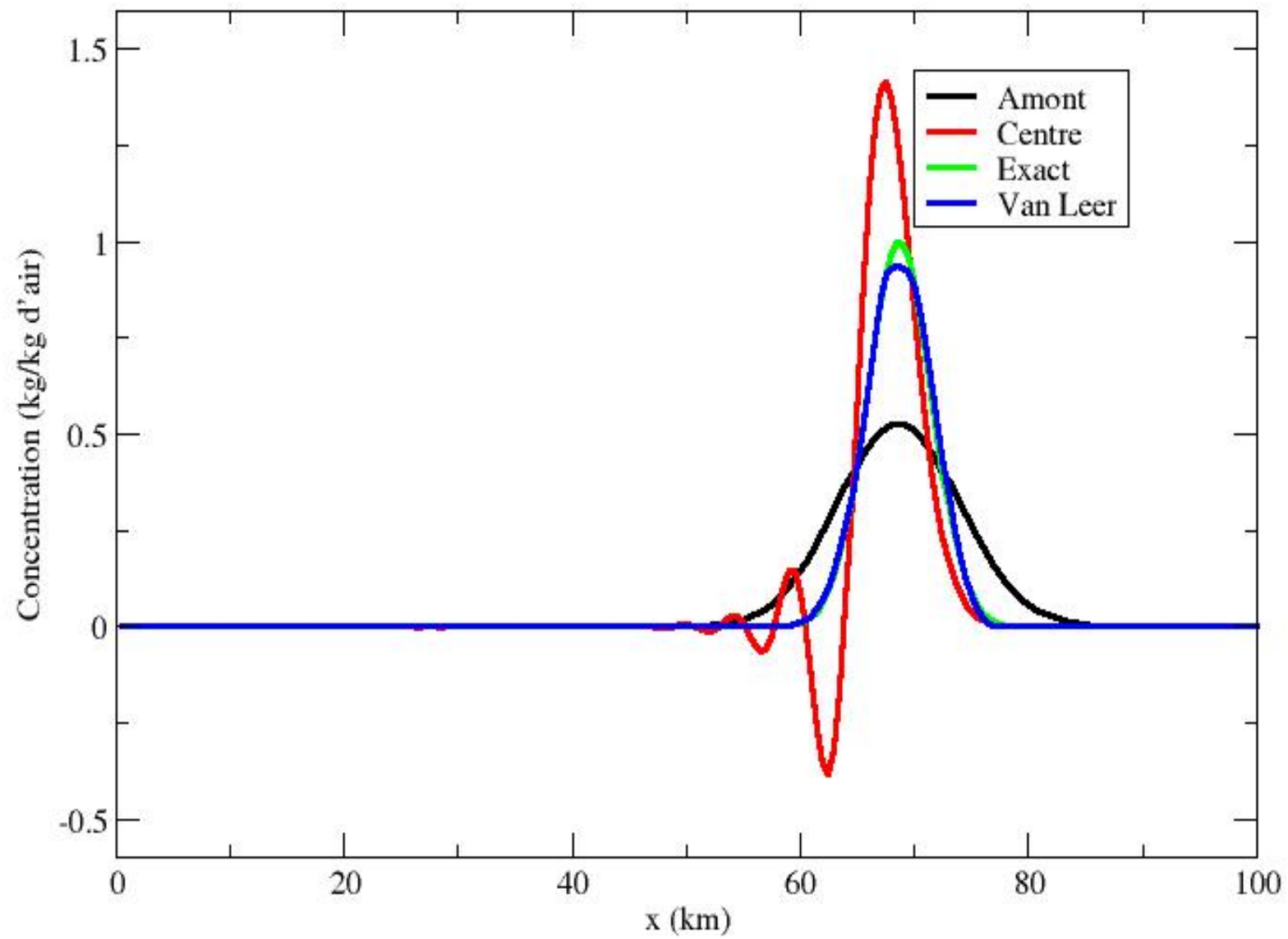
— Calcul avec schéma d'advection

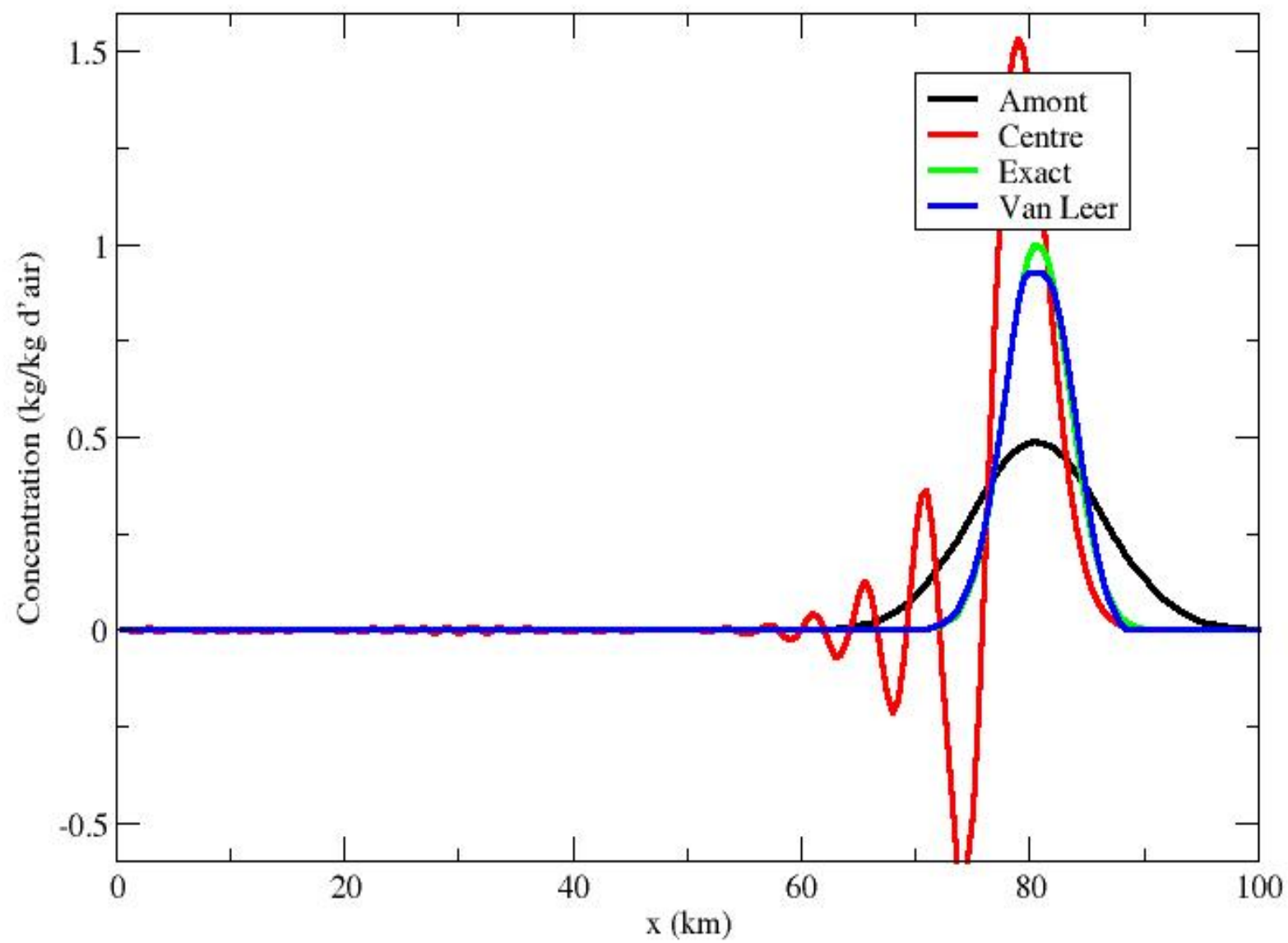
U constant, nombre de Courant $U \, dt/dx = 0,2$

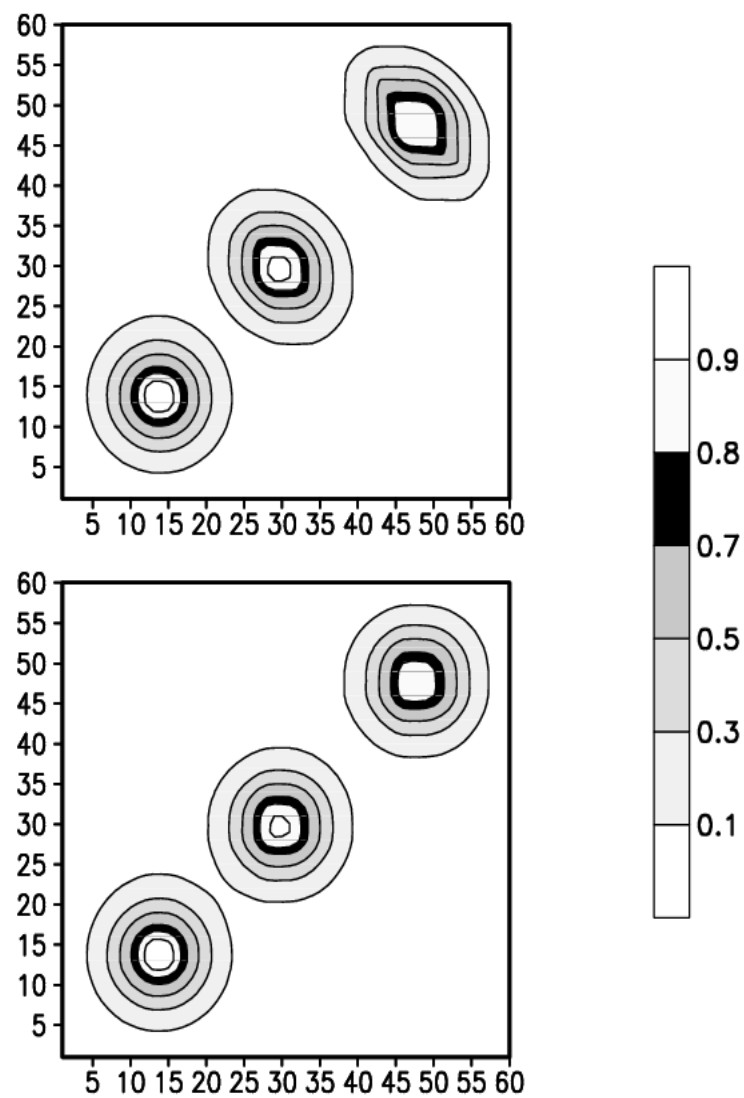
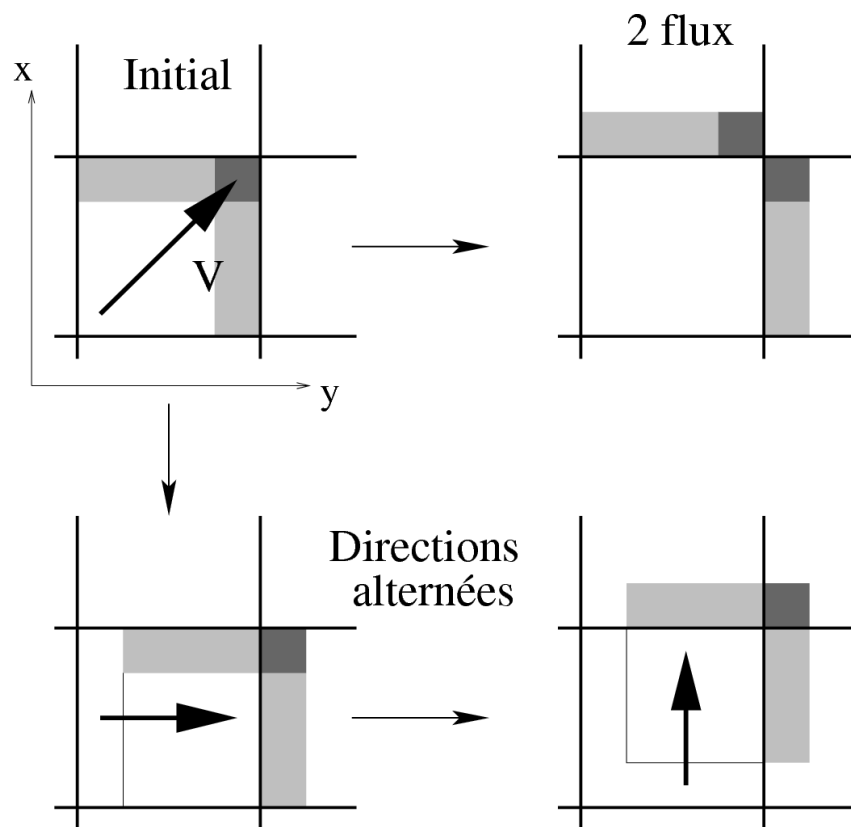
Advection avec le schéma volumes finis d'ordre 2 de Van Leer



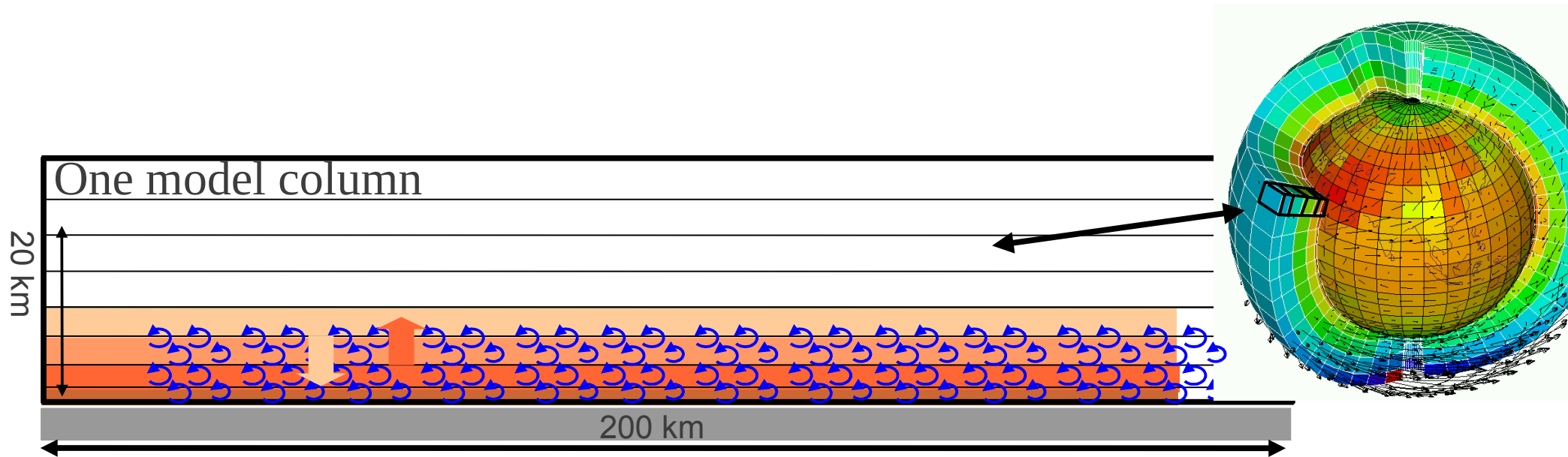








II.1 Thermal plumes and clouds



Classical approach :

→ « **Turbulent mixing** » or **diffusion**

Mixing by small scale random motion

Analogous to molecular diffusion

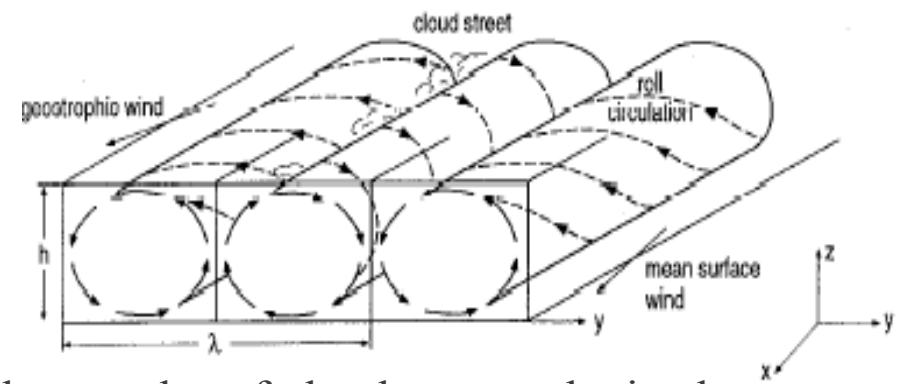
$$Dq/Dt = Sq \quad \text{avec} \quad Sq = \frac{\partial}{\partial z} \left(K_z \frac{\partial q}{\partial z} \right)$$

→ Computation of K_z : a field of research

$$K_z = f(dU/dz, d\theta/dz, e, \dots)$$

New equations, new parameters ...

Importance of organized structures Visualized by cloud streets

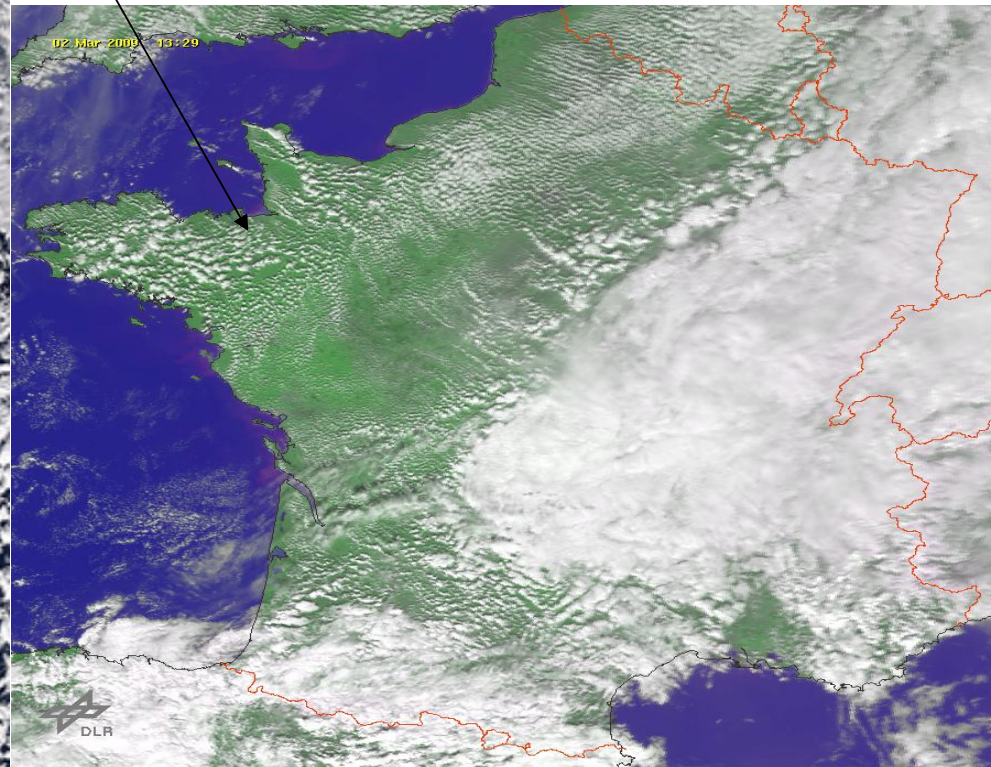


Classical examples of cloud streets obtained at the top of thermal rolls

- Polar air arriving on warm air masses
- Entry of maritime air on a warmer continent



(Atkinson & al., 1996)



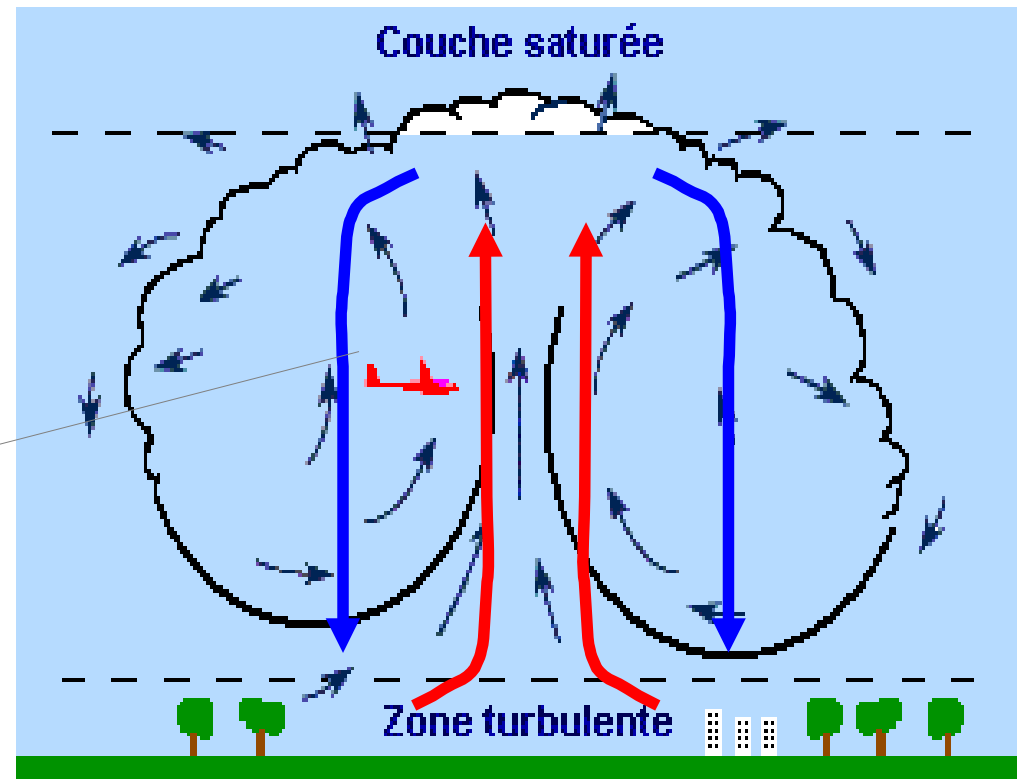
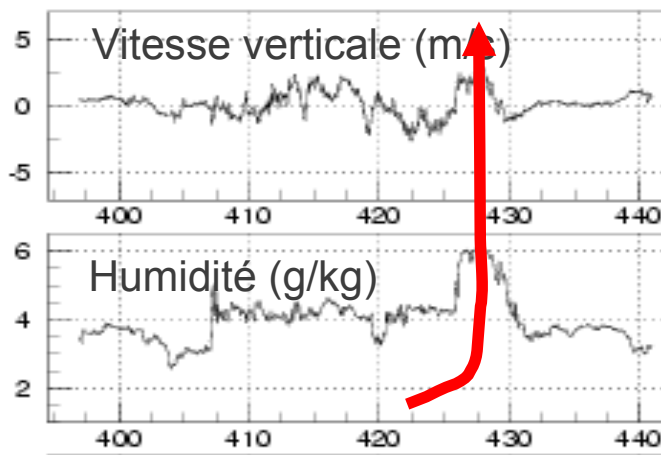
III. Illustration de la méthodologie de développement et d'évaluation sur des développements en cours

Turbulence isotrope de petite échelle → mélange turbulent

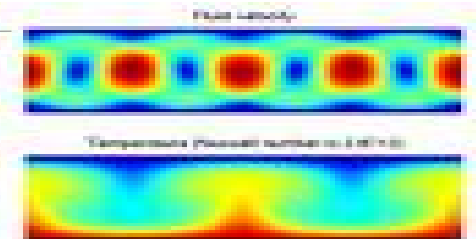
Turbulence atmosphérique :
“mésos-échelle”, organisée et anisotrope



Exemples de mesures avion en palier
(région parisienne, conditions estivales, cumulus)



- L'air chaud (léger) et humide monte de la surface sous l'effet des forces d'Archimède.
- Analogue à la convection de Rayleigh-Benard
- En montant cet air se refroidit (détente adiabatique) et ne peut plus contenir autant de vapeur d'eau.
- Si saturation : apparition de cumulus en haut du panache .



II.1 Thermal plumes and clouds

Genesis of the thermal plume approach

Mass flux schemes

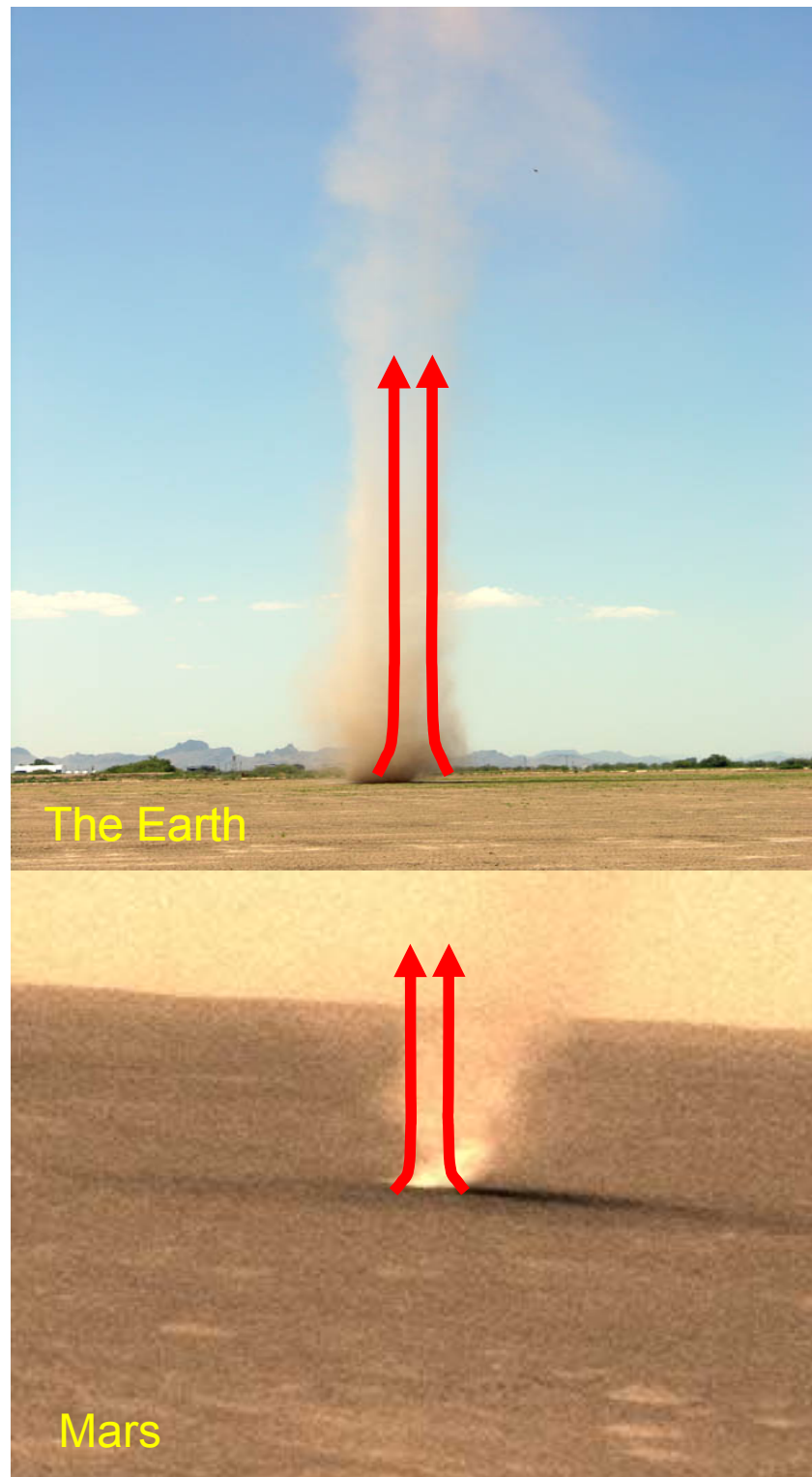
- mass flux schemes already used but essentially for clouds and deep convection
- dry convective boundary layer were given a weaker priority

Origin of the LMD thermal plume model (2002) :

- motivated by the Martian climate : Mars is a global desert with very strong and frequent dry convection
- Inspired by air plane observations during the Trac campaign (Paris area)

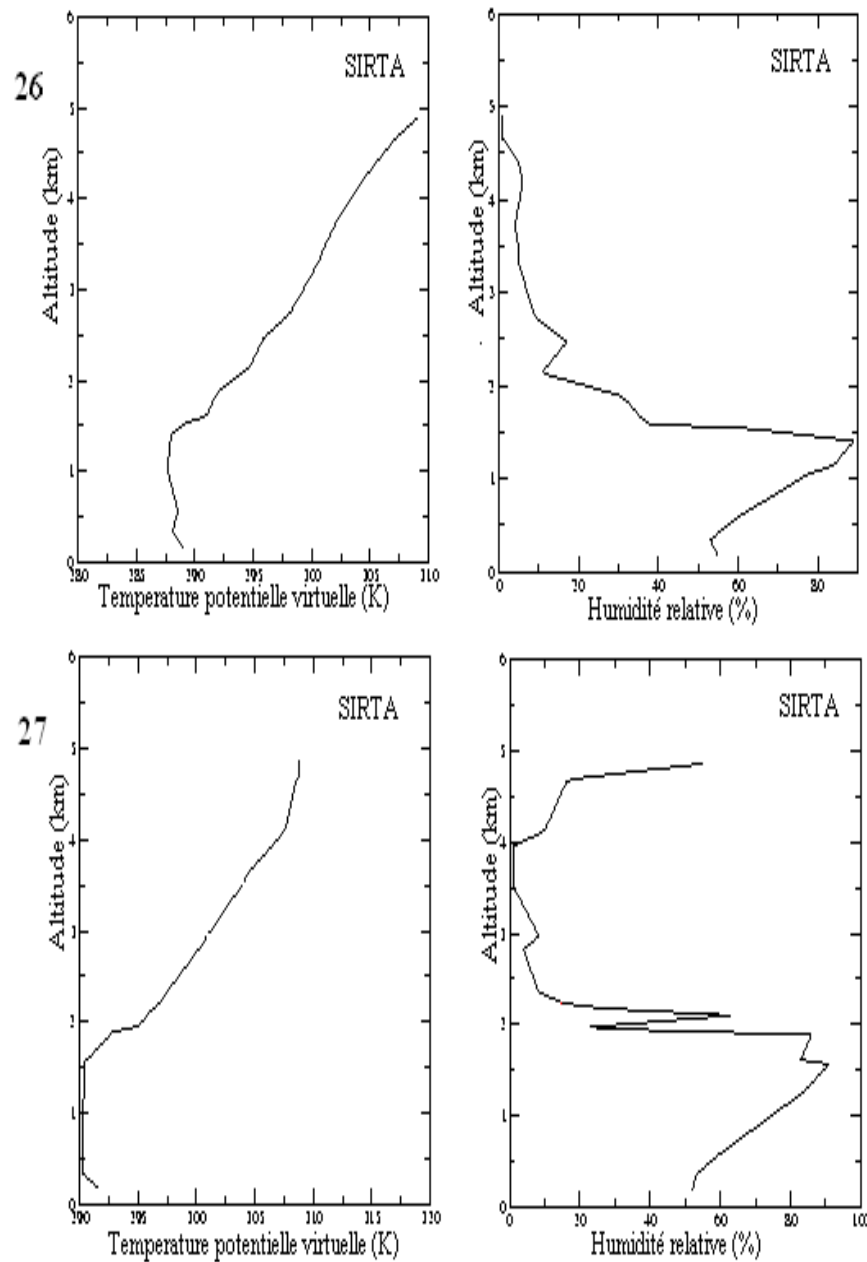
Other origins :

- First paper proposing the combination of a diffusive approach and mass flux scheme for the convective boundary layer (Chatfield, 1985)
- Independent parameterization issued from the GCSS and eurocs community (Siebesma and collaborators, 2004)

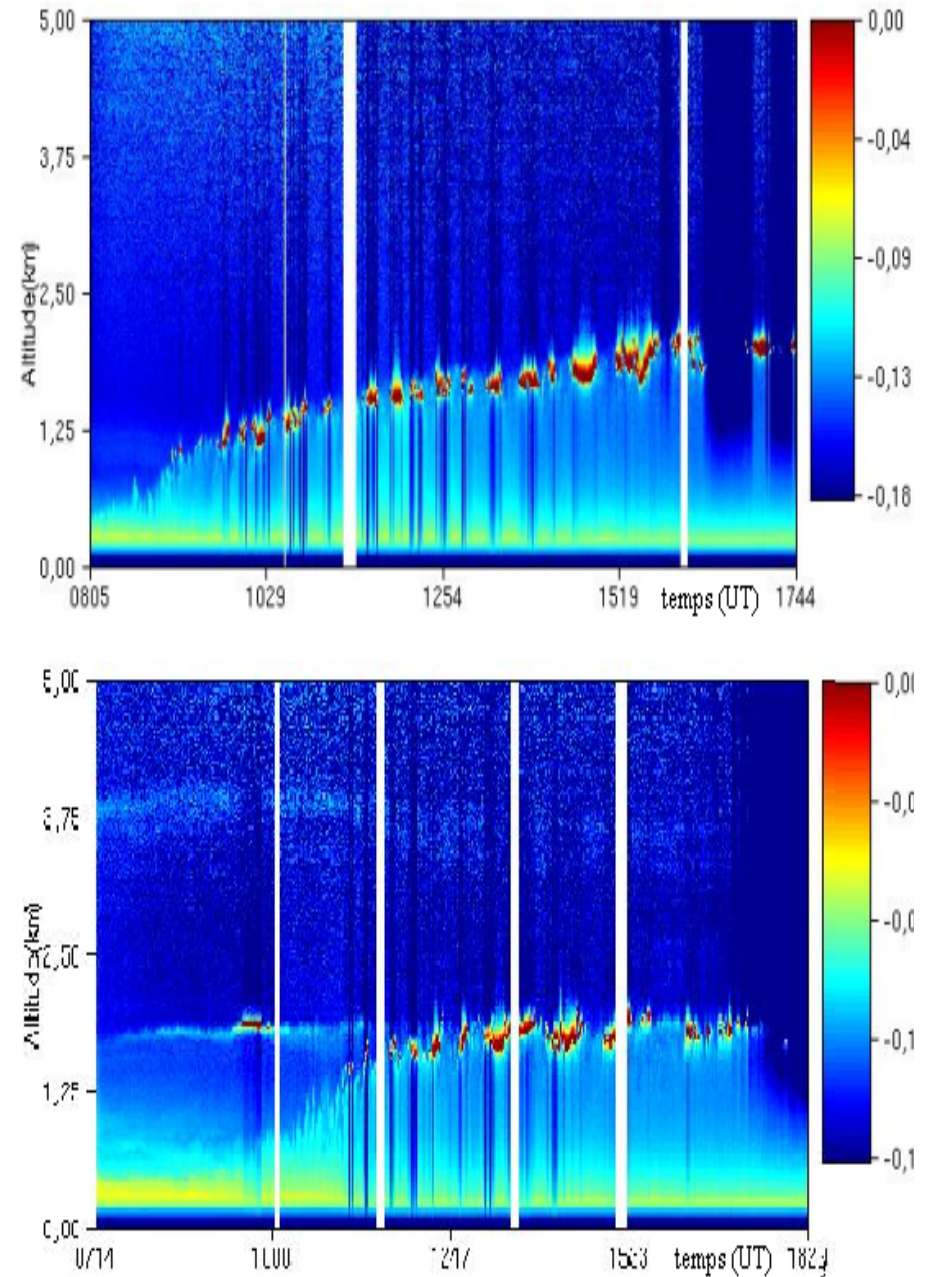


Exemple d'observations de la couche limite en région parisienne

Sondages de Trappes (Paris)

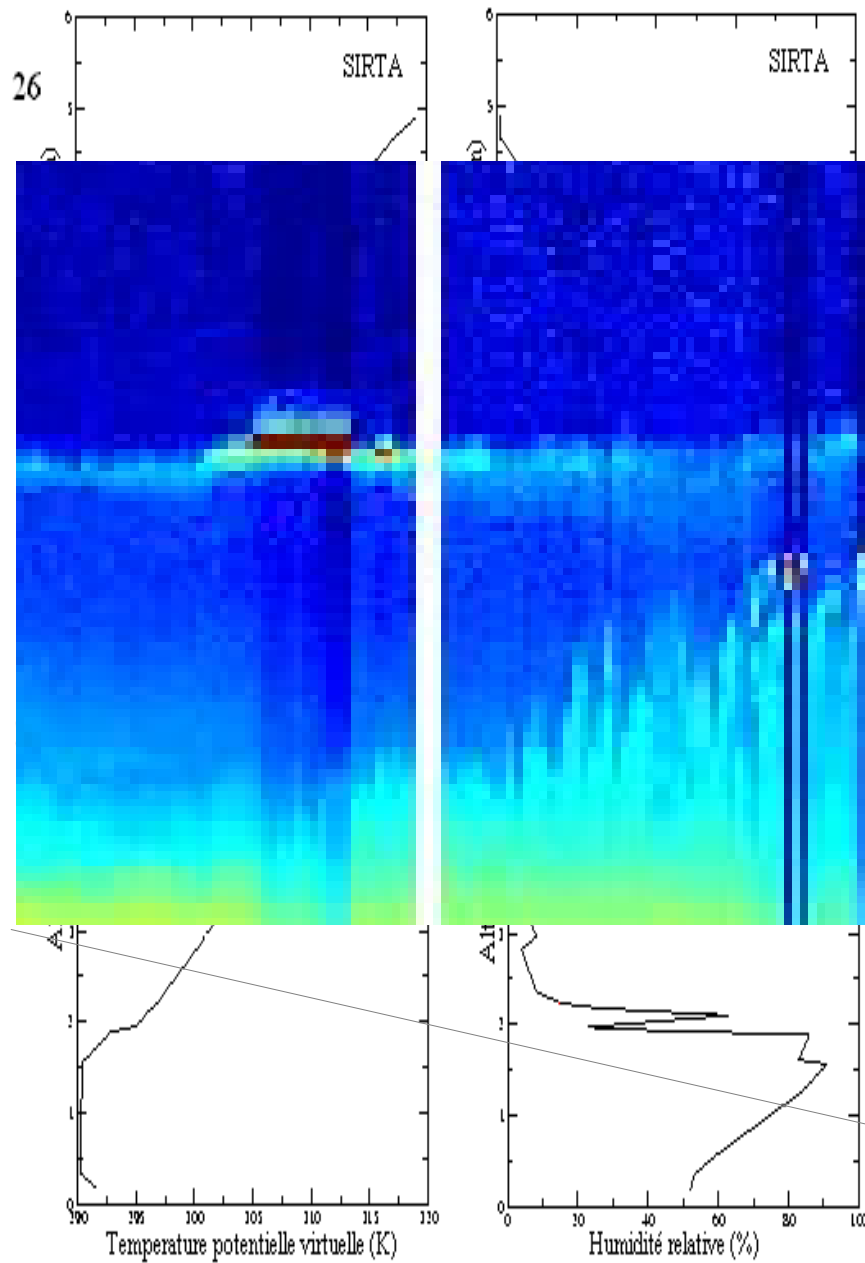


Observations Lidar

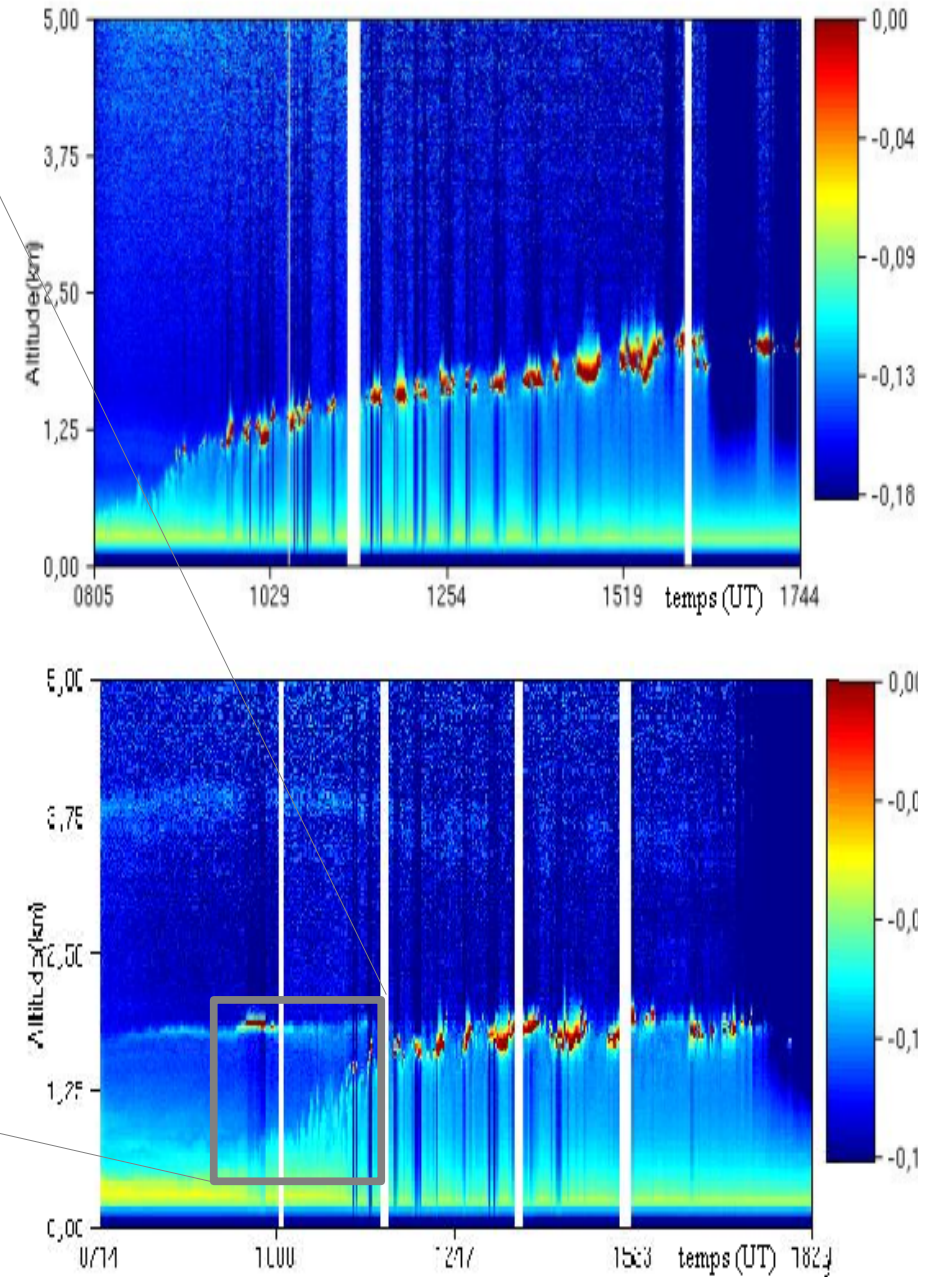


Exemple d'observations de la couche limite en région parisienne

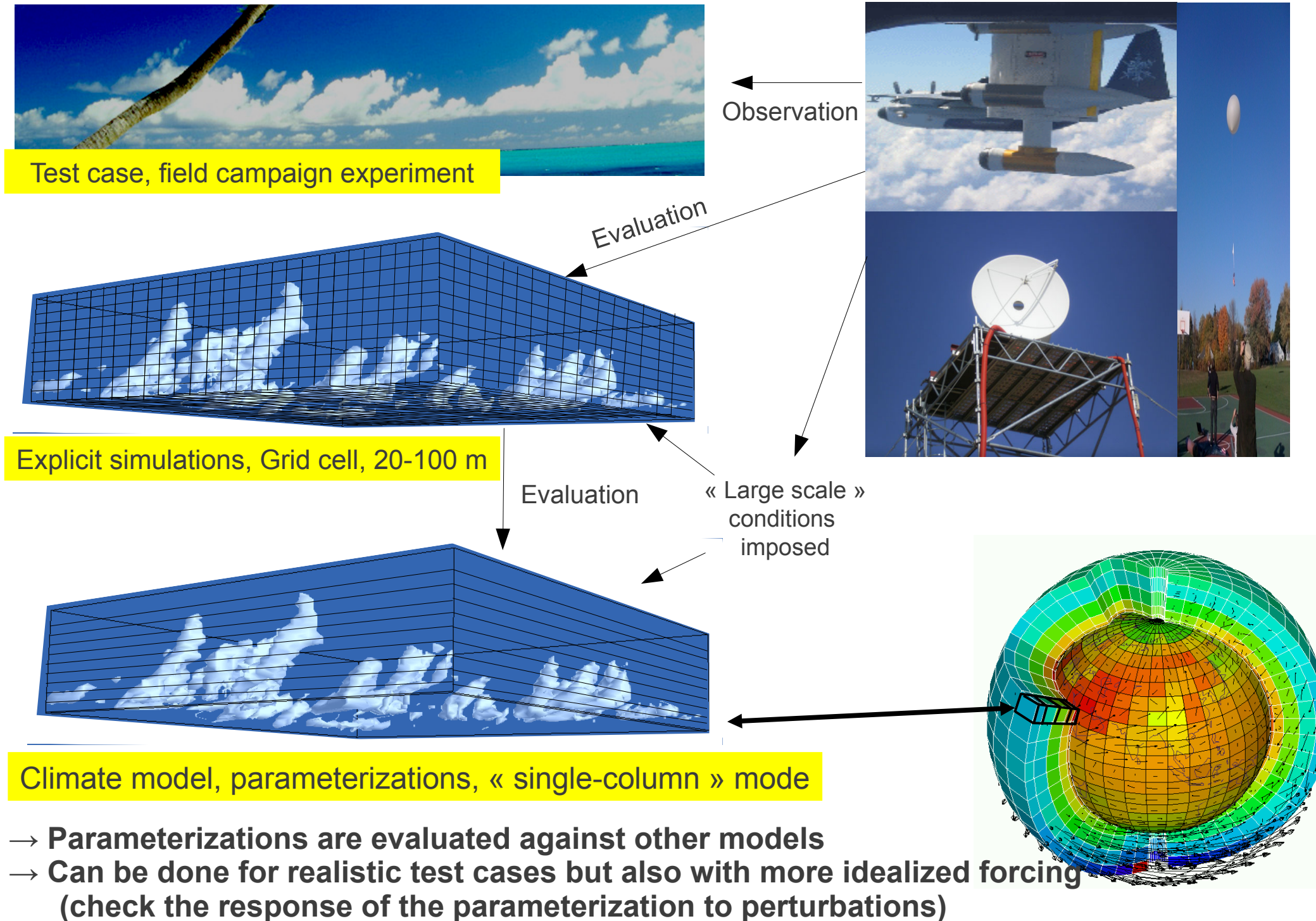
Sondages de Trappes (Paris)



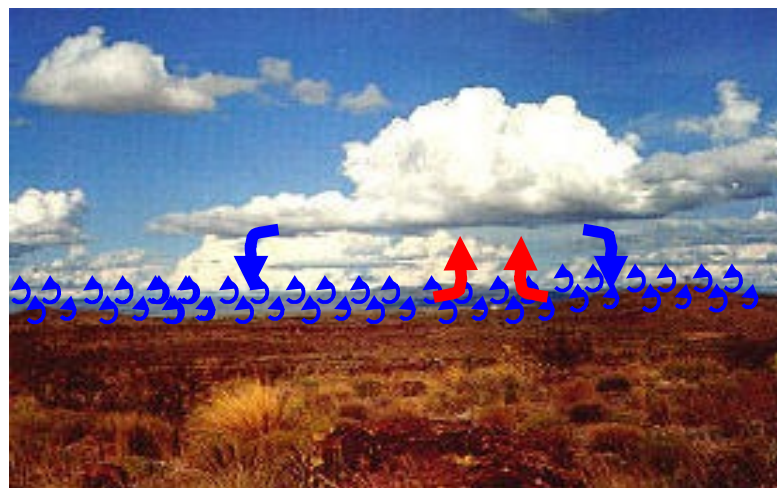
Observations Lidar



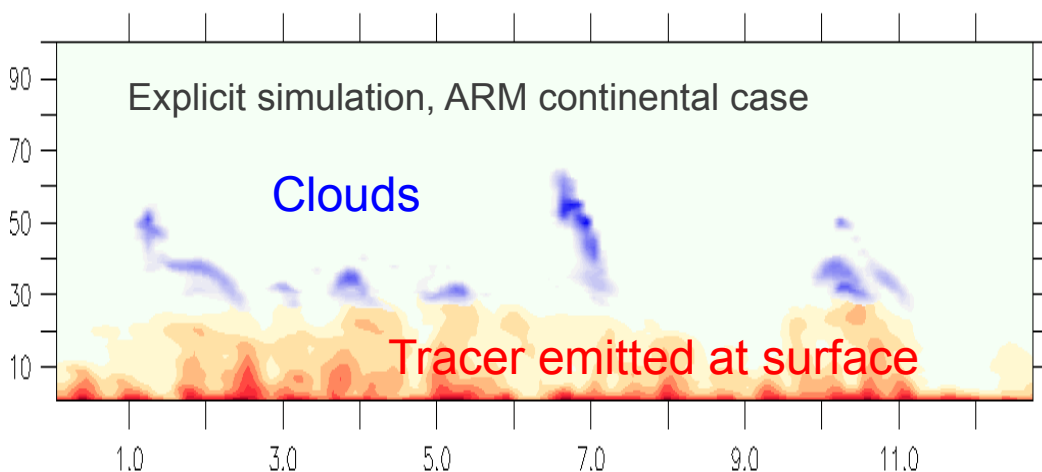
1.2 Cloud process studies and the use of high resolution explicit models



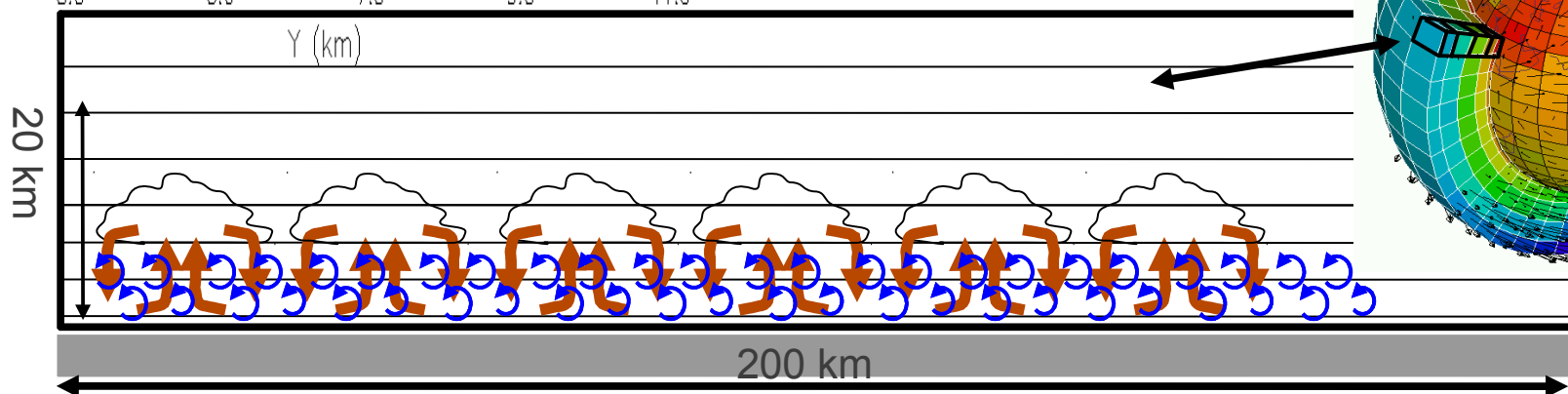
II.1 Thermal plumes and clouds



Meso-NH simulation



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Turbulent diffusion :
for isotropic small scale turbulence

Atmospheric turbulence :
“meso-scale”, organized and anisotropic

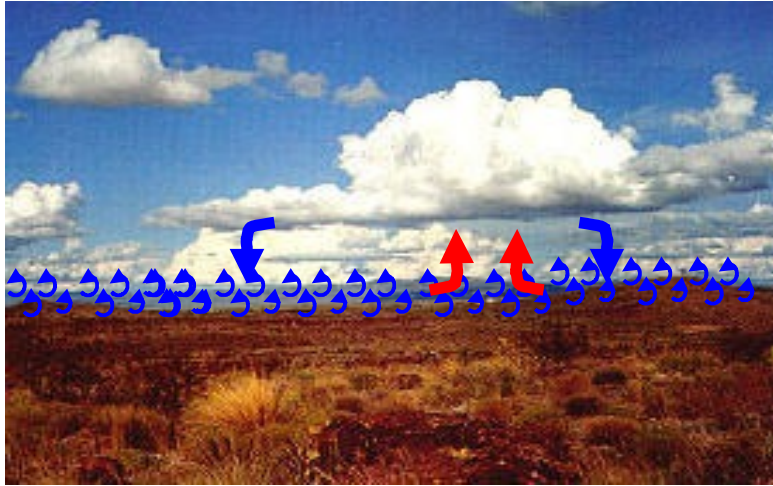
→ « **Thermal plume model** »

Each atmospheric column is divided in 2 :

- plume of air rising from the surface
- air subsiding around the plume

A « mean plume » is represented, at the top of which a « mean cumulus » can appear.

II.1 Thermal plumes and clouds



Turbulent diffusion :

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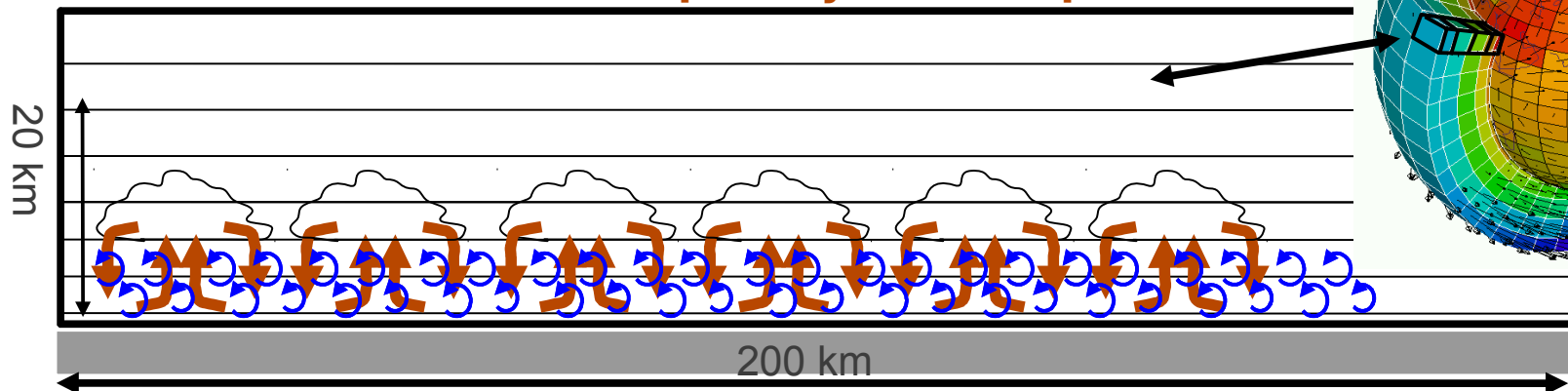
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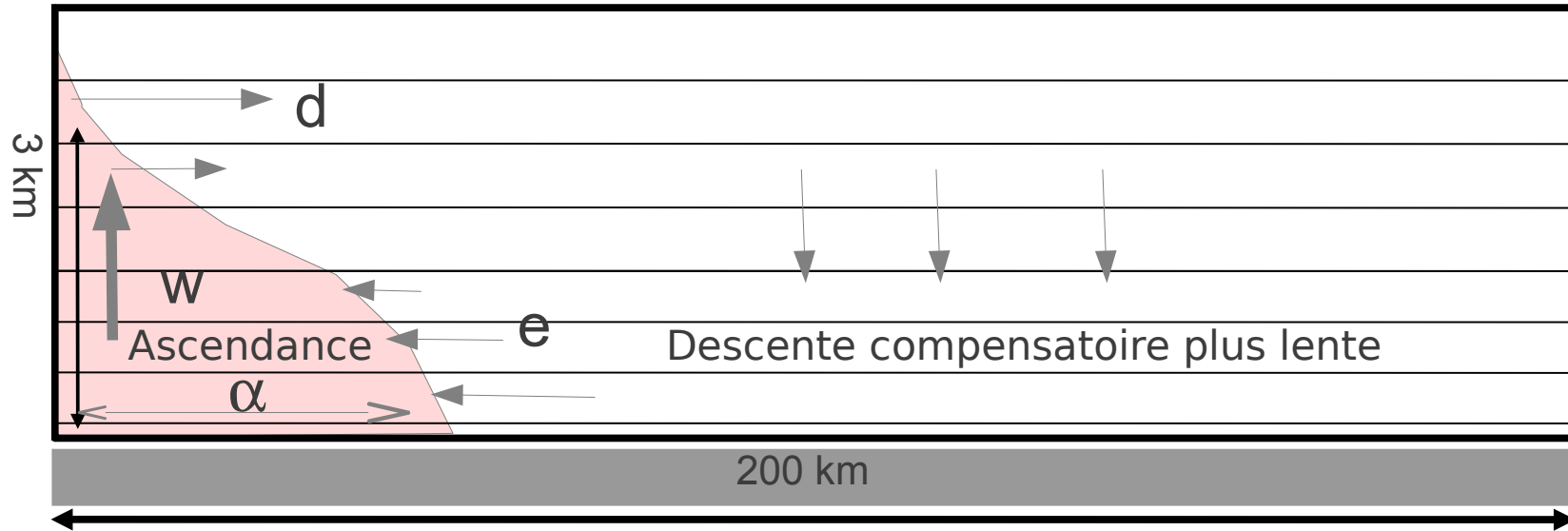
$$S_q = -\frac{1}{\rho} \frac{\partial}{\partial z} \overline{\rho w' q'} = \underbrace{-\frac{1}{\rho} \frac{\partial}{\partial z} \rho K_z \frac{\partial q}{\partial z}}_{\text{Turbulent diffusion}} + \underbrace{-\frac{1}{\rho} \frac{\partial}{\partial z} [\rho \alpha w (q_a - q)]}_{\text{Transport by thermal plumes}}$$

Turbulent diffusion

Transport by thermal plumes



III. Développement récent : le modèle du thermique



Variables internes de la paramétrisation :

w : vitesse moyenne des panaches ascendants
 α : fraction de la surface couverte par les ascendances
 e : taux d'entrée latérale d'air dans le panache (entraînement)
 d : sorties d'air depuis le panache (détrainement)
 q_a : concentration du composant q dans l'ascendance

Terme source pour les équations explicites

$$S_q = -\frac{1}{\rho} \frac{\partial}{\partial z} \overline{\rho w' q'} = \underbrace{\frac{1}{\rho} \frac{\partial}{\partial z} \rho K_z \frac{\partial q}{\partial z}}_{\text{Diffusion turbulente}} + \underbrace{-\frac{1}{\rho} \frac{\partial}{\partial z} [\rho \alpha w (q_a - q)]}_{\text{Transport par le modèle de panache}}$$

Diffusion turbulente **Transport par le modèle de panache**

4 Paramètres libres :

$$a_1 = \frac{2}{3}, \beta_1 = 0.9, b = 0.002, c = 0.012 m^{-1}, d = 0.5$$

Conservation de la masse :

$$\frac{\partial f}{\partial z} = e - d \quad \text{avec } f = \alpha \rho w$$

Conservation de la masse du composant q

$$\frac{\partial f q_a}{\partial z} = e q - d q_a$$

Equation du mouvement

$$\frac{\partial f w}{\partial z} = -d w + \alpha \rho B$$

B étant la poussée d'Archimède

$$B = g \frac{\theta_{va} - \theta_v}{\theta_v}$$

$$e = f \max(0, \frac{\beta_1}{1 + \beta_1} (a_1 \frac{B}{w^2} - b))$$

$$d = f \max(0, -\frac{a_1 \beta_1}{1 + \beta_1} \frac{B}{w^2} + c (\frac{q_a - q}{q_a})^d)$$

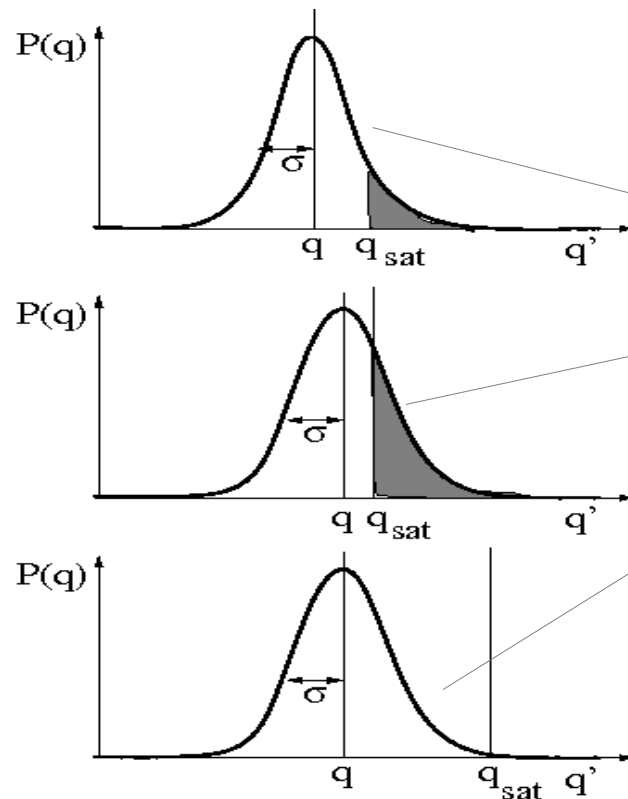
Etc ...

I. Modèles de circulation générale

Représentation des nuages

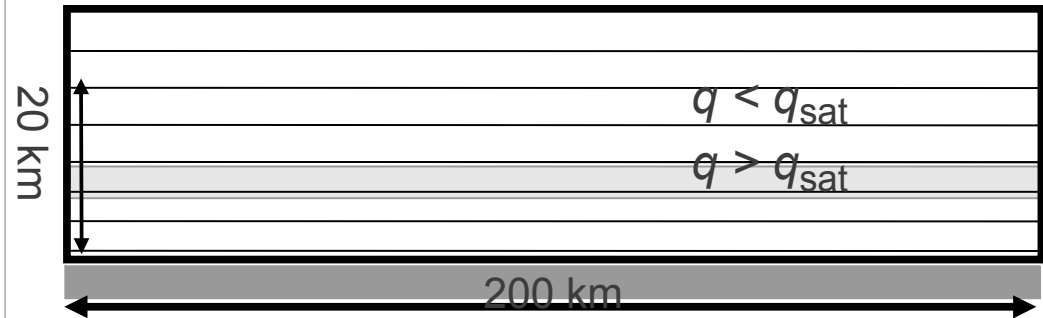
q : concentration en vapeur d'eau
 q_{sat} : concentration maximum à saturation
Si $q > q_{\text{sat}}$:
→ la vapeur d'eau condense = nuage

On connaît q et q_{sat} à l'échelle de la maille
→ Fraction de la maille couverte de nuages ?



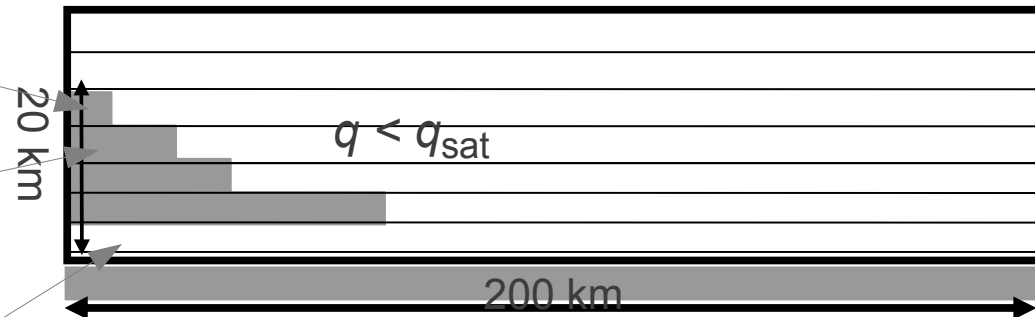
Modèle « tout ou rien » :

Si $q > q_{\text{sat}}$ maille nuageuse, sinon ciel clair.



Modèle « statistique » :

On suppose une distribution statistique de q' dans la maille autour de q



Intervient dans **Q**

→ condensation

→ prise en compte des nuages dans le code radiatif

Paramétrisation simple : gaussienne $\sigma / q = 20\%$

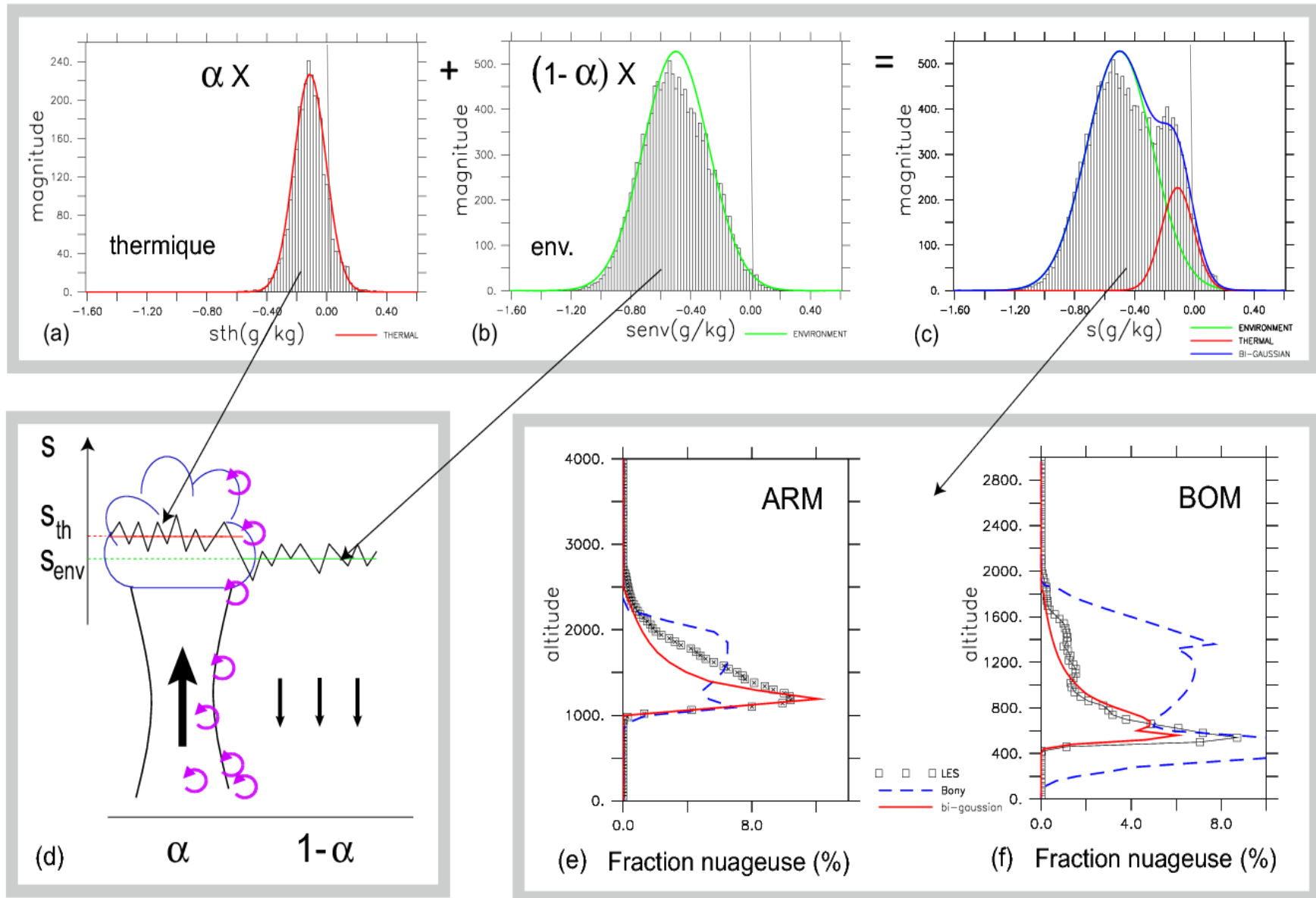
III. Développement récent : le modèle du thermique

Nouvelle paramétrisation de nuages couplée aux thermiques :

Utilisation d'une PDF bi-gaussienne pour la distribution d'eau totale sous nuageuse

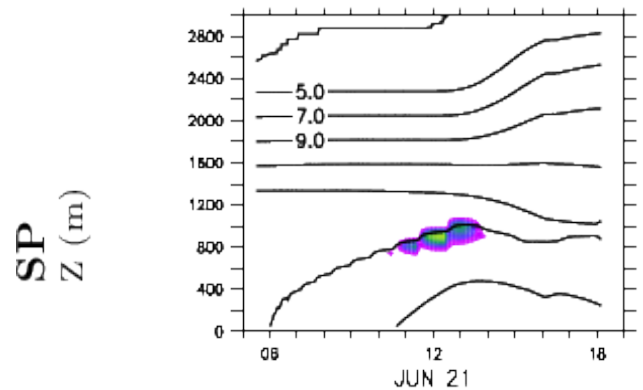
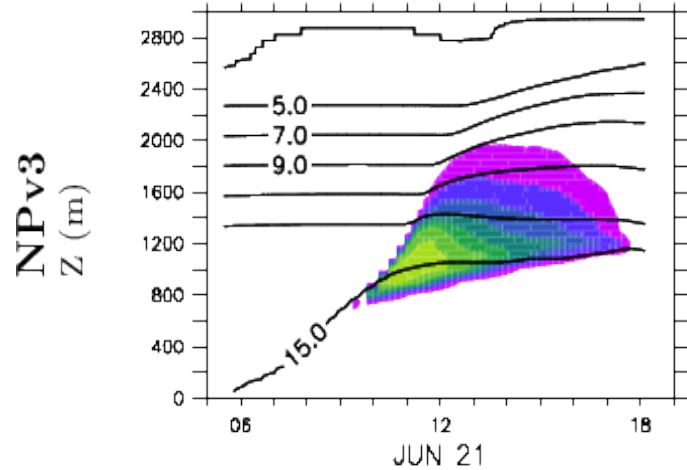
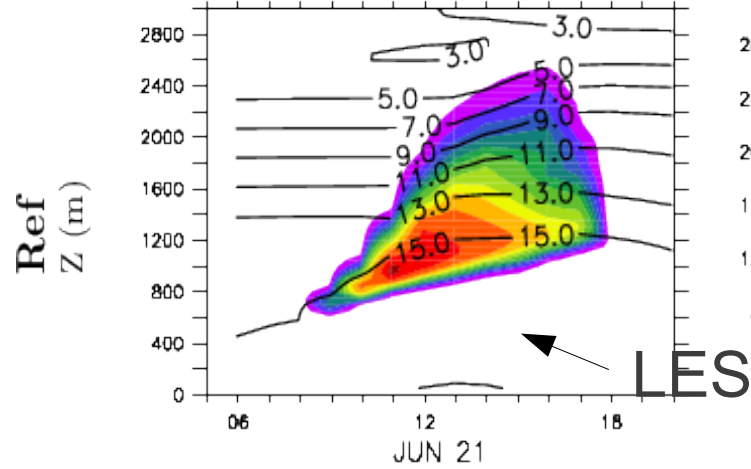
Une gaussienne pour les panaches thermiques et une pour l'environnement

Comparaison des distributions prédites par ce schéma avec les distributions des LES

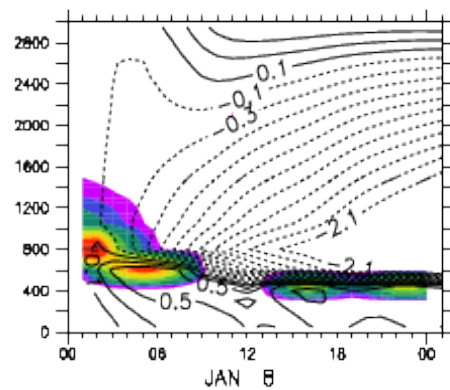
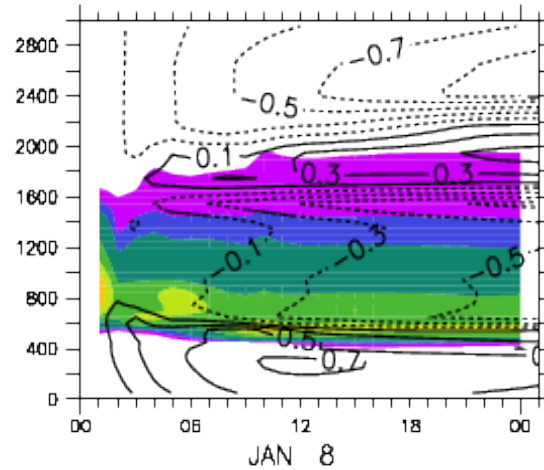
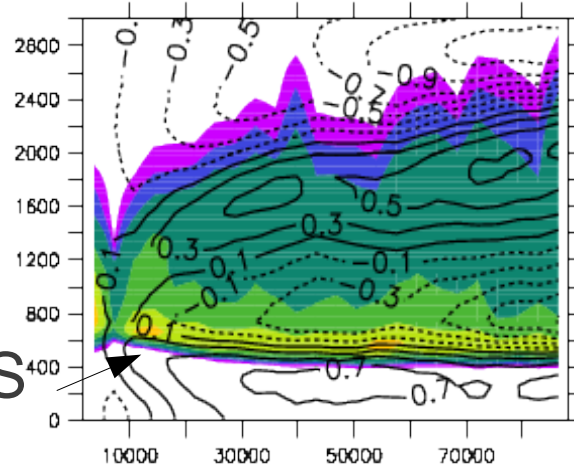


1D test cases

Eurocs Cumulus

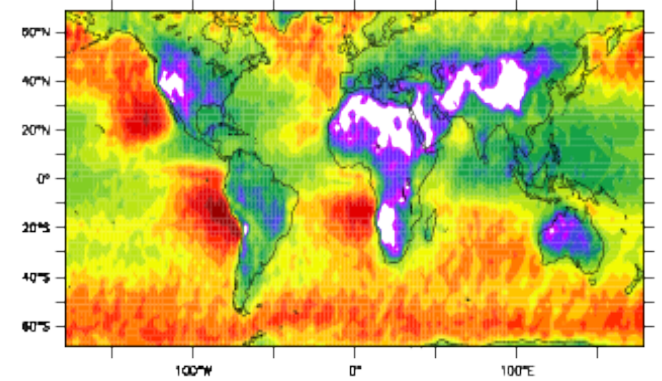


Rico

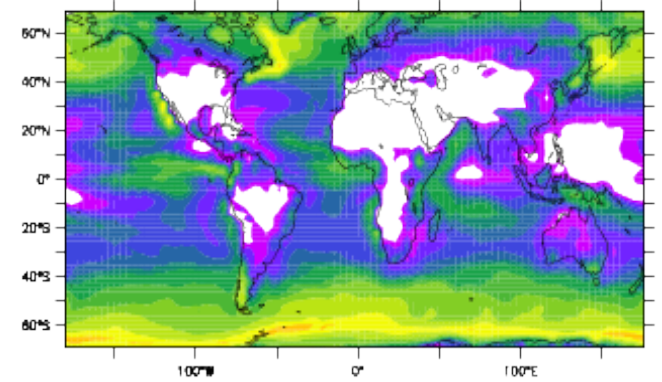
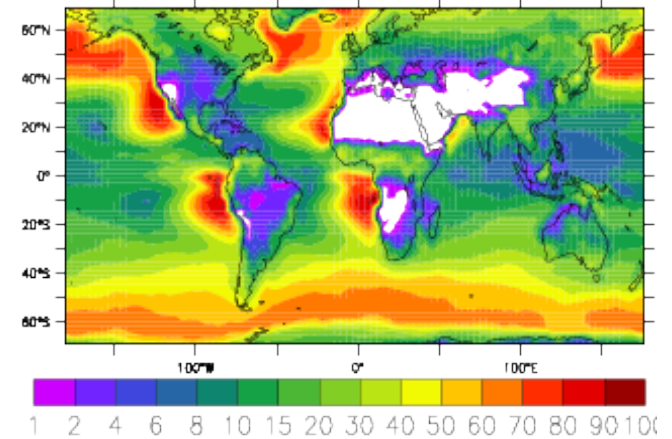


3D simulations

Calipso

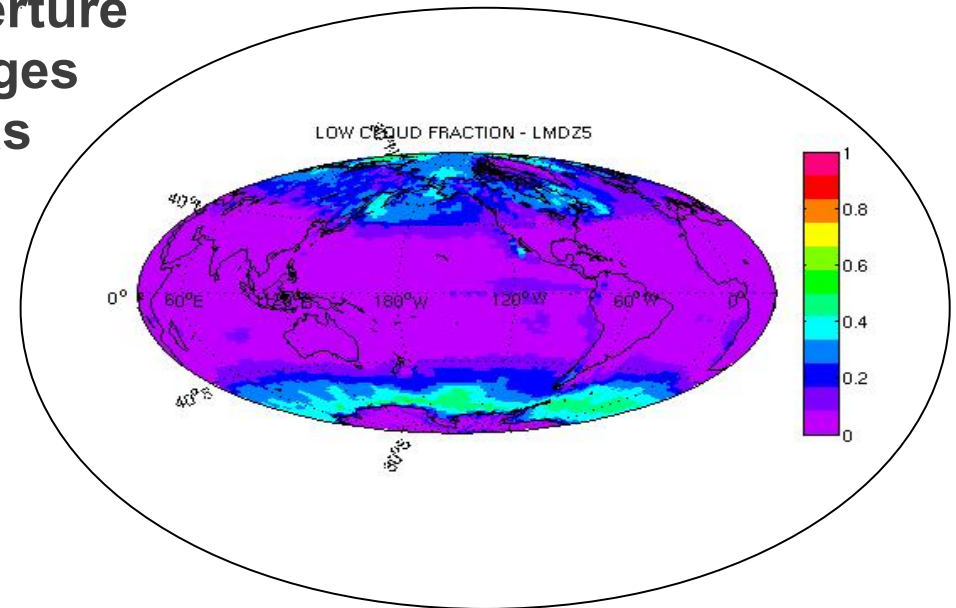
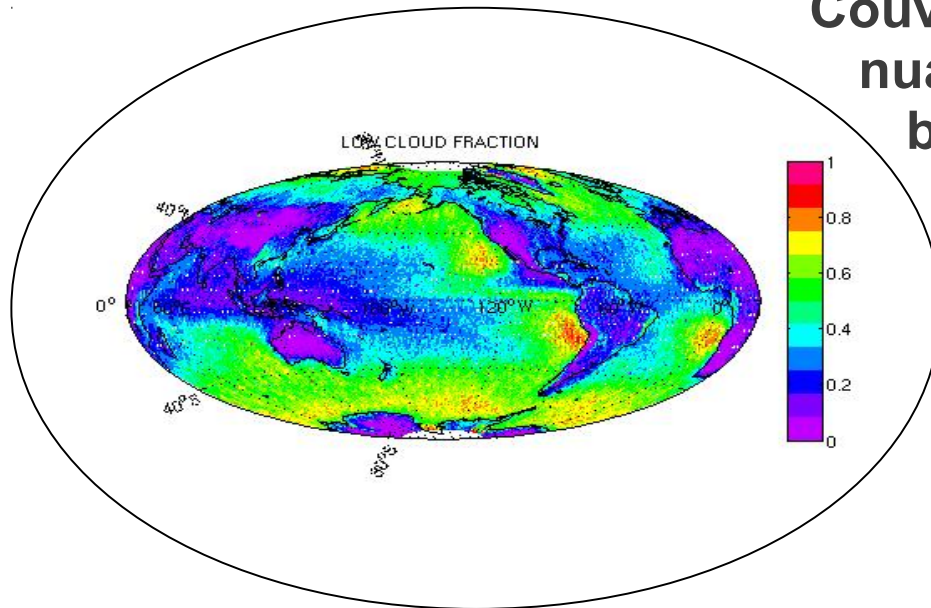


Using COSP simulator
To compare model and satellite

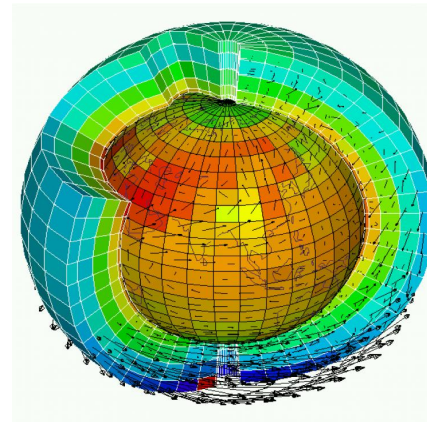




Couverture nuages bas



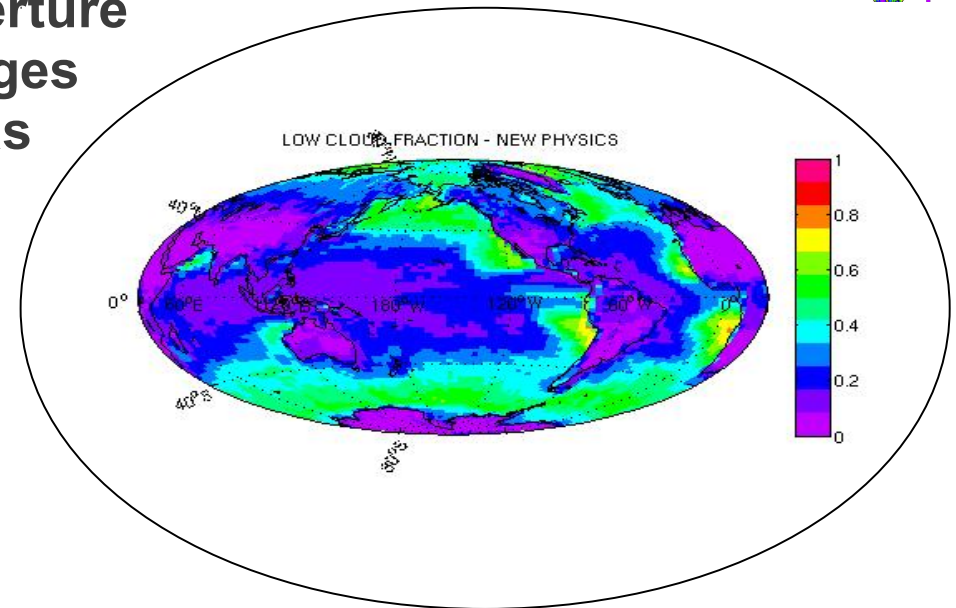
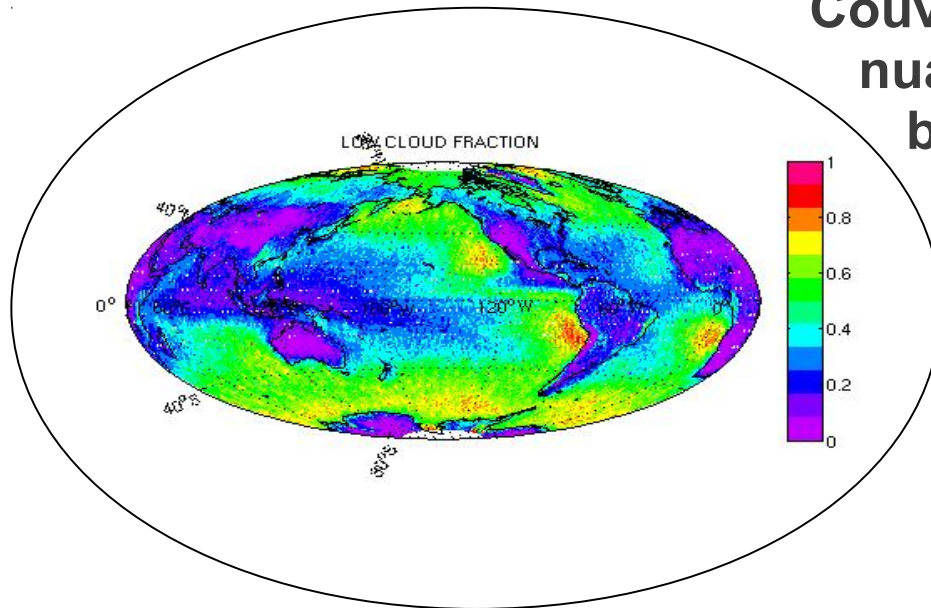
**Observations
calipso
grille LMDZ**



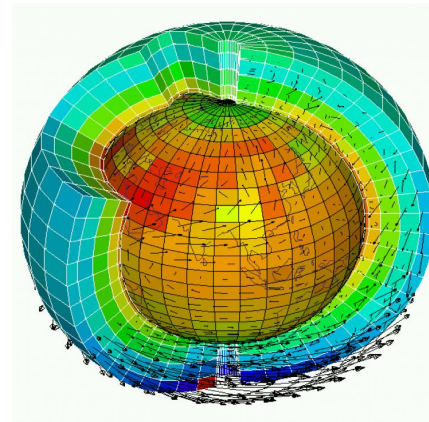
**LMDZ « physique
standard »
+ simulateur
calipso**



Couverture nuages bas



**Observations
calipso
grille LMDZ**



**LMDZ « nouvelle
physique »
+ simulateur
calipso**

III Illustration with the LMDZ climate model : Climate change projections

